University of Southern Queensland Faculty of Health, Engineering and Sciences

School of Engineering

Peak flow estimation considering future climate scenarios for flood resilience modelling of the Warrego Highway at Laidley Creek

A dissertation submitted by

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Abstract

The Warrego Highway at Laidley Creek is susceptible to inundation following periods of high catchment rainfall, causing the closure of a high priority connection within the state-controlled road network. Increased surface temperatures attributed to climate change are projected to increase the intensity of extreme rainfall events. It is therefore anticipated the extent of flooding from Laidley Creek will increase and events will become more frequent.

This study investigated the impact of increased high intensity rainfall on peak discharge characteristics using a semi-distributed node-reach runoff routing model developed in the RORB software package. The model parameters were calibrated against eight historic rainfall-runoff events encompassing a variety of peak magnitudes and event durations. The LIMB-BOM 2020 design rainfall envelope was factored to account for the projected median rises in surface temperature across two emissions scenarios at three projection horizons. Current and projected design event discharges were estimated using the Monte Carlo stochastic simulation method and compared to the estimates obtained by independent techniques including flood frequency analysis and regionalised regression methods.

The simulated discharges closely corresponded to the independent estimates for the 2% and 1% AEP events, indicating the model was accurately calibrated for low frequency, high flow events. The model projected rises in the median design discharge of 51% and 49% for the 2% and 1% AEP events respectively by 2090. However, uncertainty in the results increased for the higher frequency events, with less alignment between the independent estimates and the simulated flows.

The hydrologic simulations completed in this project form the first component of a proposed site investigation into the current and future flood resilience of the Warrego Highway at Laidley Creek. An accompanying hydraulic investigation of the existing bridge crossing section is recommended to complete the investigation.

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(This is a 2-unit research project in the Bachelor of Engineering Honours Program)

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iv

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Table of Contents

Abs	stract .		ii
Lin	nitation	ns of Use	iii
Cer	tificate	e of Dissertation	iv
Acł	nowle	dgements	v
Lis	of Tal	bles	xi
Lis	of Fig	ures	xii
Abl	oreviat	ions	xvi
1.	Intro	duction	1
1.1	Ov	verview	1
1.2	Ba	ckground	1
1.3	Sc	ope for research	2
1.4	Re	search objectives	2
1.5	Pro	oject benefits	3
2.	Litera	ature Review	4
2.1	Int	roduction	4
2.2	La	idley Creek catchment	4
	2.2.1	Catchment geography and flow characteristics	4
	2.2.2	Catchment land uses	7
	2.2.3	Warrego Highway crossing of Laidley Creek	8
	2.2.4	Flood history of Laidley Creek at the Warrego Highway	9
2.3	Re	gional hydrologic modelling	12
	2.3.1	Overview	12
	2.3.2	Introduction to Australian Rainfall and Runoff 2019 (ARR 2019)	13
	2.3.3	Design floods and rainfall events	14
	2.3.4	Event terminology	15
	2.3.5	Flood event criteria for the state-controlled road network	16
2.4	Inp	out data for hydrologic modelling	18

	2.4.1 Observed stream discharge data	18
	2.4.2 Observed rainfall data	20
	2.4.3 Design rainfall data	22
	2.4.3.1 Revised LIMB 2020 IFDs	23
2.5	Conversion of rainfall to runoff in regional catchments	24
	2.5.1 Catchment rainfall and streamflow relationships	24
	2.5.2 Design rainfall Areal Reduction Factor	25
	2.5.3 Rainfall temporal patterns	26
	2.5.4 Catchment losses	28
2.6	Hydrologic modelling methods	31
	2.6.1 At-site flood frequency analysis	31
	2.6.1.1 Annual maximum series	32
	2.6.1.2 Log-Pearson III distribution	34
	2.6.2 Runoff routing	35
	2.6.2.1 Node-reach model features	38
	2.6.2.2 Subcatchment area delineation	38
	2.6.2.3 Non-linear storage routing	38
	2.6.3 Continuous catchment simulation models	39
	2.6.4 Regression methods of analysis	40
	2.6.4.1 Parametric regression	40
	2.6.4.2 Quantile regression	40
2.7	Future climate considerations	41
	2.7.1 Overview of climate change	41
	2.7.2 Future climate impacts on hydrologic components	44
	2.7.2.1 Event intensity and frequency	44
	2.7.2.2 Temporal patterns	44
	2.7.3 Direct climate impacts within the Lockyer Valley region	44
	2.7.4 Design recommendations for future flood estimation	46

	2.7.4.1	DTMR climate change risk assessment framework	48
2.8	Summa	ıry	48
3.	Methodo	logy	50
3.1	Introdu	action	50
3.2	Design	methodology overview	50
	3.2.1 Pre	liminary methodology scope	50
	3.2.2 Bes	st practice modelling considerations	51
3.3	Data a	equisition and preparation	52
	3.3.1 Cat	chment geographic spatial data	52
	3.3.2 Gau	ıged streamflow data	56
	3.3.2.1	Rating curve and control evaluation for 143229A (Warrego Highway)	58
	3.3.2.2	Rating curve and control evaluation for 143209B (Mulgowie)	60
	3.3.3 Obs	served rainfall data	62
	3.3.3.1	Daily observed rainfall totals	62
	3.3.2.2	Hourly observed rainfall	64
	3.3.4 Des	ign rainfall event parameters	64
	3.3.4.1	LIMB-BOM IFD enveloped dataset	64
	3.3.4.2	Projected future climate IFD datasets	64
	3.3.4.3	ARR Datahub parameters	65
	3.3.5 Fina	disation of the design methodology scope	65
3.4	Semi-c	listributed runoff routing model development	65
	3.4.1 Cat	chment model generation from spatial data	65
	3.4.2 Ger	neration of calibration storm files from previous observed events	68
	3.4.3 Cal	ibration of the RORB model parameters	71
	3.4.3.1	Calibrated events with closely replicated peak flows	74
	3.4.3.2	Problematic or uncertain calibration events	77
	3.4.4 Cui	rent and future design flow simulations	79
	2111	Ensamble simulation method	80

	3.4	.4.2 MC simulation method	81
3.5	Ва	yesian Flood Frequency Analysis	82
	3.5.1	Initial plotting position from annual maximum series	82
	3.5.2	Fitted distribution functions	83
	3.5.3	Bayesian parameter estimation	85
	3.5.4	FFA design discharge results	86
3.6	Re	gression methods of peak flow estimation	86
	3.6.1	RFFE method	86
	3.6.2	P&W method	88
3.7	Re	search methodology evaluation	90
3.8	Su	mmary	90
4.	Resu	lts	91
4.1	De	esign discharge simulation for current climate conditions	91
4.2	De	esign discharge simulation for future climate scenarios	93
4.3	Se	nsitivity analysis of the model parameters	95
5.	Discu	ssion	97
5.1	Inc	dependent sanity validation of the results	97
	5.1.1	Comparison of current design discharge simulation	97
	5.1.2	Comparison of design discharges for future climate scenarios	98
5.2	Li	mitations of the results	99
	5.2.1	Data limitations	99
	5.2	.1.1 Uncertainty in the climate projections	100
	5.2	.1.2 Inadequate catchment coverage from rainfall stations	100
	5.5	.1.3 Duration of the gauged streamflow record	101
	5.2	.1.4 Simplification of catchment spatial features	102
	5.2.2	Modelling limitations	102
	5.2	.2.1 Calibration uncertainty	102
	5.2	2.2 Simulation uncertainty	104

	5.5.2.3 Validation uncertainty	105
	5.2.2.4 Limited consideration of time variable components	105
5.3	Potential model improvements	105
6.	Conclusion	107
6.1	Summary of results	107
6.2	Recommended future work	108
Refe	erences	109
App	endix A: Project specification and work plan (topic later refined)	116
App	endix B: Rainfall station datasets coverage	121
App	endix C: Design rainfall IFDs at Laidley Creek catchment centroid	123
App	endix D: ARR Datahub files - Laidley Creek catchment	131
App	endix E: RORB catchment file	137
App	endix F: Calibration storm event files	143
App	endix G: Design discharge simulation sets	167
App	endix H: Sensitivity analysis of model parameters using 2020 outputs	178
Apn	endix I: Flood frequency analysis – RMC outputs	182

List of Tables

List of Figures

Figure 2.1: Upper reaches of the Laidley Creek Catchment surrounded by Little Liverpool Range (foreground) and Mistake Mountains (background) (Starkey 2023)
Figure 2.2: Laidley Creek crossing at Thornton (Lockyer Valley Regional Council 2024b) 5
Figure 2.3: Thornton crossing overtopped during May 2022 flood event (Lockyer Valley Regional Council 2022b)
Figure 2.4: Sandy Creek floodplain at Forest Hill (Lockyer Valley Regional Council 2024b)
Figure 2.5: Upper reaches of the Laidley Creek Catchment surrounded by Little Liverpool Range (foreground) and Mistake Mountains (background) (Starkey 2023)
Figure 2.6 : Laidley Creek catchment outlet at 143229A gauging station adjacent to Warrego Highway (Department of Resources 2024a)
Figure 2.7: Thornton crossing overtopped during May 2022 flood event (Lockyer Valley Regional Council 2022b)
Figure 2.8: Sandy Creek floodplain at Forest Hill (Lockyer Valley Regional Council 2024b)
Figure 2.9: Laidley Creek flows overtopping Warrego Highway on 13/05/2022 (Department of Transport and Main Roads 2022a)
Figure 2.10: Widespread inundation of the Laidley Creek floodplain adjacent to the Warrego Highway during the January 2011 flood event (Lacey 2011)
Figure 2.11: ARR 2019 Preferred Terminology (Bates et al. 2019)
Figure 2.12: Illustration of freeboard between roadway surface and cross drainage flows (Institute of Public Works Engineering Australasia 2017)
Figure 2.13: Recommended design event AEPs for cross drainage road flood immunity (Weeks, Babister & Retallick 2023)
Figure 2.14: Example rating curve generated from gauged data (Bates et al. 2019)
Figure 2.15 : Source of incremental error through rating curve extrapolation at high discharges (Bates et al. 2019)
Figure 2.16: Design Rainfall Intensity IFD chart (example) (Bureau of Meteorology 2016b)
Figure 2.17: Short duration ARF equation for all locations in Australia (Babister et al. 2016)
Figure 2.18: Long duration ARF parametric equation for Laidley Creek (Babister et al. 2016)

Figure 2.19: Example event cumulative burst distributions (Jordan, Seed & Nathan 2019)
Figure 2.20: IL/CL model application to hyetograph (Ball, Weinmann & Boyd 2019)29
Figure 2.21: Median initial loss distribution for Australia (Ball, Weinmann & Boyd 2019)
Figure 2.22: Continuous loss distribution in Australia (Ball, Weinmann & Boyd 2019)30
Figure 2.23: Burst and storm initial losses (Ball, Weinmann & Boyd 2019)31
Figure 2.24: Time series plot of discharge data (Holland, Herschy & Archer 1998)
Figure 2.25: Storage and travel impacts between inflow and outflow hydrographs (Ball, Weinmann & Boyd 2019)
Figure 2.26: Node-reach runoff routing model processes (Ball, Weinmann & Boyd 2019)
Figure 2.27: Projected CO2 emissions (mass, left) and CO2 atmospheric concentration (ppm, right) for RCPs until 2100 (Van Vuuren et al. 2011)
Figure 2.28: Project global temperature rises associated with the RCP climate scenarios relative to the 1961-1990 baseline (Wasko et al. 2024)
Figure 2.29: Annual mean temperature anomaly in Queensland between 1910 and 2023 based on 30 year average from 1961 to 1990 (Bureau of Meteorology 2024a)
Figure 3.1: Poor alignment of Laidley Creek using the rasterised threshold flow accumulation method
Figure 3.2 : Laidley Creek catchment boundary delineated from ridgelines identified in the 25 metre contour overlay
Figure 3.3: Laidley Creek catchment with centroid location, major stream network and gauging stations shown
Figure 3.4: Site details for the 143229A and 143209B streamgauges extracted from the WMIP (Department of Regional Development, Manufacturing & Water 2024)
Figure 3.5: Monthly maximum streamflow record for 143229A gauge [generated from data produced by (Department of Regional Development, Manufacturing & Water 2024)]57
Figure 3.6 : 143229A Warrego Highway hydraulic control and surrounding stream section (Department of Regional Development, Manufacturing & Water 2024)
Figure 3.7: Well defined site rating curve from individual gaugings for 143229A station (Department of Regional Development, Manufacturing & Water 2024)
Figure 3.8: 143209B Mulgowie hydraulic control and surrounding stream section (Department of Regional Development, Manufacturing & Water 2024).

Figure 3.9: Site rating curve with significantly varied gaugings for 143209B station (Department of Regional Development, Manufacturing & Water 2024)
Figure 3.10: Full (green) and partial (yellow) rainfall station event record status
Figure 3.11: Suitable displacement between actual subcatchment centroids (black) and snapped centroid nodes to adjacent stream network (orange)
Figure 3.12: Partitioned subcatchments and adjusted stream reaches for RORB catchment file development
Figure 3.13: May 2022 event rainfall (mm) isohyetal distribution from active station datasets (stars)
Figure 3.14: Calculated hydrograph (red) for February 2022 event using regionalised kc = 22.48
Figure 3.15: Calculated hydrograph (red) for May 2022 event using regionalised kc = 22.48
Figure 3.16: Lateral separation between calculated and observed hydrographs for January 2013 event
Figure 3.17: Calibrated hydrograph for November 2008 event
Figure 3.18: Calibrated hydrograph for January 2011 event
Figure 3.19: Calibrated hydrograph for January 2013 event
Figure 3.20: Calibrated hydrograph for March 2017 event
Figure 3.21: Calibrated hydrograph for February 2022 event
Figure 3.22: Calibrated hydrograph for January 2024 event
Figure 3.23: Calibrated hydrograph for May 1996 event
Figure 3.24: Calibrated hydrograph for May 2022 event
Figure 3.25: 5% AEP, 48 hour Ensemble simulation event hydrographs
Figure 3.26: Sample set of runoff frequency curves generated from MC simulation
Figure 3.27: 143229A annual maximum discharge series 1991 – 2023
Figure 3.28: 143229A Laidley Creek at Warrego Highway AM series input and plotting position graph generated in RMC-BestFit
Figure 3.29: LP3, Gamma and Weibull distributions fitted to AM series with outlier flows identified
Figure 3.30: Posterior mode and confidence intervals of the Bayesian parametric estimation method for the LP3 distribution
Figure 3.31. REFE input tool for 1/3229A catchment characteristics 87

Figure 4.1: Distribution of MC simulated peak discharges for current climate scenario92
Figure 4.2 : Distribution of sample 1% AEP event discharges for current IFDs and RCP 8.5 scenarios
Figure 4.3: Median 1% AEP peak discharges for RCP 4.5 and 8.5 scenarios to 209095
Figure 5.1: Reduction of central catchment daily rainfall station coverage from May 1996 (left) to February 2022 (right)
Figure 5.2: Level of calibration accuracy for various flow magnitudes during January 2011 event

Abbreviations

AADT Average Annual Daily Traffic

AATOS Average Annual Time of Submergence

AIC Akaike Information Criteria

AEP Annual Exceedance Probability

AHD Australian Height Datum

AM Annual Maximum

ARF Areal Reduction Factor

ARI Average Recurrence Interval

ARR Australian Rainfall and Runoff: A Guide to Flood Estimation

BIC Bayesian Information Criteria

BOM Bureau of Meteorology

CCNHRA Climate Change and Natural Hazards Risk Assessment

CL Continuing Loss

DDD Darling Downs District

DEM Digital Elevation Model

DRDMW Department of Regional Development, Manufacturing and Water

DTMR Department of Transport and Main Roads

EY Exceedances per Year

FFA Flood Frequency Analysis

FIP Flood Information Portal

GCM Global Circulation Model

GEV Generalised Extreme Value

GP Generalised Pareto

IDW Inverse Distance Weighted

IFD Intensity-Frequency-Distribution

IL Initial Loss

IPCC Intergovernmental Panel on Climate Change

LiDAR Light Detection and Ranging

LIMB Lockyer-Ipswich-Moreton Bay

LGA Local Government Area

LVRC Lockyer Valley Regional Council

LP3 Log Pearson Type III

MC Monte Carlo

MLE Maximum Likelihood Estimation

PILF Potentially Influential Low Flow

POT Peak-over-Threshold

P&W Palmen and Weeks

QFCD Queensland Future Climate Dashboard

QWMN Queensland Water Modelling Network

RCP Representative Concentration Pathway

RCM Regional Climate Model

RFFE Regional Flood Frequency Estimation

RMC RMC-BestFit

RSME Root Mean Square Error

SRTM Shuttle Radar Topography Mission

WMIP Water Monitoring Information Portal

1. Introduction

1.1 Overview

This research project intends to assess the impact of intensifying extreme rainfall events caused by climate change on critical transport infrastructure. For this project, the Warrego Highway crossing of Laidley Creek was selected as the case study location of interest. The site was chosen because of its historical record of flooding and its significance within the Southeast Queensland highway network controlled by the Department of Transport and Main Roads.

Defined within this introduction are the design objectives which guided the scope of literature reviewed to develop a thorough understanding of the multi-disciplinary topic, including hydrologic modelling technical recommendations and requirements. The objectives also guided the formation of a suitable research methodology, which considers modelling practices utilised by industry to ensure quality in the results obtained.

Throughout this report, the inputs sourced and outputs generated in response to each design objective are verified against supporting literature and independent methods of analysis to ensure the finalised results are as accurate as possible. The uncertainty and limitations associated with climate scenarios projections, as well as different modelling methods, are emphasised throughout this report. The results are presented in a format suitable for a future hydraulic investigation of the case study location as an extension to this research project.

1.2 Background

The Warrego Highway between Ipswich and Toowoomba supports approximately 26500 vehicle movements per day (Department of Transport and Main Roads 2023) as the primary connection for communities and freight between metropolitan Southeast Queensland and regional Southern Queensland.

The section of the highway corridor surrounding Laidley Creek is prone to inundation because of its low-lying elevation. Previous high rainfall events within the Laidley Creek catchment have resulted in major flooding, causing closures which prohibit traffic flow along a vital component of the state-controlled highway network.

The intensity of extreme rainfall is projected to increase in the coming decades due to raised surface temperatures attributed to climate change (Bureau of Meteorology & CSIRO 2022; Department of Transport and Main Roads 2024f).

Infrastructure upgrades to the Warrego Highway crossing of Laidley Creek have been explored to enhance flood immunity during peak events, in conjunction with improving motorist safety at a notoriously dangerous intersection. The lengthy period associated with the planning and design of upgrades to significant public infrastructure means accurate estimates indicative of future climate conditions are essential to achieving desired improvements to the site. Accurate planning and design is vital to prevent intensified damage to infrastructure from an evolving climate.

1.3 Scope for research

Broad projections of future climate conditions are widely published in literature and continually revised as modelling capabilities are improved. However, impacts of climate change on rainfall and the subsequent effects on catchment hydrology particularly at a regionalised scale are emerging, but remain less publicised.

It is hypothesised that inundation of the Warrego Highway crossing site will be exacerbated under future climate scenarios. However, no case study has yet to incorporate these projections as a component of a site specific hydrologic investigation of the Laidley Creek catchment. A clear gap in current knowledge exists with an ideal scope that underpins this project.

Hence, this research project aims to quantify the effects of increased extreme rainfall intensity on peak event discharge within the Laidley Creek catchment under multiple climate projection scenarios. It is intended this would be achieved through the development of a hydrologic model calibrated to the topography of the Laidley Creek catchment and current climate conditions. Revision of the model inputs representative of anticipated future rainfall intensities would yield updated design discharge estimates to be compared with current design discharges.

1.4 Research objectives

A series of sequential research objectives were established to advance the case study, as outlined below and expanded upon within the corresponding components of this report.

Objective 1: Background review of literature

Completion of a comprehensive review of literature to identify fundamental parameters, applications and assumptions covering the multi-disciplinary aspects associated with this project. Topics to be reviewed include catchment hydrology, climate science, transport engineering and infrastructure management. Academic, technical and government sources of reference material should be considered.

Objective 2: Documentation of the design research methodology

Documentation of a definitive research methodology that thoroughly describes the formation of a suitable hydrologic model, referencing the findings and assumptions introduced in the literature review.

Objective 3: Data acquisition and processing in preparation for technical modelling

Acquisition of high quality spatial and hydrologic data sets to produce a subcatchment model and storm event files compatible with a designated software platform.

Objective 4: Calibration of the design model

Generation of design discharge estimates from model specific parameters calibrated to historic peak flow events and verified against independent computational methods.

Objective 5: Simulation of future design discharges

Simulation of revised design discharge estimates representative of the impacts of climate change on hydrologic processes at a localised scale, such that the differences in hydrologic processes between current and future climate conditions are quantified.

Objective 6: Communication of the model outcomes

Preparation of this report communicating the development and findings of the research, including uncertainties in the outputs. An evaluation of the adopted research approach, specifically its strengths and limitations and a recommended course of action is incorporated.

1.5 Project benefits

Several beneficial outcomes are anticipated from undertaking this research project. Foremost is an enhanced understanding of current and future catchment hydrologic processes and interactions within the Laidley Creek catchment.

The design discharge estimates obtained through this research project are a crucial input for a two-dimensional hydraulic modelling study of the bridge crossing site. While outside the scope of this project, detailed hydraulic investigations are necessary to assess the current and future levels of flood immunity achieved by a bridge structure. The accuracy of these models is reliant on the accuracy of the results determined in the preceding hydrologic modelling (this report).

In addition to the intended use of the results of this research, an improved understanding of catchment hydrology could assist the development and revision of floodplain management strategies to reduce the impact of intensifying rainfall events on local communities.

2. Literature Review

2.1 Introduction

This chapter provides a synthesised background of literature reviewed across various topics relevant towards the development of various components of the project methodology. This review begins by introducing the characteristics of the Laidley Creek catchment and the Warrego Highway crossing at Laidley Creek. The application of Australian Rainfall and Runoff 2019 for modelling with observed and design events is explored. The application of event criteria for infrastructure design and management within the state controlled road network is examined. Then, the fundamental processes of regional hydrology that contribute to the generation of runoff from rainfall are reviewed. Subsequently, catchment flood modelling approaches representative of regional hydrologic processes are introduced, including at-site flood frequency analysis, runoff routing, continuous simulation and regression methods of analysis. The advantages and limitations of each approach are examined within the context of this research project. Future global climate change scenarios are introduced and the impacts on catchment hydrologic processes are contextualised to Australian conditions. Relevant published methodologies that detailing how future climate projections are incorporated within the hydrologic modelling for the design and management of state controlled road infrastructure are reviewed. Finally, this chapter concludes by outlining the evident knowledge gap that exists in current literature forming the basis of this dissertation.

2.2 Laidley Creek catchment

The Laidley Creek catchment is located within the Lockyer Valley Regional Council (LVRC) local government area (LGA) and the Darling Downs District (DDD) of the state-controlled road network managed by the Queensland Department of Transport and Main Roads (DTMR). The following sections outline the geographical characteristics of the catchment and the flood history at the catchment outlet adjacent to the Warrego Highway.

2.2.1 Catchment geography and overview of flow characteristics

The Laidley Creek catchment encompasses an area of 462 km² consisting of all the tributaries of Laidley Creek upstream of the Department of Regional Development, Manufacturing and Water (DRDMW) operated streamgauge 143229A titled 'Laidley Creek at Warrego Highway' (Department of Regional Development, Manufacturing & Water 2024). Nestled between the ridgelines of the Little Liverpool Range and the Mistake Mountains, Laidley Creek originates in steep, densely-vegetated terrain where elevations exceed 1000 - 1100 m AHD (Department of Resources 2024b), as depicted in Figure 2.1.



Figure 2.1: Upper reaches of the Laidley Creek Catchment surrounded by Little Liverpool Range (foreground) and Mistake Mountains (background) (Starkey 2023)

Overland flows generated from the steep hillslopes are collected within a distinct valley containing Laidley Creek. Low volume, channelised flows pass through the rural localities of Townson, Thornton and Mulgowie, where minor bridge crossings are located, as exemplified in Figure 2.2. When flow volumes increase substantially during peak flood events, these crossings are overtopped, as observed in Figure 2.3. A second DRDMW operated streamgauge (143209B) is located at Mulgowie which captures an upstream subcatchment area of 167 km² (Department of Regional Development, Manufacturing & Water 2024).



Figure 2.2: Laidley Creek crossing at Thornton (Lockyer Valley Regional Council 2024a)



Figure 2.3: Thornton crossing overtopped during May 2022 flood event (Lockyer Valley Regional Council 2022b)

Laidley Creek continues north past the outer edges of the townships of Laidley and Forest Hill, where the landscape is dominated by expansive, low-lying floodplains and rolling hills, as characterised by Figure 2.4 below. The 6.9 GL capacity off stream reservoir named Lake Dyer is situated in this portion of the catchment (SEQWater 2024). Flowing parallel to Laidley Creek is its major tributary, Sandy Creek, which originates from the slopes of Mount Berryman and flows north until its confluence with Laidley Creek downstream of Forest Hill. Laidley Creek continues north until the 143229A gauge located at the Warrego Highway crossing. Flows through the streamgauge outlet continue 4.5 kilometres downstream until the confluence of Laidley Creek and Lockyer Creek at Glenore Grove. Lockyer Creek then flows into the Brisbane River at Lowood, just downstream from the Wivenhoe Dam spillway.



Figure 2.4: Sandy Creek floodplain at Forest Hill (Lockyer Valley Regional Council 2024b)

2.2.2 Catchment land uses

The dominant land uses within the catchment are conservation, rural agricultural and rural residential zones according to the LVRC Flood Information Portal (FIP) (Lockyer Valley Regional Council 2022a), as shown in Figure 2.5. The predominantly rural nature of the catchment means the proportion of impervious surface area is negligible in the context of a regional hydrology analysis. The high proportion of natural, pervious surfaces and narrow road corridors depicted in Figures 2.1 – 2.4 support this concept.

The FIP provides detailed flood information, including flow depth, velocity, level and hazard for a range of flood events at an individual property level for the entire LVRC LGA. The information serves as a useful validation source for the modelling within this dissertation.

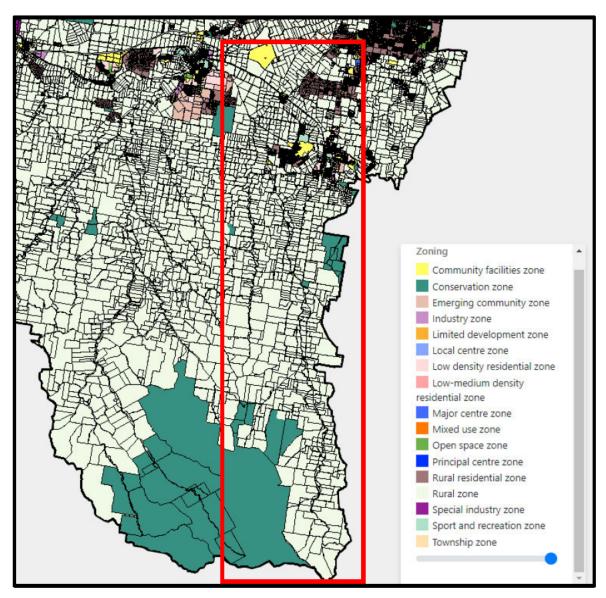


Figure 2.5: LVRC FIP land use zoning within the Laidley Creek catchment - approximate extent annotated (Lockyer Valley Regional Council 2022a).

The agricultural industry has a significant presence in the Lockyer Valley region, which in the 2020-21 financial year generated \$375 million of produce for both domestic supply and international exportation (Australian Bureau of Statistics 2021). Commodities produced in the region include vegetables, representing 86% of the regional agricultural output, as well as hay, grains and nursery flowers; and beef and poultry farming (Australian Bureau of Statistics 2021). These commodities require practices such as irrigated broadacre cropping, dryland cropping and livestock grazing, which are heavily reliant on water resource availability and distribution throughout the catchment.

During periods of drought with insufficient water availability for irrigation; crop propagation, growth and survival is continually threatened until harvest. Conversely, periods of intense rainfall and extensive flooding lead to inundation of the typically low lying, flat cropping fields, causing widespread losses of produce. Both climatic extremes threaten employment opportunities and the viability of businesses in the agricultural industry, which impacts the supply of produce into communities.

2.2.3 Warrego Highway crossing of Laidley Creek

The Warrego Highway between Ipswich and Toowoomba is referred as section 18A of the state-controlled road network managed by DTMR. This section of the network is the primary connection for communities and freight between metropolitan Southeast Queensland and regional Southern Queensland, as well as western and interstate destinations.

The Warrego Highway is classified as a national highway route and forms a critical component of the National Land Transport Network, supporting B-Double and Higher Mass Limits vehicular movements (Department of Transport and Main Roads 2019). The average annual daily traffic (AADT) on this section of the highway is 26534 vehicles per day including 21.5% heavy vehicles (Department of Transport and Main Roads 2023). 18A facilitates the transportation of \$19 billion worth of freight annually (Infrastructure Australia 2023) as a vital connection between regional primary industries and producers; and the Port of Brisbane (Department of Transport and Main Roads 2012). Freight transported along the route includes seasonal harvest produce, livestock, mining plant and equipment, fuels, building supplies and machinery, and general consumer supplies for western communities (Department of Transport and Main Roads 2012). The route also serves "as a strategic military ... link between key military installations in southern Queensland, including Amberley RAAF base, Oakey Army Aviation Centre and Borneo Barracks" (Department of Transport and Main Roads 2012).

The Laidley Creek catchment outlet coincides with the DRDMW gauging station 143229A, which is situated adjacent to the Warrego Highway at the Forest Hill – Fernvale Road intersection (between Plainland and Crowley Vale at an approximate chainage of 48.0 km), see Figure 2.6. Laidley Creek passes under the Warrego Highway at the Jack Martin Bridge.



Figure 2.6: Laidley Creek catchment outlet at 143229A gauging station adjacent to Warrego Highway (Department of Resources 2024a)

2.2.4 Flood history of Laidley Creek at the Warrego Highway

Social accessibility and economic productivity characteristics of the region are dependent on the connectivity of the Warrego Highway between Toowoomba and Ipswich, which is compromised by insufficient flood immunity above stream crossings and low lying floodplains. As identified by DTMR, the area surrounding the Laidley Creek outlet is subject to repeated major flooding which causes closures of the Warrego Highway (Department of Transport and Main Roads 2012). The road corridor between chainages 47.8 km and 48.2 km is prone to flooding, especially at both the eastbound and westbound approaches to the Jack Martin Bridge. As pictured in Figure 2.7, the existing surface vertical depression at chainage 47.9 km (at the

intersection of Forest-Hill Fernvale Road) is the lowest level of the Warrego Highway within the immediate vicinity of Laidley Creek and is most vulnerable to inundation.



Figure 2.7: Warrego Highway westbound approach towards Laidley Creek crossing (Department of Transport and Main Roads 2024b)

During peak flow events, transverse flows from Laidley Creek overtop the road surface and cannot be sufficiently dissipated by the existing minor box culvert drainage. Therefore, traffic is prohibited from passing through the crossing. Flooding was recurrent enough to warrant the installation of a monitoring camera, from which imagery from previous flooding events has featured in mainstream media and DTMR publications, see Figures 2.8 and 2.9 below.



Figure 2.8: Laidley Creek flows extensively overtopping Warrego Highway on 26/02/2022 (Department of Transport and Main Roads 2022b)



Figure 2.9: Laidley Creek flows overtopping Warrego Highway on 13/05/2022 (Department of Transport and Main Roads 2022a)

The extent of inundation and hence the length of roadway closed inherently varies between flood events due to the naturally unique rainfall distribution patterns and catchment conditions of each event. The more widespread floods, exemplified by Figure 2.8, have a longer roadway closure duration causing greater social and economic disruptions compared to less significant events. Aerial imagery captured during the January 2011 flood event, shown in Figure 2.10, demonstrates the extent of floodplain inundation from Laidley Creek (situated in the foreground along the tree line) during a peak event. The extent of flooding is consistent with the event characteristics modelled in the FIP (Lockyer Valley Regional Council 2022a).



Figure 2.10: Widespread inundation of the Laidley Creek floodplain adjacent to the Warrego Highway during the January 2011 flood event (Lacey 2011)

Listed in Table 2.1 are the previous events when the Warrego Highway has been inundated by transverse flows of Laidley Creek. The source that confirms the roadway was overtopped during each event, consisting of archival flood records for historic events, and photographic evidence in more recent times, are referenced.

Table 2.1: Observed Laidley Creek overtopping events of Warrego Highway

Start date	End date	Peak discharge at	Source of record to verify roadway	
		143229A (m ³ /s)	inundation/overtopping	
3/05/1996	6/05/1996	496.6	(Bureau of Meteorology 2010a)	
10/01/2011	12/01/2011	1387.1	(Lacey 2011)	
27/01/2013	29/01/2013	1041.5	(Wordsworth 2013)	
25/02/2022	28/02/2022	1097.7	(Department of Transport and Main Roads	
			2022b)	
12/05/2022	14/05/2022	521.1	(Department of Transport and Main Roads	
			2022a)	
30/01/2024	30/01/2024	366.7	(Transport and Main Roads Queensland	
			2024)	

2.3 Regional hydrologic modelling

2.3.1 Overview

Regional hydrologic modelling is a numeric representation of the distribution of water within the natural environment throughout the water cycle. Modelling simulates complex physical processes including rainfall, evaporation and infiltration losses to statistically estimate hydrologic processes of surface runoff, stream flow and groundwater flow (Singh & Woolhiser 2002). Modelling allows engineers and policymakers to understand catchment hydrology and implement measures to optimise resource distribution, particularly during flood or drought.

A catchment is an area bounded by a natural topographic rise which causes all surface flow to drain to a common outlet, often through at least one channel (Department of Environment & Innovation 2021). Larger regional catchments are comprised of smaller areas known as subcatchments. A rural catchment is characterised as a predominantly naturally discharging basin with a high proportion of pervious surfaces allowing infiltration of rainfall into the ground (Ladson 2014).

A hydrologic model is classified by its period of simulation as either a single event method or a continuous simulation method. Event methods simulate a singular flood event by simplifying the physical processes behind the conversion of rainfall to runoff as a set of numeric parameters (Babister, Retallick & Testoni 2019). Generally, a loss model is utilised to simulate rainfall excess from a singular storm event, while a hydrograph routing model simulates the conversion of rainfall excess to streamflow considering the spatial characteristics of a catchment (Nathan et al. 2019). With this approach, the prevailing catchment boundary conditions are specified by calibrating the model against previously observed data. Continuous simulation methods convert extended duration, continuous climatic datasets into a output streamflow dataset for the corresponding time series, from which the frequency and extent of flooding is extracted by statistical analysis (Nathan et al. 2019).

2.3.2 Introduction to Australian Rainfall and Runoff 2019 (ARR 2019)

Hydrologic modelling in Australia is standardised by design guidelines published by Engineer's Australia in Australian Rainfall and Runoff: A Guide to Flood Estimation (ARR) 2019. ARR 2019 is comprised of nine separate books. The technical guidelines presented in Books 1-5 and 7 (Ball et al. 2019; Ball, Weinmann & Boyd 2019; Bates et al. 2019; Jordan, Seed & Nathan 2019; Nathan et al. 2019; Babister, Retallick & Testoni 2019) document a variety of modelling approaches to estimate regional hydrologic characteristics. The development of each model from first principles concepts is discussed in detail and contextualised with respect to its intended purpose, required data preparation and model performance limitations.

ARR 2019 provides specific technical guidance about the requirements any catchment hydrologic assessment. ARR 2019 defines the parameters used to establish a flood hydrograph model representative of an event within a catchment, which are discussed throughout this review of literature. The modelling approaches presented within ARR 2019 are developed from the following considerations:

- Catchment geography and spatial characteristics.
- Event rainfall depth; in the form of both historic observed rainfall and design inputs.
- Event characteristics regarding the time and spatial distributions of rainfall.
- The processes which influence the conversion of rainfall to runoff.
- Observed streamflow data, if the catchment is gauged.

ARR 2019 also provides guidance for engineers to incorporate climate change projections into design flood estimation techniques. In addition, DTMR has published its own technical guidelines to standardise the adoption of these projections within modelling for a variety of different departmental projects. This guidance is examined in subsequent sections of this review.

2.3.3 Design floods and rainfalls

The objectives of catchment management and engineering design projects usually adopt risk-based failure or exceedance criteria. The hydrology field utilises hypothetical scenarios known as design flood and/or rainfall events with a specified frequency of occurrence to stipulate flood size characteristics (Bates et al. 2019). Because typical event characteristics, which consist of peak discharge, level and volume, are dependent on variable channel properties such as cross section and surface roughness, probability exceedance criteria provides a common mode of analysis for catchment scale modelling. The principles and methodologies presented in ARR 2019 were predominantly formulated from the concept of design events (Bates et al. 2019).

The probability relationship between design flood and rainfall events is not entirely direct. Flood frequency methods of discharge analysis directly estimate flood characteristics when the probability of a certain event magnitude is exceeded (Bates et al. 2019). However, the exceedance probability associated with a design rainfall event does not necessarily correlate to the corresponding exceedance probability for a flood event. In every modelling scenario, the representation of all processes that contribute to the conversion of rainfall to runoff introduces some joint uncertainty. The prevailing causes and effects associated with this conversion are directly related to the event conditions at the time of occurrence and are usually unique in space and time. Consequently, the "true probability of the derived flood characteristic may be ... biased with respect to the true flood magnitude with the same probability as the design rainfall, especially at low (exceedance) probabilities" (Bates et al. 2019). For example, a rainfall event within a saturated catchment may result in a large volume of runoff and cause significant flooding, however the same rainfall event could occur in a dry catchment may and yield minimal runoff. The process of preserving an event exceedance probability from a design rainfall and transforming it to a design flood is known as AEP neutrality (Bates et al. 2019). ARR 2019 stipulates in Book 3, Section 2.3.6.2 that caution should be exercised when utilising transformative methods of probabilistic-exceedance modelling between rainfall and discharge to avoid large margins of error (Ball et al. 2019).

2.3.4 Event terminology

The frequency of design flood and rainfall events are expressed in terms of descriptor classifications. The utilisation of each classification is dependent on the intended application of the model, as well as the frequency of occurrence of the flood or rainfall characteristic, which is, categorised from very frequent to extreme (Bates et al. 2019). Average recurrence interval (ARI), annual exceedance probability (AEP) and average number of exceedances per year (EY) are the frequency descriptors used in ARR 2019. ARI is defined as "a statistical estimate of the average period in years between the occurrence of a flood of a given size or larger" (Ladson 2014) and is expressed in years. Alternatively, AEP is defined as "the likelihood of a flood of given size or larger (occurring) in any year" (Ladson 2014) and is expressed in either percentage or 1 in *X* years. Frequencies higher than 50% AEP are expressed as *Z* EY to avoid confusion around successive seasonal events (Bates et al. 2019). Outlined in Figure 2.11 is the preferred industry terminology from ARR 2019.

Frequency Descriptor	EY	AEP (%)	AEP (1 in x)	ARI
	12			
	6	99.75	1.002	0.17
Very Frequent	4	98.17	1.02	0.25
very i requerit	3	95.02	1.05	0.33
	2	86.47	1.16	0.5
	1	63.21	1.58	1
	0.69	50	2	1.44
Frequent	0.5	39.35	2.54	2
rrequent	0.22	20	5	4.48
	0.2	18.13	5.52	5
	0.11	10	10	9.49
Rare	0.05	5	20	19.5
Rate	0.02	2	50	49.5
	0.01	1	100	99.5
	0.005	0.5	200	199.5
Van Doro	0.002	0.2	500	499.5
Very Rare	0.001	0.1	1000	999.5
	0.0005	0.05	2000	1999.5
	0.0002	0.02	5000	4999.5
Extreme				
			PMP/ PMP Flood	

Figure 2.11: ARR 2019 Preferred Terminology (Bates et al. 2019)

2.3.5 Flood event criteria for the state-controlled road network

Several flood event metrics are used to assess the resilience of road infrastructure against inundation, which also assist the design of infrastructure upgrades and future corridor planning.

The flood immunity of a roadway is one such metric and is defined as the event probability of the flow level "that just reaches the height of the upstream shoulder of the road, or where the road is kerbed, the top of the inlet pit" (Weeks, Babister & Retallick 2023). In these scenarios, the trafficked pavement surface is above the flood level and remains dry. An allowance for model uncertainty, blockage effects and extremely rare flood magnitudes, known as freeboard, is usually prescribed between the design flood level and the structure level of interest (Weeks, Babister & Retallick 2023), as illustrated in Figure 2.12.

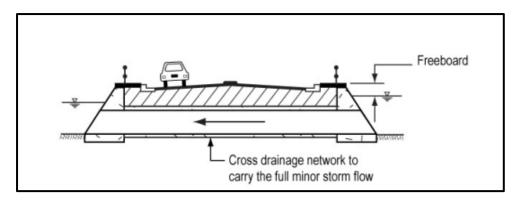


Figure 2.12: Illustration of freeboard between roadway surface and cross drainage flows (Institute of Public Works Engineering Australasia 2017)

The level of flood immunity stipulated for a specific element of road network infrastructure is dependent on many considerations, as detailed by Weeks, Babister & Retallick (2023) in Austroads Guide to Road Design Part 5: Drainage – General and Hydrology Considerations. Such considerations include:

- The strategic function of the road link within the broader network, including funding priorities.
- The availability of suitable, alternate routes during an overtopping flood event.
- Any requirements to maintain emergency access along the road link.
- Any adverse flood impacts on surrounding properties from development within and surrounding the road corridor. For example, a lower level of immunity would facilitate transverse flows across the road surface reducing upstream afflux.

Similar considerations dictate the magnitude of freeboard required for new road designs. Generally, road designs that avoid overtopping flows during a flood event provide 300 mm of freeboard between the water surface level and the pavement subgrade or bridge soffit (Weeks, Babister & Retallick 2023).

The recommended event frequencies for flood immunity against transverse flows for the design of new road infrastructure, including bridge decks, are specified by Weeks, Babister & Retallick (2023) and adopted by DTMR (Department of Transport and Main Roads 2024e). An extract of the recommendations is provided in Figure 2.13 below.

Element	Austroads road classification	AEP
Cross drainage (culverts & bridges)	Controlled Access Highways Includes: Motorways & Freeways (National/State/Territory)	1%
	Arterial Roads Classes 1 & 2 Includes: National/State/Territory Highways, Urban Arterial Roads	2 – 1%
	Arterial Road Class 3 Includes: State/Territory main roads	2%
	Local Roads Classes 4 & 5	10 – 5%
	Urban Collector/Distributor Roads	10 – 2%
	Urban Local Roads	10%

Figure 2.13: Recommended design event AEPs for cross drainage road flood immunity (Weeks, Babister & Retallick 2023)

In this instance, the Warrego Highway is a National Highway and future flood immunity is recommended to be modelled for the 2 and 1% AEP events. To facilitate infrastructure management and future design projects, this research will determine the design discharges at the Laidley Creek outlet for the current and future climate scenarios, with an emphasis on ensuring the 2% and 1% events are accurately simulated.

Another flood event metric used by DTMR is the average annual time of submergence (AATOS), which is defined as the expected average duration per year the roadway is submerged by flood flows of any depth (Weeks, Babister & Retallick 2023). The frequency of overtopping flow events as well as the duration of each event are factors that contribute to AATOS (Weeks, Babister & Retallick 2023). The AATOS of a section of road is an alternate assessment of flood immunity that correlates flood characteristics to disruptions to traffic and incurred damage to infrastructure. Therefore, this metric is calculated by DTMR to evaluate the economic impact of disruptions caused by flooding when developing a business case for infrastructure upgrade projects (Department of Transport and Main Roads 2024e). AATOS is dependent on the catchment response to intense rainfall events which is affected by catchment area, shape, baseflow conditions and soil properties (Weeks, Babister & Retallick 2023).

2.4 Input data for hydrologic modelling

The types of accessible input data for catchment hydrologic modelling are explored in this section. The applicability of the data for different modelling approaches is examined and limitations of use are discussed.

2.4.1 Observed stream discharge data

Flood analysis requires routine collection of real flood data parameters at locations along defined flow paths, including water level stage above a datum (typically the channel base or ground surface), and subsequently discharge, the volume of water passed per unit time. Observed streamflow measurements are taken by instruments known as streamgauges. In Queensland, streamgauge monitoring is operated by the Department of Regional Development, Manufacturing & Water (DRDMW) and observations are recorded via remote electronic water level sensors and collated on the Water Monitoring Information Portal (WMIP). Hydraulic control devices such as a weir or sluice gate regulate flow in open channels to facilitate gauge measurement (Ladson 2014). Rectangular, v-notch and crump weirs are most commonly used in Queensland (Department of Regional Development, Manufacturing & Water 2024) and are typically either artificial concrete masonry structures or formations of natural channel sediment acting as a constriction (Ladson 2014). Empirical equations calibrate the observed flow depth to the equivalent discharge through the control device with known sectional properties (Caroline & Afshar 2014):

Discharge through a rectangular weir:
$$Q = C_d \frac{2}{3} b \sqrt{2gh}^{3/2}$$
 (2.1)

Discharge through a V notch weir:
$$Q = C_d \frac{8}{15} \sqrt{2g} \tan(a/2) h^{5/2}$$
 (2.2)

Discharge through a crump weir:
$$Q = \left(\frac{2}{3}\right)^{3/2} C_d C_v b \sqrt{gh}^{3/2}$$
 (2.3)

Where Q is the weir discharge (m³/s), h is the water depth above weir (m), C_d is the coefficient of discharge, g is the constant acceleration due to gravity (9.81 m²/s), b is the weir width (m), a is the notch weir angle (deg) and C_v is the coefficient of velocity for modular or non-modular flow, depending on the submerged condition of weir (Caroline & Afshar 2014).

In practice, empirical equations facilitate the development of discharge rating relationships. However, modern technological advancements have simplified this process. Measurements of discharge velocity in a channel by current meters are correlated to discharge at various times with associated river stage heights (Ladson 2014). A rating curve is generated by plotting stage-discharge data points and subsequently applying a best-fit curve to estimate discharge at a given

stage for a certain channel cross section (Bates et al. 2019). An example rating curve developed from a set of gauged data is shown in Figure 2.14.

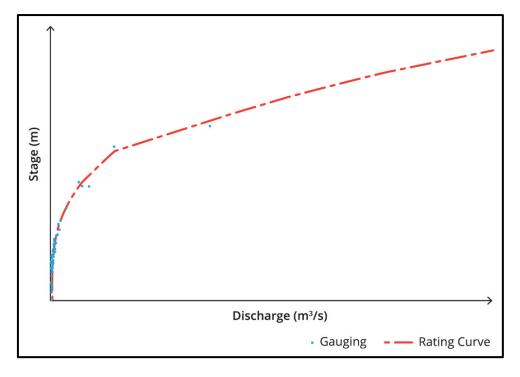


Figure 2.14: Example rating curve generated from gauged data (Bates et al. 2019)

A reliable rating curve is one that is defined thoroughly by a widespread range of flow gaugings but features densely grouped gaugings for similar flow magnitudes. A rating curve is typically well defined for low, frequently occurring flows, provided a constant control is maintained and a sufficient number of gaugings are taken. Vegetation growth, sedimentation and erosion; construction, dredging and damming; as well as downstream backwater effects significantly impact flows through a gauge (Bates et al. 2019) and modify the true site rating. Measurements of higher flows are typically less accurate for two reasons: the infrequency of rare events prohibits sufficient gauge recordings, and water level unsteadiness during intense flooding results in rapid fluctuations of flow slope, leading to variations in discharge from the steady flow condition (Bates et al. 2019). Estimating the discharge of peak flow and hypothetical design events larger than the maximum gauged level at a site requires extrapolation of the rating curve. However, extrapolating the defined curve beyond the range of gauged measurements is subject to error (Ladson 2014). This error is dependent on the nature of the true rating curve compared to the extrapolated estimate, as illustrated in Figure 2.15.

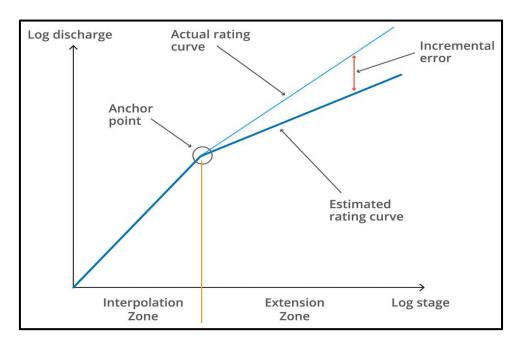


Figure 2.15: Source of incremental error through rating curve extrapolation at high discharges (Bates et al. 2019)

Streamflow data is required to be obtained and managed within various time scales for different hydrologic models, which are discussed in subsequent sections of this report. For example, flood frequency analysis computes flow quantiles from either the annual or daily maximum recorded flow, a continuous catchment simulation utilises daily total runoff volumes, while an hourly or sub-hourly increment is used for event based models which require an enhanced time resolution to generate a short duration hydrograph.

2.4.2 Observed rainfall data

The depth of precipitation that reaches the surface is recorded by an instrument known as a rain gauge. In Australia, observational data for 8000 current rainfall gauging stations is managed by the Bureau of Meteorology (BOM) Climate Data Services branch (Bureau of Meteorology 2010b, 2024b). Rainfall is also recorded at a limited number of DRDMW streamgauges. Historical rainfall data is available for an additional 11000 closed stations, including records that date back to the 1800s (Bureau of Meteorology 2010b).

Two standardised gauging instruments are used at sites across Australia. The oldest instrument is the eight inch rain gauge, which collects 25 mm of rainfall within an internal measuring cylinder while excess precipitation is captured by an outer container (Bureau of Meteorology 2007). Each gauge is placed 300 mm above the surface level to minimise the effects of wind on the accuracy of gauging (Bureau of Meteorology 2007). Observations of the rainfall

received in the preceding 24 hours to 9am are made by local volunteer personnel, who report the rainfall depths to BOM at the end of each month.

More recently, many of these stations have been replaced by contemporary automatic weather stations, which use an eight inch tipping bucket rain gauge. This instrument is capable of providing automatic readings on the depth and rate of rainfall to a precision of 0.2 mm and is self-emptying (Bureau of Meteorology 2007). This approach is more accurate and allows for more frequent rainfall observations to be recorded compared to the traditional manually documented gauging system. These advancements in gauging technology has allowed the introduction of rainfall observations at a minute, sub-hourly and hourly scale, as well as rainfall intensity measurements in the form of a pluviograph (Bureau of Meteorology 2024b).

It is rare that a station with a long operational history has maintained an entirely complete rainfall record. Common explanations provided by the Bureau of Meteorology (2010b) about the gaps in observed data sequences include:

- Closure of a station.
- Upgrade of station capacity.
- Damage sustained to an instrument which requires repair works.
- Absence of observer to take rainfall measurement.
- Automatic weather station failure.

The singular nature of a rainfall station means the observation is a point rainfall reading rather than a spatial rainfall measurement. A network of rainfall gauges is typically required to estimate the actual rainfall volume received in a regional-sized catchment over a specified interval with a sufficient degree of confidence (US National Weather Service n.d.). Various areal rainfall averaging techniques have been developed to compute catchment rainfall volumes, including the numeric and graphical methods used by the US National Weather Service (n.d.):

- Numeric methods; including arithmetic mean and distance-weighted methods.
- Graphical methods; including isohyetal (contour) and Thiessen polygon (areaweighted) methods.

Each method has its own associated advantages and disadvantages. However, the accuracy of catchment rainfall estimation is primarily dependent on the capability of each rain gauge to accurately record each rainfall event as well as ensuring sufficient coverage of gauge stations throughout the catchment (US National Weather Service n.d.). Sharp et al. (2021) noted that

many remote parts of Australia, including rural and mountainous areas in Queensland, had poor rain gauge coverage over an extended period of time due to the inherent geographic isolation and difficulties associated with access to these areas.

2.4.3 Design rainfall data

Design rainfall data is a fundamental component of many flood modelling approaches used in Australia. Observed rainfall recordings were processed to perform rainfall frequency analysis to relate the rainfall characteristics of intensity, the depth of rainfall in a certain time; frequency, the AEP of the rainfall event; and the duration of the rainfall event (Ladson, 2014). Intensity-frequency-duration (IFD) relationships are also referred to as design rainfall bursts between 1 EY and 1% AEP (Jordan, Seed & Nathan 2019) and have a similar purpose as design discharges for probabilistic hydrologic modelling. Design rainfall is expressed in the form $^{\rm Y}{\rm I}_{\rm D}$, where Y is the AEP and D is the duration of the event (Ladson 2014). Statistical exceedance design rainfalls are useful at overcoming the effects of short-term seasonality and climatic variability (Bates et al. 2019). Because numerous factors have an influence on peak flood events, the AEP of a design rainfall may not correspond to the equivalent design discharge.

Location-specific IFD relationships are retrievable from the Design Rainfall Data System (2016), managed by BOM. The database features IFD raster data at a grid cell size of 0.025° latitude and longitude (Bureau of Meteorology 2016a). The system uses cell interpolation to translate point rainfall observations into IFD relationships covering the entire Australian continent, including locations where gauged point data is unavailable or limited (Johnson et al. 2016). Such locations are geographically isolated, often due to high elevation, which prevents access to install and monitor rainfall gauges. These locations typically receive higher rainfalls than low-lying areas which contain the majority of gauge stations. (The et al. 2012), Consequently, most data is sourced from locations at lower elevations, meaning "at high elevations, rainfall is likely to be underestimated when it is spatially interpolated without reference to elevation" (The et al. 2012). The BOM IFD system used thin-plate spline smoothing of a digital elevation model (DEM) to incorporate topography into the design rainfall grid (Johnson et al. 2016), yielding more accurate IFD relationships.

IFD charts provide design rainfall depth in mm or intensity in mm/hr for events ranging from 1 minute to 7 days in duration and are available in either a tabular or chart format (Bureau of Meteorology 2016a; Jordan, Seed & Nathan 2019). Presented as Figure 2.16 is an example BOM IFD rainfall intensity chart, which illustrates that the IFD curves for different frequencies

are approximately parallel when plotted in a log-log scale. Ladson (2014) describes the two relationships that exist between IFD parameters:

- For a certain event duration, the intensity increases as the frequency reduces.
- For a certain event frequency, the intensity decreases as the duration increases.

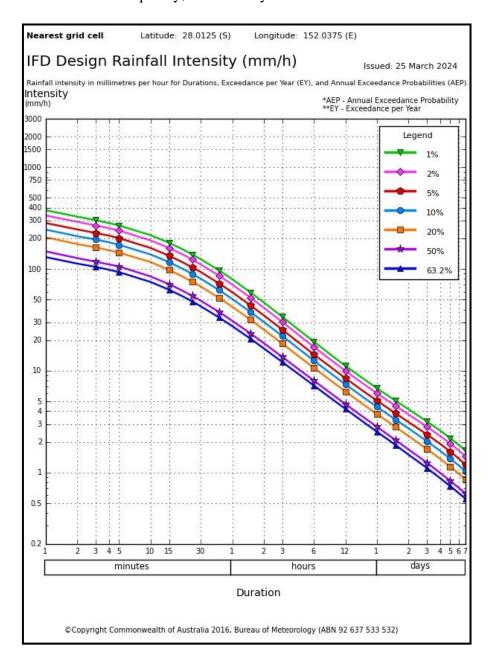


Figure 2.16: Design Rainfall Intensity IFD chart (example) (Bureau of Meteorology 2016b)

2.4.3.1 Revised LIMB 2020 IFDs

A revised set of gridded IFD relationships, known as the Lockyer-Ipswich-Moreton Bay (LIMB) 2020 IFDs, were computed for four LGAs within South-East Queensland including the LVRC after considerable differences were discovered between council rain gaugings and the BOM 2016 IFDs (Babister 2021). The methods used to derive the 2020 LIMB IFDs

prioritised localised rainfall data at a sub-daily time step to achieve "a reduction in local biases across all AEPs, durations and areas, compared to the ARR 2016 IFDs" (Babister 2021). Two alternate recommendations were provided for design flood modelling with the study region: either the revised LIMB 2020 IFDs should be adopted, or for conservative design of high rainfall events, an envelope consisting of the maximum IFDs between the LIMB 2020 and the existing BOM 2016 datasets would also be an acceptable approach (Babister 2021).

2.5 Conversion of rainfall to runoff in regional catchments

The following section introduces the fundamental hydrologic considerations for modelling the conversion process of rainfall to runoff within Australian catchments. Such processes are only applicable to non-urbanised catchments where the vast proportion of surfaces are pervious.

2.5.1 Catchment rainfall and streamflow relationships

Gauged streamflow is the combination of continual groundwater discharge into a stream, known as baseflow, and quickflow. From the commencement of a storm, stream discharge is increased from direct rainfall capture and overland surface drainage. These additions are known as quickflow, often referred to as runoff. Ladson defines quickflow as "the rapid component of catchment runoff that occurs in response to rainfall" (2014). The runoff depth generated is considered as rainfall excess, the total event rainfall depth minus any losses during the conversion process (Ball, Weinmann & Boyd 2019). The subsequent peak discharge yielded from a design rainfall event is computed in two different approaches:

- The development of a time-step discharge hydrograph from a rainfall excess hyetograph at a singular point of interest, or
- Catchment runoff routing through the combination and translation of flood hydrographs, requiring specialised software.

The underlying assumption of the singular discharge hydrograph method is that the runoff originating from the entire catchment area is instantaneously and concurrently contributing to discharge, measured at a point location (Ball, Weinmann & Boyd 2019). The resultant discharge from an excess rainfall intensity during a time-step is defined by equation 2.4 as:

$$Q_i = 0.278 I_{e_i} A (2.4)$$

Where Q_i is the discharge (m³/s) generated during a time-step i, I_{e_i} is the excess rainfall intensity (mm/hr) in the correspond time-step and A is catchment area (km²) (Ladson 2014).

Equation 2.4 is derived from a similar approach to the time of concentration rational method (Department of Transport and Main Roads 2024a), which considers the time of drainage for runoff originating from the most distant location of a small catchment to reach outlet for peak discharge estimation. For large catchment scale analysis with flow regimes that are complex, multifaceted networks of surface and channel discharges, the time of concentration is unable to be easily determined (Ladson 2014). Instead, the critical event duration is determined through a comparison of the generated peak discharges from a range of trial durations for a specified AEP (Ball, Weinmann & Boyd 2019). The duration length with the maximum discharge provides an insight into the catchment response. The factors that predominantly affect the conversion of rainfall to runoff are:

- Inconsistent spatial distributions of rainfall.
- Temporal distribution patterns of rainfall.
- Catchment losses.

These considerations are elaborated upon below.

2.5.2 Design rainfall Areal Reduction Factor

The IFD design rainfalls are specific only to a certain point location and are generally not reflective of the average rainfall intensity across a significantly large catchment. The average areal rainfall accounts for the assumption that "larger catchments are less likely than smaller catchments to experience high intensity storms simultaneously over the whole of the catchment area" (Jordan, Seed & Nathan 2019). The ratio between catchment average and point rainfall is represented by the Areal Reduction Factor (ARF). ARR 2019 stipulates that the ARF shall be applied to reduce upstream rainfall data for catchment discharge estimation (Jordan, Seed & Nathan 2019).

ARF is dependent on AEP, event duration and to a lesser extent, catchment size. Longer duration events also consider regionalised impacts (Jordan, Seed & Nathan 2019) through variable equation coefficients. The ARR Datahub (Babister et al. 2016) generates outputs to determine the ARF of any design scenario considering a maximum event duration of 7 days and a catchment size less than 30,000 km² (Jordan, Seed & Nathan 2019)

For short duration events less than 12 hours in duration, ARF is determined by a singular equation for all locations in Australia. The short duration ARF equation, as reproduced from the ARR Data Hub in Figure 2.17, is dependent on the catchment area, as well as the event AEP and duration in minutes.

$$\begin{split} \text{Short Duration ARF} \\ ARF &= Min \left[1, 1 - 0.287 \left(Area^{0.265} - 0.439 \text{log}_{10}(Duration) \right) . Duration^{-0.36} \right. \\ &+ 2.26 \times 10^{-3} \times Area^{0.226} . Duration^{0.125} \left(0.3 + \text{log}_{10}(AEP) \right) \\ &+ 0.0141 \times Area^{0.213} \times 10^{-0.021} \frac{(Duration - 180)^2}{1440} \left(0.3 + \text{log}_{10}(AEP) \right) \right] \end{split}$$

Figure 2.17: Short duration ARF equation for all locations in Australia (Babister et al. 2016)

Events lasting over 24 hours are classified as long duration and the ARF is governed by regionalised parameters in conjunction with catchment area, design AEP and event duration. Laidley Creek is situated within the 'Semi-arid inland QLD' ARF classification region. The long duration ARF parameters and equation from the ARR Data Hub are shown in Figure 2.18.

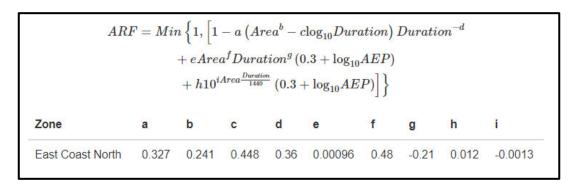


Figure 2.18: Long duration ARF parametric equation for Laidley Creek (Babister et al. 2016)

For a design scenario with a duration between 12 and 24 hours, the interpolation methods described in Chapter 4.3 of ARR 2019 Book 2 by Jordan, Seed and Nathan (2019) are used to determine ARF.

2.5.3 Rainfall temporal patterns

The temporal pattern of a rainfall event represents the distribution of rainfall intensity over a design duration, which can significantly vary between events of similar magnitude (Jordan, Seed & Nathan 2019). Previous approaches to modelling adopted a simple uniform distribution of constant rainfall intensity, however more recent literature has demonstrated temporal pattern selection has a significant impact on the catchment response, specifically the magnitude of peak discharge and shape of a flood hydrograph (Loy 1990; Ball 1994; Ladson 2014). Ball (1994) established that generated hydrographs attributed to variable rainfall rates had higher peak discharges than those of constant intensity. Further, Loy (1990) detailed that the peak discharge generated from runoff routing models featuring different temporal patterns for design

events differed by over 50%. Temporal pattern selection is a crucial consideration to preserve model probability neutrality between rainfall and discharge.

Temporal 'loading' characterises the nature of a temporal pattern, depending on when the majority of rainfall is received during the event. Specifically, loading is classified by the proportion of the event duration at which 50% of the cumulative event rainfall total has been received (Visser et al. 2023). Book 2 of ARR 2019 (Jordan, Seed & Nathan 2019) classifies loading into three categories, as exemplified in Figure 2.19:

- Front loaded Minimum of 50% rainfall received in the first 40% of event duration.
- Centrally loaded 50% of rainfall received between 40% and 60% of event duration.
- Back loaded Minimum of 50% rainfall received from 60% of event duration onwards.

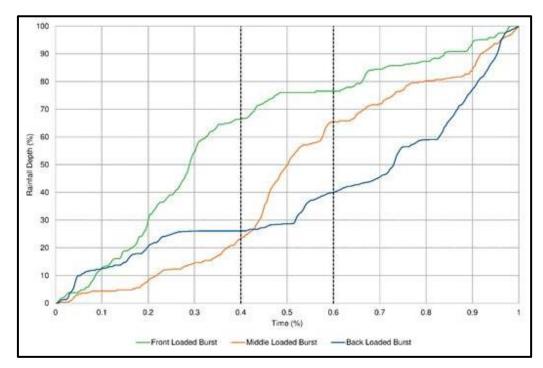


Figure 2.19: Example event cumulative burst distributions (Jordan, Seed & Nathan 2019)

The third ARR 2016 revision project focussed on the development of an ensemble of temporal patterns for design durations and AEPs representative of variable conditions experienced at any location (Loveridge, Babister & Retallick 2015). This project was a progression from the outdated practices of ARR 1987, which utilised a singular temporal pattern for each design rainfall to represent a typical storm event (Loveridge, Babister & Retallick 2015). The ARR 2019 ensemble of temporal patterns are regionalised into 12 subareas of the Australian continent. Laidley Creek is situated within the 'East Coast North' region. Each regional set of temporal patterns encompass durations from 15 minutes to 7 days across 4 burst AEP

categories 'very rare' to 'frequent' (Jordan, Seed & Nathan 2019). Patterns are retrievable from the ARR Data Hub (Babister et al. 2016). The current proportion of design pattern loadings in the Central Slopes region is detailed in Table 2.2. Centrally loaded patterns are most frequent for events both greater and less 6 hours in duration.

<u>Table 2.2: East Coast North burst loading proportion by duration (Jordan, Seed & Nathan 2019)</u>

Region	Duration	Front Loaded	Middle Loaded	Back Loaded
		(%)	(%)	(%)
East Coast	≤ 6 hours	28.9	56.5	14.6
North	≥ 6 hours	23.4	48.5	28.1

2.5.4 Catchment losses

Catchment loss processes are responsible for the proportion of rainfall that is not directly converted to runoff during an event. These processes include:

- Vegetation interception
- Atmospheric evapotranspiration
- Ground surface infiltration
- Surface depression and channel storage

Empirical loss models represent the effects of losses in yielding runoff discharge. Event-specific loss analysis is complex in nature, is time consuming, and requires substantial resources to obtain sufficient site data about the loss processes. Therefore, current practice advises loss values for rural catchments should be inferred from regional information as investigated by Hill, Zhang and Nathan (2016).

ARR 2019 recommends the Initial Loss (IL)-Continuing Loss (CL) model as most suitable for rural catchment design flood modelling (Ball, Weinmann & Boyd 2019). The IL-CL model assigns constant depth values to both the IL and CL for an event. IL is considered as the beginning storm losses (as listed above) that occur prior to the infiltration capacity of the ground being exceeded (Ball, Weinmann & Boyd 2019). Once the surface is saturated and surface runoff begins, a CL is adopted across the remaining storm event duration. The IL-CL model is applied directly to the time-stepped rainfall hyetograph to produce a rainfall excess hydrograph, refer to Figure 2.20. Ball, Weinmann and Boyd noted that IL and CL "do not vary systematically with the severity of the event (therefore) ... loss is independent of AEP" (2019).

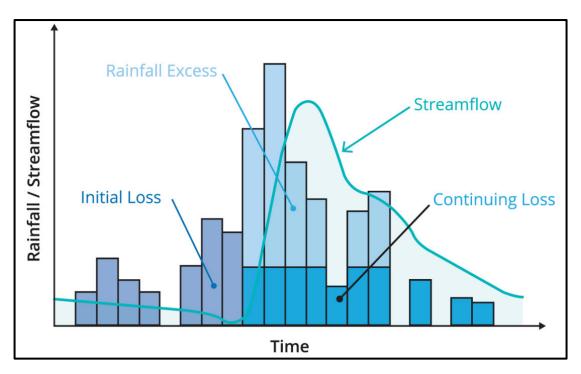


Figure 2.20: IL/CL model application to hyetograph (Ball, Weinmann & Boyd 2019)

Figures 2.21 and 2.22, both extracted from ARR 2019 Book 5 by Ball, Weinmann and Boyd (2019), illustrate the spatial distribution of regionalised losses within non-arid rural catchments, at a discrete incremental scale across Australia. Median initial loss is displayed in Figure 2.21, while continuing loss depth is shown in Figure 2.22.

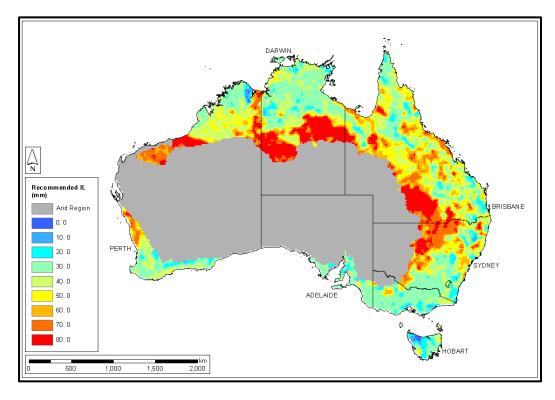


Figure 2.21: Median initial loss distribution for Australia (Ball, Weinmann & Boyd 2019)

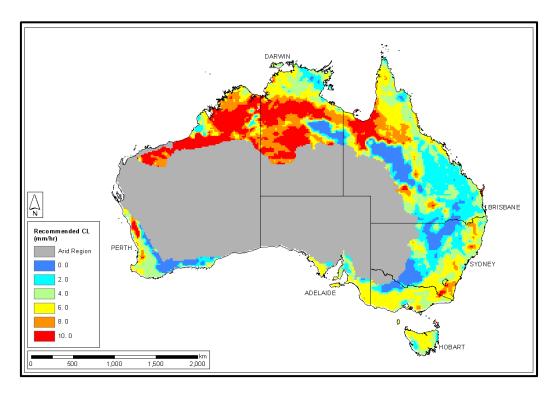


Figure 2.22: Continuous loss distribution in Australia (Ball, Weinmann & Boyd 2019)

Considerations for burst and storm differences are required for event-based modelling. The BOM design IFDs are considered as critical 'bursts' while the ARR 2019 IL depths are for an entire 'storm' event (Ladson 2016). The burst component represents the most intense period with the lowest probability occurrence in the entire event, and generally reflects the whole storm event in larger catchments (Ball, Weinmann & Boyd 2019). The pre-burst component occurs prior to the burst and provides an indication of pre-event conditions including surface saturation and storage, which impacts initial loss characteristics (Ladson 2016). Pre-burst rainfall depths are also retrievable from the ARR Data Hub (Babister et al. 2016). For modelling with BOM IFDs, an initial storm loss IL_s from ARR Datahub require conversion to an initial burst loss IL_b by subtracting the pre-burst rainfall depth PB per equation 2.5.

$$IL_b = IL_S - PB (2.5)$$

(Jordan, Seed & Nathan 2019) noted that for non-urban, non-coastal catchments, the pre-burst rainfall typically only represents a small proportion of the total storm event and had minimal contribution to catchment runoff response. Figure 2.23 illustrates the components of a relatively short duration rainfall event with a high proportion of pre-burst rainfall. The ability to distinguish the individual components of calibration storm events is a crucial preparatory phase to ensure the accuracy of subsequent design event analysis.

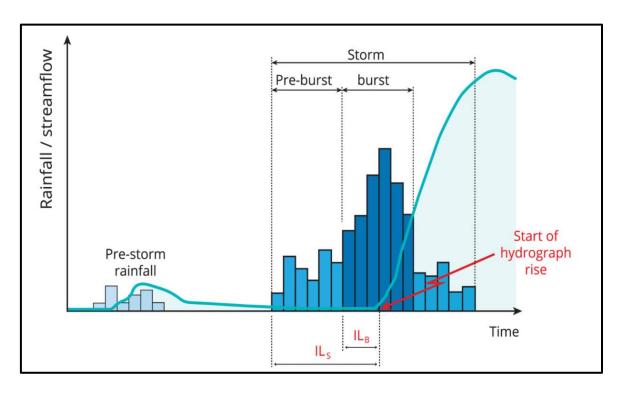


Figure 2.23: Burst and storm initial losses (Ball, Weinmann & Boyd 2019)

2.6 Hydrologic modelling techniques

Discussed in this section are the different hydrologic modelling techniques recommended by ARR 2019 and DTMR to undertake a catchment analysis. The theory and applications of each approach are introduced and the required data inputs are referenced to the previous section. The advantages and limitations of each approach are also discussed in detail.

2.6.1 At-site Flood Frequency Analysis

At-site Flood frequency analysis (FFA) correlates flood frequency and recorded discharge by fitting a probability distribution to a continuous time series of recorded streamflow discharge. Ladson describes FFA as a "statistical analysis of data (to form) useful inferences ... on the magnitude and frequency of future flood events" (2014), while Rima et al. describes FFA as "a widely used statistical technique for estimating design floods" (2022). FFA can only be undertaken within gauged catchments.

Flood peaks are considered as independent random variables in time. From a time-series plot, such as the example presented in Figure 2.24, peak flood discharge data can be classified in two series. These namely are the Annual Maximum (AM) series and Peak-Over-Threshold (POT) series. An AM series is comprised of the highest single flood discharge in each yearly period on record (Ball et al. 2019), capital Q only in Figure 2.6. A POT series consists of all flood discharge peaks that exceed a stipulated threshold discharge (Ladson 2014), capital Q

and lowercase q in Figure 2.24 provided Q also exceeds the threshold. A POT series typically includes K events for N years on record, where N < K < 3N (Ladson 2014).

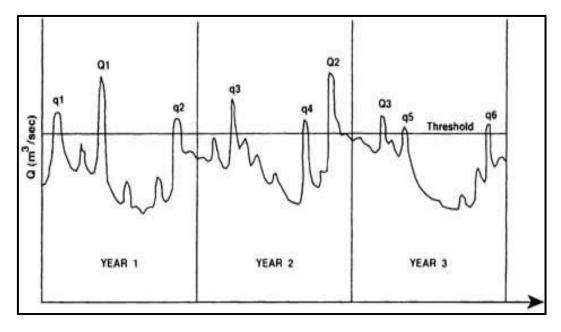


Figure 2.24: Time series plot of discharge data (Holland, Herschy & Archer 1998)

ARR 2019 recommends using the AM series for FFA when considering design events with an AEP rarer than 10% as the modelling associated with this approach is much simpler yet still generates almost identical estimates to the POT series (Ball et al. 2019). Additionally, estimations of the rarest events using POT "may be compromised in order to obtain a good fit to the smaller peaks where the bulk of the data lies" (Ball et al. 2019). The POT series is typically only preferred for common floods for urban stormwater and structure design.

2.6.1.1 Annual Maximum Series

An AM series consists of the peak discharge in each water year, considered in Queensland as the 12-month period spanning from October to September, as these months when lowest average flow is typically experienced and flooding is unlikely (Ladson 2014). All other flood peaks, regardless of their magnitude compared to the maximum peak in another water year, are omitted. Typically, the AM series will contain an equivalent number of data points to record duration in years, however this may not occur in situations with missing data in the gauged records. Ball et al. describes the correlation between adopted data from nearby catchments to produce an estimate of the magnitude of a unrecorded significant flood, and the actual discharge, as "often insufficient" (2019). In instances where the largest discharge definitely occurred outside the period of missing data, the peak discharge for that water year is still usable in the AM series. However, when the unrecorded flood "cannot be estimated with reasonable

certainty" (Ball et al. 2019), the entire water year should be omitted from the FFA. This is because FFA projections assume that the causes and effects of past flood events remain current for future conditions. Consequently, many instances of data omission will impact the accuracy of the fitted probability distribution.

The annual maximum discharges are considered as individual events, provided events where concurrent discharges in short succession were experienced have adopted a singular peak discharge. Event flood probability is described through the application of a probability distribution function to a probability plot of AM gauged discharge data against initial estimates of associated AEP. The probability notation $P(Q \le q | \theta)$ expresses the exceedance probability of a flood q being greater than or equal to a magnitude Q, conditional on input parameters θ (Ball et al. 2019). By ranking the peak annual discharges in descending order, the estimated AEPs are derived from the plotting position method (Ladson 2014; Ball et al. 2019). ARR 2019 recommends usage of the Cunnane AEP plotting position equation for modelling consistency (Ball et al. 2019):

$$P(i) = \frac{i - 0.4}{n + 0.2} \tag{2.6}$$

Where P is the Cunnane plotting AEP in decimal form, i is the gauged flood rank and n is the total number of flood events gauged (Ladson 2014).

Several probability distributions have been previously adopted for FFA. The best distribution is dependent on the complex, often uninterpretable arrangement of the true discharge data.

Consequently. several valid probability distributions have been previously fitted to FFA studies, as noted by Ladson (2014). These distributions include:

- Normal
- Exponential
- Log Pearson III (LP3)
- Generalised Extreme Value (GEV)
- Generalised Pareto (GP)
- Gumbel

The adoption of some models is location dependent, for example LP3 is commonly used in the United States (Singh 1998). In Australia, ARR 2019 recommends either the LP3 or GEV distributions (Ball et al. 2019; Jordan, Seed & Nathan 2019). However, this recommendation is only an informed suggestion, as a selection of the most appropriate distribution function is arbitrary in nature, given any "rigorous analytical proof that any particular probability

distribution for floods is the correct theoretical distribution (does not exist)" (Ball et al. 2019). While the true distribution of flood frequency is indeterminable, Rahman et al. (2013) conducted analyses of 15 distributions with recorded AM series data from Australian catchments, which determined LP3, GEV and GP were the most appropriate. Ladson (2014) and Ball et al. (2019) argued that the results of such empirical studies cannot correlate to conclusive evidence towards the universal adoption of a certain distribution model, because of sample variability effects arising from the short duration of available gauge records.

The 2013 study conducted by Rahman et al. identified that within Queensland, LP3 with method of moments was the best-fitting distribution across 56 gauging sites.

2.6.1.2 Log Pearson III Distribution

The LP3 distribution is characterised by three parameters, namely scale, shape and location (Desvina et al. 2019). These parameters are determined by using the indirect method of moments to calculate the statistical moments of the dataset containing the logarithms of the Cunnane discharges (Ladson, 2014). The statistical moments considered are:

- mean, representative of the central axis of the data,
- skew, a measure of data symmetry, whether the data is mostly large or small, and
- standard deviation, a measure of data spread or separation (Singh 1998).

From the statistical moments, the LP3 distribution is represented by the general equation:

$$\log(Q_Y) = M + K_Y(g) \times S \tag{2.7}$$

Where Q_Y is the discharge of a 1 in Y AEP event, M, g and S are the mean, skew and standard deviation respectively of the log(Q) series, and K_Y is the frequency factor of the 1 in Y AEP (Ball et al. 2019). Equations for the moments of mean, skew and standard deviation for sample data x_i are:

$$\bar{x} = \frac{1}{n} \sum_{i=1}^{n} x_i \tag{2.8}$$

$$S = \left[\frac{\sum_{i=1}^{N} (x_i - \bar{x})^2}{n-1} \right]^{0.5}$$
 (2.9)

$$g = \frac{n\sum_{i=1}^{N} (x_i - \bar{x})^3}{(n-1)(n-2)s^3}$$
 (2.10)

Where $x_i = \log q_i$, the logarithm of the Cunnane discharges (Ladson, 2014).

 K_Y is approximated by the Wilson-Hilferty transformation:

$$K_{Y}(g) = \begin{cases} \frac{2}{g} \left[\left\{ \frac{g}{6} \left(Z_{Y} - \frac{g}{6} \right) + 1 \right\}^{3} - 1 \right] & \text{if } g \neq 0 \\ 0 & \text{if } g = 0 \end{cases}$$
 (2.11)

Where Z_Y is the frequency factor for the standard normal distribution, given in ARR 2019 Book 3 Table 3.2.2 (Ball et al. 2019). Approximations of $K_Y(g)$ are provided in Table 7.7 of Haan (1977).

2.6.2 Runoff routing

Runoff routing is a rainfall event-based method of runoff hydrograph estimation. The approach simulates the flow characteristics during a flood event through a "series of conceptualised ... storages (to) represent the attenuation and translation effects of a catchment on the rainfall-excess hyetograph" (Laurenson, Mein & Nathan 2010). The translation and attenuation between the peaks of an inflow hyetograph and an outflow hydrograph, as shown in Figure 2.25, is indicative of the impacts of catchment spatial characteristics located between the rainfall inputs and the downstream location of analysis. Translation is representative of the lag time of flows along a stream or channel, which is directly proportional to the reach length (Ball, Weinmann & Boyd 2019). Attenuation considers the reduction in the magnitude of the flood hydrograph peak caused by concentrated and distributed storages located within a catchment, as well as transmission losses (Ball, Weinmann & Boyd 2019). Routing through additional storages diffuses the hydrograph, causing a further reduction in the peak flow while extending the duration of the discharge event (Laurenson, Mein & Nathan 2010), as evident in Figure 2.25. Forms of temporary flood storage spread throughout a catchment include:

- Depressions within overland hillsides
- Stream channels and banks
- Floodplains

Contemporary flood hydrograph modelling conceptualises these distributed storage elements as a combined network (Laurenson, Mein & Nathan 2010; Ball, Weinmann & Boyd 2019). Additionally, lakes, basins, reservoirs and dams contained within a catchment are considered concentrated storage elements as the relationship between these storages and inflow-outflow characteristics are usually more direct (Ball, Weinmann & Boyd 2019).

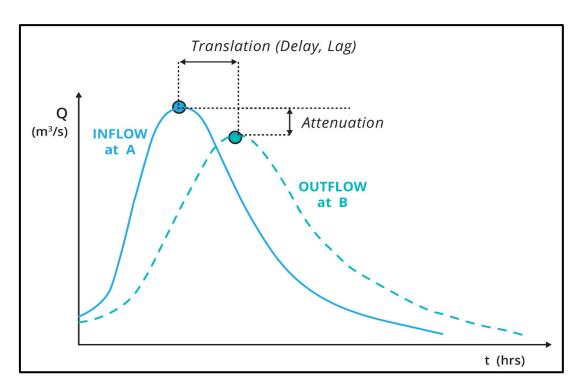


Figure 2.25: Storage and travel impacts between inflow and outflow hydrographs (Ball, Weinmann & Boyd 2019)

The spatial representations of a catchment are classified by the form and level of detail that the temporary flood storage and associated hydrologic phenomena is conceptualised. Homogenous lumped models use the time-area or unit hydrograph approaches to determine flows from small catchment or individual subcatchments within a larger catchment area (Ball, Weinmann & Boyd 2019). These approaches consider spatially uniform rainfall depths and patterns; baseflow and loss characteristics.

Semi-distributed routing methods account for the variability of these factors by partitioning the larger catchment into homogenous subcatchments. The runoff generated by each subcatchment is routed downstream through a quantified reach storage (Laurenson, Mein & Nathan 2010). The nature of semi-distributed methods complex hydrologic phenomena that influence the timing and magnitude of runoff during a flood event to be modelled. For example, non-linear catchment responses through branched networks and significant storages are conceptualised by such methods (Ball, Weinmann & Boyd 2019).

Node-reach models are the most adopted form of semi-distributed models due to their relative ease of establishment whilst being sufficiently representative of catchment configurations of varying size and complexity. In a typical node-reach model, a rainfall input is applied across each subcatchment which is converted from an excess hyetograph to an inflow hydrograph. These hydrographs are routed through reaches within the catchment, where the runoff

production is determined at critical nodes (Ball, Weinmann & Boyd 2019) ultimately concluding at the outlet where the total magnitude and timing of the runoff hydrograph is computed from all contributory flows (Laurenson, Mein & Nathan 2010; Ball, Weinmann & Boyd 2019), as demonstrated by the conceptualised model diagram in Figure 2.26 below.

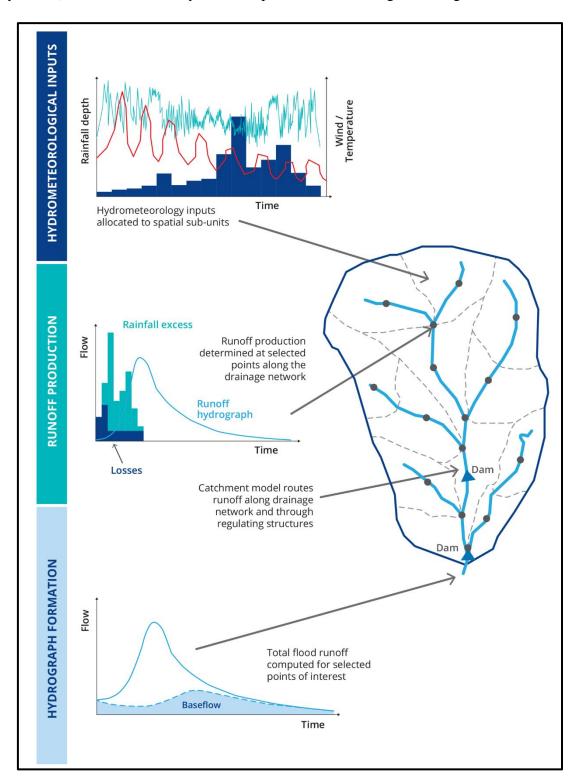


Figure 2.26: Node-reach runoff routing model processes (Ball, Weinmann & Boyd 2019)

2.6.2.1 Node-reach model features

Laurenson, Mein and Nathan (2010) outline the types of nodal features within semi-distributed models, including:

- Input nodes for excess hyetographs/inflow hydrographs at subcatchment centroids.
- Input nodes for inflow hydrographs from upstream subcatchments entering another subcatchment.
- Junction nodes between reaches.
- Reservoir storage bodies.
- Output nodes at gauging stations and locations of computed runoff hydrographs.

Linear reaches represent the stream routing links between node features.

2.6.2.2 Subcatchment area delineation

The subcatchment delineation method is a simplification of the spatial composition of a broader catchment. In a similar approach to catchment delineation, subcatchments should be defined by significant topographic features that bound sub-areas with homogenous flood characteristics (Ball, Weinmann & Boyd 2019). The number of subcatchments used was first demonstrated by Weeks (1980) to affect the model, specifically that as the number of subcatchments increased, the hydrograph peak increased and was postponed. Ball, Weinmann and Boyd (2019) suggested delineating between 10 and 100 subcatchments was common contemporary practice, depending on the total catchment area. Laurenson, Mein and Nathan (2010) recommended a minimum of five subcatchments be delineated upstream of a node where a runoff hydrograph was to be computed.

2.6.2.3 Non-linear storage routing

The general form of a storage discharge relationship correlating inflow and outflow through a routed reach is expressed by Ball, Weinmann and Boyd (2019) as:

$$S = KQ \tag{2.12}$$

Where S is the storage volume of a routed reach, Q is the discharge through the reach and K is the lag parameter between inflow into and outflow from the reach. However, the relationship between storage and discharge in routed features is typically non-linear (Ball, Weinmann & Boyd 2019), best represented by a power function first proposed by (Laurenson 1986):

$$S = kQ^m (2.13)$$

Where k is a dimensionless coefficient deduced from K and m is a dimensionless exponent indicative of the relative efficiency of storage and discharge with respect to flow magnitude.

Considering equations 2.12 and 2.13, the lag parameter K can be expressed in the form:

$$K = \frac{s}{o} = kQ^{m-1} \tag{2.14}$$

Equation 2.14 indicates three scenarios that dictate how lag time is affected by changes in the m parameter value. For the scenario m = 1, the relationship between storage and discharge is linear, meaning increases in storage and discharge occur at an identical rate and the lag time is constant for all storage-discharge combinations (Laurenson, Mein & Nathan 2010). Where m < 1, the lag time decreases with increasing magnitude of discharge, meaning the flow is characterised as efficient. Conversely, where m > 1, the lag time increases as storage increases, indicating flow is inefficient. Typical values of m for natural catchment streams were determined to be between 0.6 and 1.0, with 0.8 being widely adopted as an average value for modelling (Laurenson, Mein & Nathan 2010; Ball, Weinmann & Boyd 2019).

2.6.3 Continuous catchment simulation models

Continuous catchment simulation models, such as AWBM, Sacramento and Simhyd (Podger 2004), convert a complete series of climatic data inputs, usually spanning many years, into a set of runoff flows for the corresponding period as a reflection of "the full spectrum of flood and drought conditions (experienced)" (Nathan et al. 2019). Continuous simulation models use parametric values to represent catchment hydrologic characteristics including surface storage capacities and distributions of baseflow recharge and streamflow (Nathan et al. 2019).

Rainfall, potential evapotranspiration and gauged stream discharge datasets at a daily interval are required to develop a continuous simulation of a catchment. The input datasets must be complete for the entire simulation period, meaning that gap filling and stochastic approaches (Nathan et al. 2019) are usually required to address incomplete (missing data) or inadequate (short duration) data records.

A calibration of the provided climate inputs against the observed runoff data is undertaken by using a selected data optimisation technique. Then, a selected objective function assesses the calibration measure of fit and governs the criteria of which the optimal set of model parameters is determined from (Podger 2004). Once the simulated streamflows are generated, the frequency and magnitude attributes of previous flood events are determined using similar methods to FFA (Nathan et al. 2019).

An analysis of continuous simulation models performed by Ling et al. (2015) demonstrated that the method was unable to concurrently replicate a long term hydrograph in conjunction

with rare flood event characteristics to a sufficient level of accuracy. When calibrated against a comprehensive flow record, "the highest flow peaks were under estimated and the flood frequency curve calculated from simulated annual maximum series provided a very poor fit to the observed flood frequency curve" (Ling et al. 2015). Conversely, calibration against rare flood peaks reduced the model fit against the objective function criteria, while only yielding "slight improvements in matching the observed flood frequency curves" (Ling et al. 2015).

2.6.4 Regression methods of analysis

Regression based methods are simple techniques to estimate peak flow at the outlet of a catchment. Two such Australian regression methods are discussed below.

2.6.4.1 Parametric regression

Regional Flood Frequency Estimation (RFFE) is an alternate method of peak discharge estimation in both gauged and ungauged catchments. RFFE uses a regionalised least squares regression of the LP3 distribution parameters attributed to data obtained from 853 gauged catchments (Ball et al. 2019). The method transfers readily accessible flood characteristics, listed below, from a regional set of gauged catchments to any location through regional estimation equations. These equations provide consistent results with gauged records and are dependent on five variables available from BOM, WMIP and ARR Datahub:

- Area of the catchment
- ^{50%}I_{6h} and ^{2%}I_{6h} design intensities
- Ratio of ${}^{2\%}I_{6h} / {}^{50\%}I_{6h}$
- Shape factor, the "shortest distance between catchment outlet and centroid divided by the square root of catchment area" (Ball et al. 2019).

In a similar nature to the development of IFDs from a finite number of data stations, uncertainty exists in the RFFE method in discrete gauged data extrapolation to ungauged catchments. RFFE is not suitable for urban catchments or areas greater than 1000 km², because the technique was developed from data gauged in rural catchments with areas less than this threshold (Ball et al. 2019).

2.6.4.2 Quantile regression

The Palmen and Weeks (P&W) quantile regression technique is similar to RFFE, however regression equations were directly fitted to Queensland gauged data characterised by catchment area and ^{2%}I_{72h}, rather than the distribution parameters. The regression equations yielded "reasonable results (against) independent streamflow data (and were) superior (to) the Main

Roads Rational Method" (Palmen & Weeks 2011). P&W is only appropriate for rural catchment areas less than 1000 km² (Palmen & Weeks 2011) and is typically less adopted compared to RFFE (Ball et al. 2019).

2.7 Future Climate Considerations

2.7.1 Overview of climate change

Climate change has drastically affected catchment hydrologic processes in regions globally. Current Queensland climate variability is experienced across seasonal, yearly and decadal oscillations, which are correlated to the patterned occurrence of rainfall events (Mora et al. 2013). In the upcoming 20 to 50 years (Mora et al. 2013), the effects of climate change are expected to become "increasingly pronounced (and) potentially significant" (Alluvium 2019). Long-term water resources planning must consider the projected climatic conditions and provide adaptations to mitigate adverse impacts on human populations and natural ecosystems, as well as the agricultural and industrial sectors.

Since recordkeeping began in 1910, the Australian surface temperature has risen by approximately 1.5°C (Bureau of Meteorology & CSIRO 2022). This increase is attributable to industrial greenhouse gas emissions from the energy, transportation, manufacturing and agricultural sectors (United Nations Sustainable Development Goals Summit 2023). The extent of further change is dependent on the magnitude of future emissions and the subsequent climatic response. Representative Concentration Pathways (RCPs) were developed by the Intergovernmental Panel on Climate Change (IPCC) as a projection of four characterised scenarios of greenhouse gas atmospheric concentrations and total emissions mass until 2100 (Van Vuuren et al. 2011), as indicated in Figure 2.27. Each scenario is classified as RCP X, where X is the global radiative forcing energy causing warming in 2100, relative to pre-industrial levels, in units of watts per square metre (IPCC 2018).

RCP 8.5 is the most extreme scenario, representing a future with minimal changes to current emission rates, causing CO2 concentrations to continually rise. The intermediate scenarios, RCP 4.5 and RCP 6.0 depict scenarios featuring a short-term moderate rise, followed by a decline and eventual stabilisation in emissions by 2100. RCP 2.6 is the most ambitious scenario, where a decline in total CO2 concentration is forecast in the near future due to rapid reductions in industrial greenhouse gas emissions (Alluvium 2019). The RCPs 4.5 and 8.5 are recommended for Queensland climate projections to provide a realistic envelope of future potential emissions (Alluvium 2019; Bates et al. 2019).

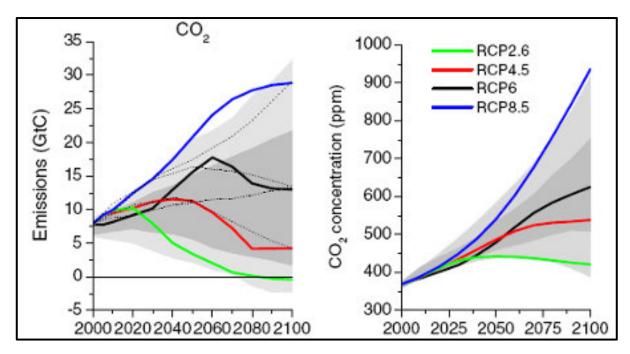


Figure 2.27: Projected CO2 emissions (mass, left) and CO2 atmospheric concentration (ppm, right) for RCPs until 2100 (Van Vuuren et al. 2011).

From the RCPs, three dimensional grids were used to develop representations of the physical processes that occur on and between the land and ocean surfaces as well as the atmosphere. These models are known as General Circulation Models (GCMs) and are simplifications of physics concepts such as the conservation of mass, energy and momentum within a closed system (Alluvium 2019). The GCMs were developed at a global scale, having a horizontal resolution of 200 to 300 km and 20 to 50 vertical layers distributed from the atmosphere to the surface level (CSIRO and Bureau of Meteorology 2020).

Illustrated in Figure 2.28 are the projected global mean temperature rises associated with the four RCP scenarios relative to a thirty year baseline from 1961 to 1990. Alluvium (2019) specified that for the two design emission scenarios, the increase in global average surface temperature by 2100 is predicted to range between:

- RCP 4.5: 1.1 2.6°C
- RCP 8.5: 2.6 4.8°C

The variation in predicted surface temperature increases for each RCP is caused by differences and uncertainties associated with the parameters of the unique GCMs. The period used to establish a relative baseline temperature can vary between models. The 2015 projections released by the Australian Government were derived from 47 individual GCMs (CSIRO and Bureau of Meteorology 2020). GCMs are continually refined and are seen as a motive for the global community to commence measures to mitigate climate change.

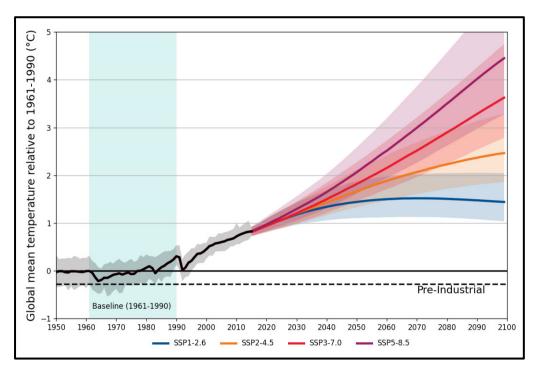


Figure 2.28: Project global temperature rises associated with the RCP climate scenarios relative to the 1961-1990 baseline (Wasko et al. 2024)

The margin of error associated with each global temperature change as presented in Figure 2.28 (and even the predictions by Alluvium (2019)) characterises the uncertainty of the individual GCMs as well as the nature of such forward-forecasting modelling, which is based upon continual revision with up-to-date climate observations and inputs.

GCMs are continually refined and are seen as a motive for the global community to commence measures to mitigate climate change. Contextualising the coarse atmospheric simulations of the GCMs for the geographic conditions of Queensland is achieved through detailed statistical downscaling. Downscaling translates the 200 km resolution GCMs to finer scale 10 to 50 km grids known as Regional Climate Models (RCMs) (CSIRO and Bureau of Meteorology 2020). The finer resolutions provides "greater detail and more accurate representation of localised extreme events" (Alluvium 2019). The GCMs and RCMs are capable of modelling components of hydrology including rainfall, pan-evaporation, runoff and channel flows (Alluvium 2019). However, the models are typically too coarse and are not calibrated to topographic runoff models and streamflow data (Alluvium 2019), making their direct hydrologic outputs unreliable. An opportunity exists for hydrologic modelling to be undertaken based on future climate projections obtained from the GCMs and RCMs.

2.7.2 Future climate influence on hydrologic components

Temperature rise is projected to modify the occurrence and extent of extreme rainfall events in Australia. The flood estimation aspects of IFD relationships and rainfall temporal patterns are most likely to be impacted (Bates et al. 2019) as discussed in the following subsections.

2.7.2.1 Intensity and frequency

Due to increased surface and atmospheric temperatures in the future, an increase in the intensity of extreme rainfall events is projected. The Bureau of Meteorology and CSIRO (2022) has already observed an increase in the intensity of extreme one hour events by approximately 10%, because for every degree of temperature warming observed, the atmospheric moisture content has risen by 7%, thus increasing the available energy to generate intense rainfall events (Dowdy 2020). Conversely, total annual rainfall and continuous steady rainfall patterns are occurring less frequently, which was attributed to "a trend towards higher surface atmospheric pressure ... and a reduction in the number of cold fronts that produce rainfall" (Bureau of Meteorology & CSIRO 2022). Bates et al. (2019) emphasised that national trends are highly variable and regionally sensitive.

2.7.2.2 Temporal patterns

Visser et al. (2023) investigated the relationships between temperature, intensity and duration on the temporal distribution patterns of Australian rainfall events. The research identified that a rise in storm temperature coincided with an increased proportion of front-loaded events (Visser et al. 2023). Additionally, the study found that a "majority of high-average intensity precipitation events are associated with front-loaded storms" (Visser et al. 2023). It was observed that events with a duration spanning less than 6 hours were more commonly front-loaded events compared to events that exceeded 6 hours (Visser et al. 2023).

The following sections outline the direct climatic impacts on the Laidley Creek catchment and guidance provided in state and federal literature for engineers and designs to incorporate climate change scenarios into flood estimation.

2.7.3 Direct climate impacts within the Lockyer Valley region

The mean annual temperature anomaly for Queensland has consistently been above the 30 year average between 1961 and 1990 of 23.2°C since 1985 (Bureau of Meteorology 2024a), as illustrated by Figure 2.29. This trend has even more prevalent in recent years as temperatures continue to rise, as indicated by the highest recorded five year running average of temperature anomaly (Bureau of Meteorology 2024a).

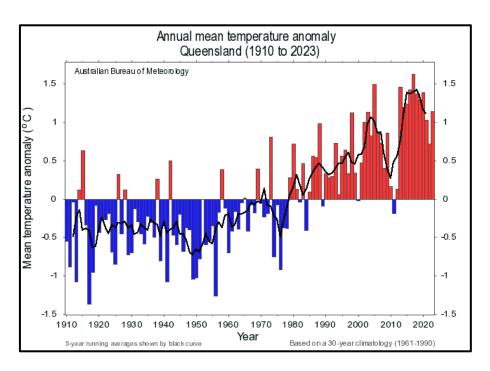


Figure 2.29: Annual mean temperature anomaly in Queensland between 1910 and 2023 based on 30 year average from 1961 to 1990 (Bureau of Meteorology 2024a)

In addition to increasing surface temperature observations, the Bureau of Meteorology and CSIRO (2022) has observed a reduction in the frequency of low-pressure system development responsible for delivering continuous rainfall to the Lockyer Valley region. Consequently, a reduction in the total annual rainfall between the period 1970 to 2022 has been observed for the entire Southern Queensland region, including the LVRC LGA (Bureau of Meteorology 2023). However, while the Bureau of Meteorology and CSIRO (2022) project the total number of rainfall events will decrease, the proportion of rainfall received from intense events in the region will increase. The resulting change in runoff for equivalent AEP events within Laidley Creek under future climate conditions has not been examined at present.

The changes to climate conditions are anticipated to significantly vary between different regions in Australia. The Queensland Future Climate Dashboard (QFCD) is an interactive, gridded database of average climate projections for four upcoming 20 year periods centred around 2030, 2050, 2070 and 2090 (Queensland Government 2024). Factors impacted by climate change, including surface temperature and total rainfall are classified by LGA or major river basin for both design RCPs (Queensland Government 2024). From the dataset, projections for the Lockyer Valley region are expecting increased surface temperatures, less total rainfall and increased concentrations of higher intensity rainfall (Queensland Government 2019).

2.7.4 Design recommendations for future flood estimation

The widespread impact of climate change on climatic components and hydrologic processes is complex in nature and is regionally specific. Consequently, ARR 2019 provides a uniform national approach to future flood estimation methodologies for engineering design purposes in Book 1, Chapter 6. This section of ARR 2019 acknowledges that ongoing research is required to "reduce key uncertainties" (Bates et al. 2019) in the design flood factors from climate change. Current understanding is predominantly concentrated around changes in rainfall intensity from climate change, while other factors are relatively unevaluated at present (Bates et al. 2019). Therefore, Book 3 of ARR 2019 recommends that long-term flood risk is considered exclusively from an increase in rainfall intensity over the project service duration. This recommendation is adopted by DTMR in its Hydrologic and Hydraulic Modelling Technical Guideline (Department of Transport and Main Roads 2024c). Book 2 of ARR 2019 supports this notion, as Jordan, Seed and Nathan emphasised that the AEP probabilistic design rainfall terminology is "equally applicable to both stationary and non-stationary climatic environments" (2019).

The temperature outputs obtained from the GCMs and RCMs are considered more reliable than the rainfall outputs. Therefore ARR 2019 has formulated the process that an adjustment factor is applied to design IFDs proportional to surface temperature projections (Bates et al. 2019). This approach ensures the predictions are representative of temperature rises as the primary cause of extreme rainfall increase and are consistent with the IPCC projections.

The expected rise in extreme rainfall intensity ranges between 2 and 28% per degree of warming (Guerreiro et al. 2018; Bates et al. 2019; Bureau of Meteorology & CSIRO 2022), however ARR 2019 has adopted median recommended rates of change per degree of temperature change depending on the storm event duration.

The IFD adjustment factor p is expressed by equation 2.15 (Wasko et al. 2024) and applies to events between 1 EY and the probable maximum precipitation (PMP) flood in Australia.

$$p = \left(1 + \frac{\alpha}{100}\right)^{\Delta T} \tag{2.15}$$

Where α is the median rate of change estimate in rainfall intensity per degree of temperature rise (%/°C) for a given event duration, per Table 2.3, and ΔT is the mean projected increase in surface temperature relative to the 1961 – 1990 baseline (°C) for a given emissions scenario and projection period, per Table 2.4.

Table 2.3: Rate of change α in rainfall intensity per degree of temperature rise (%/°C) and 66% certainty range in the value (Wasko et al. 2024)

Event duration	Median estimate	Likely range
(hours)	α (%/°C)	[~66%] (%/°C)
≤1	15	7.0 - 28.0
1.5	13.7	6.1 – 25.6
2	12.8	5.5 – 24.0
3	11.8	4.7 – 22.0
4.5	10.8	4.0 – 20.3
6	10.2	3.6 – 19.2
9	9.5	3.1 – 17.8
12	9.0	2.7 – 16.9
18	8.4	2.3 – 15.7
≥ 24	8	2.0 – 15.0

Table 2.4: Mean surface temperature increase projections ΔT for design RCP scenarios relative to 1961-1990 baseline (Wasko et al. 2024)

Climate Scenario	RCP 4.5	RCP 8.5
2030 (near-term)	1.2	1.3
2050 (medium-term)	1.7	2.1
2090 (long-term)	2.4	4.1

The existing design rainfall intensity I is multiplied by the adjustment factor p to yield the projected future design rainfall intensity I_p per equation 2.16 (Wasko et al. 2024):

$$I_p = I \times p \tag{2.16}$$

Alternate surface temperature projections were published as localised gridded datasets within the QFCD (Queensland Government 2024) as listed in Table 2.5. These localised projections generated from downscaled RCMs are potentially more reflective of future conditions in the Laidley Creek catchment than globally averaged increases. However, the approach adopted by ARR 2019 as summarised relies on a common reference baseline period. The difference reference period used between Tables 2.4 and 2.5 is likely responsible for some extent of the differences between corresponding temperature rise values.

<u>Table 2.5: Mean annual temperature rise (°C) relative to reference period 1986 – 2005 for</u> Lockyer Valley (Queensland Government 2024).

Year	RCP 4.5 scenario	RCP 8.5 scenario
2030	0.92	0.93
2050	1.4	1.8
2090	2.1	4.1

2.7.4.1 DTMR climate change risk assessment framework

In June 2024, DTMR released the Climate Change and Natural Hazards Risk Assessment (CCNHRA) Guideline, followed by Engineering Policy (EP) 170 – CCNHRA in July 2024. Together, these documents provide "policy direction, context and background information for considering and responding to climate change and natural hazards risks on (departmental) infrastructure projects" (Department of Transport and Main Roads 2024d).

The CCNHRA framework requires at least two climate projections be incorporated into the management and planning of state controlled infrastructure (Department of Transport and Main Roads 2024f). This approach allows for both short, medium and long term impacts associated with climate change to be assessed. The selection of specific climate projection dates (2030, 2050 and 2090 typically) is dependent on the forecasted design life of the component or asset (Department of Transport and Main Roads 2024d). One projection must consider climate conditions beyond the operational lifespan of the asset (Department of Transport and Main Roads 2024f).

The stipulated minimum design life for asset components within the state controlled road network varies depending on the component purpose and criticality. Pavements have a general design life of 30 years, bridge drainage structures are designed for 50 years of operation while abutments are prescribed the longest design lifespan of 100 years (Department of Transport and Main Roads 2024f, 2024a).

2.8 Summary

From an extensive study of current literature, it has been established that the impacts of climate change at a localised scale in the Laidley Creek catchment are not yet understood. This presents an evident research gap appropriate for this dissertation to address. This literature review identified the unique characteristics of the Laidley Creek catchment that contribute to the

occurrence of complex regional hydrologic processes. The relationship between natural stream hydrology and the Warrego Highway were introduced, including flood immunity requirements for the state controlled road corridor. Standardised modelling techniques to comprehend and model rainfall and discharge processes were outlined. The development of future global emissions models and the corresponding expected temperature increases was discussed at length, including the necessary downscaling process to contextualise the projections to Australian conditions. This theory backgrounded the expected impacts on hydrologic processes within the Lockyer Valley. Recommendations provided by ARR 2019 and DTMR explain how infrastructure design and management should consider future climate scenarios, including anticipated increases in rainfall intensity. This guidance will assist the development of the methodology for this dissertation. However, the specific effects of this projected rise in rainfall intensity across the Laidley Creek catchment are yet to be measured.

3. Methodology

3.1 Introduction

The methodology documented in this chapter describes the structured modelling approach developed from the comprehensive literature review to fulfil the objectives of this research project. The chapter begins by defining the preliminary scope of research following the literature review and contextualises the objectives of the methodology with best practice modelling principles. The data acquisition and processing works required to develop the model inputs are documented, including any limitations in the availability and quality of data. The generation and calibration phases of the RORB runoff-routing model to current climate conditions in the Laidley Creek catchment are detailed. The application of independent modelling approaches as a comparison of the results produced by the runoff-routing model is described. The procedures implemented and software applications throughout the methodology are documented for clarity and evaluated against the design objectives. Finally, the project methodology is critically evaluated against a range of alternate research techniques as a measure of its credibility and validity.

3.2 Design methodology overview

The methodology for this research project was developed from insights and industry practices gained through the literature review phase to fulfil the project objectives. The review backgrounded the translation of physical processes into numeric hydrological parameters for the development of empirical modelling techniques, and subsequently, the advanced modelling practices and software applications used currently in Australia. Finally, the impacts of climate change on catchment hydrology and methods to incorporate these impacts into hydrologic models were outlined.

As such, this methodology consolidates the broad literature review into a concise course of action to develop and evaluate hydrologic models which address the current gap in knowledge regarding the anticipated changes in peak streamflow discharge from high intensity rainfall events. All data preparatory and model development works were completed on a personal laptop with a high-performance CPU and GPU capable of running hydrologic software packages.

3.2.1 Preliminary methodology scope

The methodology of this project was designed to develop hydrologic models from accessible yet reputable data sources to simulate current and future climate peak flow characteristics.

The major practical components of the research methodology outlined in this chapter are:

- 1. Acquisition and processing of high quality spatial and hydrologic data into suitable formats for direct input to different model packages.
- 2. Generation of the routing model catchment and storm files from the processed data according to model conventions. The freely available RORB package was identified early in the project as a suitable platform to facilitate modelling.
- 3. Calibration of the routed model catchment parameters based on a wide range of previously observed rainfall runoff relationships.
- 4. Simulation of a collection of design rainfall events for the standard AEPs to define the magnitude of streamflow yielded for current climate conditions. These simulations were then repeated for revised design rainfall parameters representative of anticipated future climate conditions. Results of both iterations are presented in Chapter 4.
- 5. Comparison of the results (where applicable) to independent hydrologic methods, including FFA and regression methods.

Initially, it was proposed that the scope of this research project could incorporate two-dimensional hydraulic modelling at the Jack Martin Bridge and the flood-prone intersection of the Warrego Highway and Forest Hill-Fernvale Road to understand flood flow characteristics through the roadway section. However, the resolution of available aerial elevation data was insufficient to produce accurate cross sections of Laidley Creek to develop such models. Additionally, the resource requirements to survey surface elevation data at multiple channel sections were not viable for such a student research project. Hence the project scope was limited to hydrologic modelling, with the results available for future hydraulic analyses.

3.2.2 Best practice modelling considerations

The practical components of the research outlined above must be supported by best practice modelling considerations to ensure that the results produced are accurate and reliable. Ensuring the quality of modelling is high also ensures outputs are useful for future research endeavours. A range of positive modelling practices and considerations for water resources and catchment modelling were synthesised by Jakeman et al. (2018) in the *Good Modelling Practice* discussion paper published by the Queensland Water Modelling Network (QWMN). The paper describes the outcome of best practice modelling as a reduction in model uncertainties while documenting "any uncertainties and assumptions for user transparency" (Jakeman et al. 2018). Conversely, poor modelling practices have negative implications which affect the reliability of

the outputs. Many of the practices described in the QWMN paper are worthwhile adopting in a contextualised capacity to enhance the methodology of this project.

Such practices include:

- A thoroughly documented introduction of the acquired datasets and their attributes, including a critical review of the quality and reliability of the data based upon any limitations and its applicability for modelling. Data management and processing should also be explained thoroughly.
- A justified explanation of the selected modelling and calibration approach, including an evaluation of any assumptions and alternative methods, ultimately deducing why the chosen method is most appropriate to conceptualise the hydrologic and catchment characteristics with respect to the research objectives.
- An extensive model testing program to verify the model suitability across a comprehensive range of conditions.
- Sensitivity analyses of the input data and model parameters to evaluate the sensitivity and uncertainties embedded within the results.
- A sanity check of the model results against independent methods of analysis to evaluate the reasonableness of the adopted method.
- Complete and concise reporting of the model results with an emphasis on the applicability of the outputs in context of the project objectives, and any uncertainties within the data. The attachment of the input and output model files within the report appendices for clarity and repeatability is a crucial practice.

3.3 Data acquisition and preparation

The input data required to develop the hydrologic models was retrieved from official primary sources including the DRDMW operated WMIP, the BOM managed DRDS and the ARR Data Hub. Data retrieval from government and academic sources with documented quality assurance procedures, measurement recordkeeping and transparent statements about the limitations in the datasets ensures the highest possible input quality for modelling. The types of data retrieved, any associated limitations, and subsequent processing measures used in preparation for hydrologic modelling are described below.

3.3.1 Catchment geographic spatial data

As an open-source, free geographic information system package, QGIS was chosen to process catchment spatial data for this project. Primarily, the software was used to delineate the boundary and determine the flow routes of Laidley Creek and its tributaries throughout the

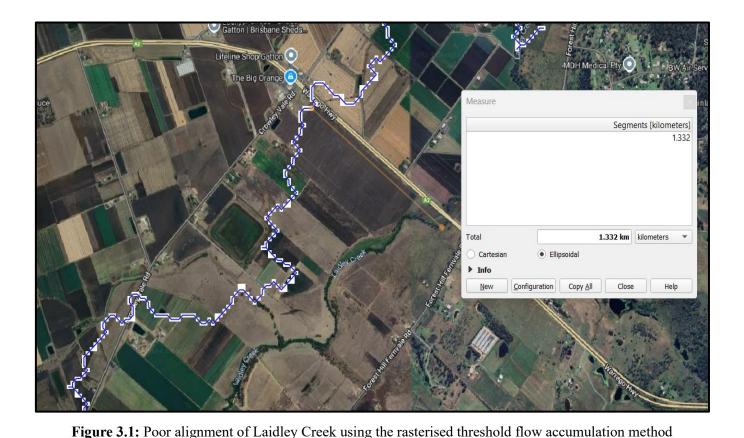
catchment. In addition, the locations of rainfall and streamflow stations were visualised in the software to calculate spatially averaged metrics for subsequent modelling.

A suitable digital elevation model (DEM) sourced from the Shuttle Radar Topography Mission (SRTM) 1 Arc-Second Global V3 dataset published by the Earth Resources Observation and Science Center of the United States Geological Survey (2014) was downloaded as a GeoTiff file covering an indicative outline the catchment region. The file corresponding to the S28° E152° DEM was imported into QGIS and reprojected to the World Geodetic System 1984 (WGS84) coordinate reference system. WGS84 is geographically accurate to within 2 metres which is sufficient for the scale of this catchment.

Two approaches were considered to generate the catchment boundary and stream routes. The first approach is preferred as it utilises hydrologic simulation methods available through open-source QGIS plugins (including SAGA, GRASS and PCRaster) to determine spatial flow characteristics based upon a terrain analysis of the DEM raster cells. The general procedure for this approach consists of the following stages:

- 1. Filling surface depressions in the DEM to maintain correct hydrologic functionality.
- 2. Computation of flow accumulation and direction based upon the gradient and orientation of the terrain at a rasterised cell level.
- 3. Calibration of flow accumulation to the existing stream features shown within satellite and mapping imagery of the catchment
- 4. Delineation of the catchment streams above the calibrated threshold level.
- 5. Delineation of the catchment boundary upstream of an identified outlet based upon the DEM cell elevations within the catchment.

The advantage of this approach is that the automatic processing of the catchment boundary and streams is a relatively quick procedure to execute, and provided the input DEM is of sufficient vertical resolution, the results are usually more accurate than manual delineation. However, the main disadvantage of the flow accumulation method is that modelling expansive floodplains with very flat terrain or artificial storages often results in misaligned flow routes (Al-Muqdadi & Merkel 2011; van der Kwast 2018). When the automatic stream delineation method was applied for the Laidley Creek catchment, the outlet was displaced 1.3 kilometres north-west of its known location and the general flow path did not align with the existing route of Laidley Creek, as illustrated in Figure 3.1. The catchment area was computed as 450.6 km², which represented a 2.5% underestimation of the 462 km² area stated within the WMIP details for station 143229A (Department of Regional Development, Manufacturing & Water 2024).



The results obtained by this method were deemed unsuitable for further modelling given the misaligned Laidley Creek flow path, considering the proximity of the actual outlet site to the Warrego Highway crossing. Hence, the second, manual catchment delineation method was adopted to reduce the irrationality of the delineated catchment features. A 25 metre interval contour layer was generated from the DEM cell elevation attribute and peaks in the terrain were identified by a set of spatial point features. A second 10 metre contour layer clarified the positioning of any ridgelines and peak elevation where the terrain was difficult to interpret. The

catchment boundary was then established by connecting the identified high elevation locations

with the outline of a polygon shapefile feature, as illustrated in Figure 3.2.

The area of the Laidley Creek catchment upstream of the 143229A station was computed as 452.3 km². This represented a minor improvement on the automatically generated catchment area. The landmass situated between the northern catchment boundary and Lockyer Creek was identified as largest contributing explanation of the difference between the manually delineated method and WMIP catchment areas of 452.3 km² and 462 km² respectively. The terrain within this section is very flat and bounded to the south by the raised corridor of the Warrego Highway. Surface levels and flow patterns have been modified by closed system irrigation and artificial storage networks used by agricultural operators. Therefore, this section was deemed to have

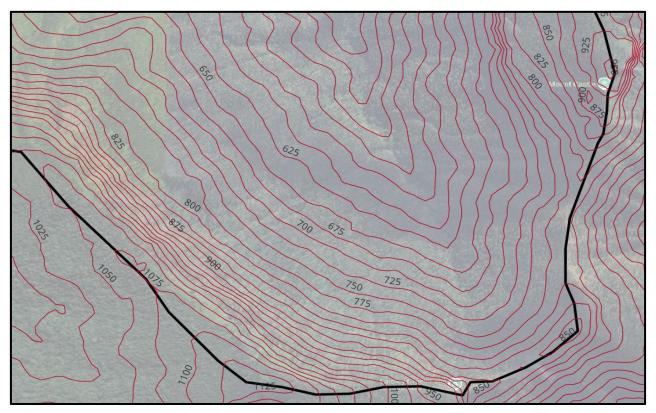


Figure 3.2: Laidley Creek catchment boundary manually delineated from ridgelines identified in the 25 metre contour overlay.

negligible impact on flood conditions at the 143229A station, such that the manually computed catchment area of 452.3 km² was representative of the contributory upstream catchment and its flow characteristics during peak flood events. Any minor boundary errors could hence be considered as having negligible impact on the catchment definition.

The catchment centroid position was determined by executing the QGIS centroid function with the catchment polygon feature as an input layer. The centroid was located at the coordinates 27.727°S 152.355°E, which almost identically correlated with the catchment centroid coordinates of 27.7266 °S 152.357 °E identified by Rahman et al. (2015) in the RFFE method.

To resolve the irrational stream delineation of the flow accumulation method, major stream flow paths were identified from the Open Street Map (OSM) spatial extent query operation using the catchment polygon as an overlay boundary. This approach yielded a series of line features that were accurately aligned to the defined flow paths observed from satellite imagery.

The delineated catchment boundary, centroid and internal stream features are identified in Figure 3.3, with the locations of the two DRDMW operated gauging stations along Laidley Creek shown for reference.

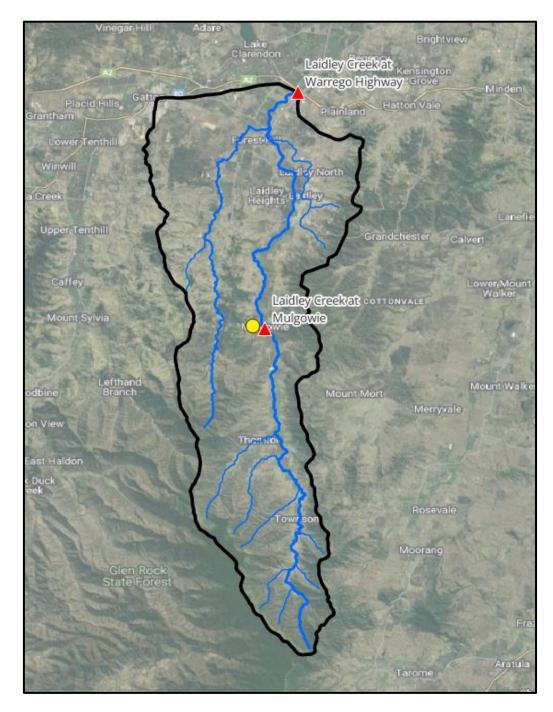


Figure 3.3: Laidley Creek catchment with centroid location, major stream network and gauging stations shown.

3.3.2 Gauged streamflow data

Historical streamgauge data was retrieved for the two Laidley Creek gauges from the WMIP website. The portal provided basic site-specific information essential for developing subsequent models. As listed in Figure 3.4, such details included the latitudinal and longitudinal coordinates, the base elevation relative to the Australian Height Datum (AHD), as well as the upstream catchment area, among other technical characteristics for each gauge.

Site no.	143229A	143209B
Zone	56	56
Easting/Northing	439678.023/6952145.579	437333.126/6932548.553
Latitude	27°33'11.4"S	27°43'47.9"S
Longitude	152°23'20.4"E	152°21'51.2"E
Site commence	31/10/1990	06/03/1967
Site ceased		
Zero gauge	76.313	132.620
Datum	AHD	AHD
Control	Two Metre Crump	Control Weir
Cease to flow level	0.505	1.000
Maximum gauged level	7.654	6.220
Maximum gauge date	28/01/2013	11/01/1968
Distance from stream mouth	5.000 km	31.000 km
Catchment area	462 sq.kms	167 sq.kms
Gaugings	59 gaugings between 01/11/1990 and 04/10/2023	224 gaugings between 06/03/1967 and 05/06/2024

Figure 3.4: Site details for the 143229A and 143209B streamgauges extracted from the WMIP (Department of Regional Development, Manufacturing & Water 2024)

Monthly, daily and hourly datasets of stream discharge and level from each gauged site were extracted from the WMIP in spreadsheet form. Data at varying time intervals was obtained to facilitate the development of different model inputs. The monthly maximum streamflow record of for the primary station at 143229A is shown in Figure 3.5, which importantly, shows that the four highest flow peaks correspond with the four largest events listed in Table 2.1. The baseflow component of runoff was zero or near-zero flow.

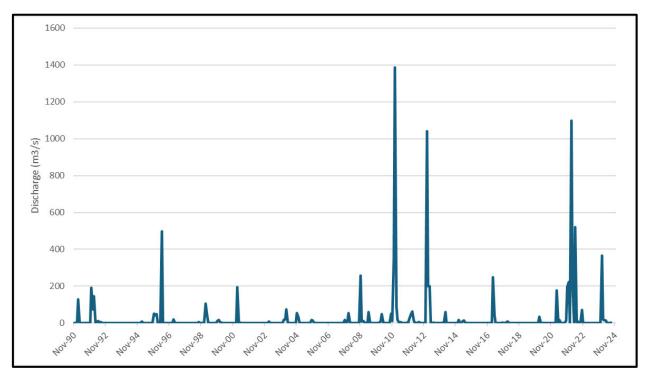


Figure 3.5: Monthly maximum streamflow record for 143229A gauge [generated from data produced by (Department of Regional Development, Manufacturing & Water 2024)]

Data quality codes assigned by DRDMW signify the quality and reliability of the individual data readings. The majority of historical data has been audited and validated by DRDMW and is considered sufficiently reliable for modelling. The most recent data recordings (typically within the last 12 to 24 months) are typically unverified, indicated as a code 40 or 130 (DRDMW 2023). A review of these dates against observed rainfall data confirmed the readings were accurate. Code 15 is used to differentiate flows less than minimum threshold of the gauge from regular flow recordings (DRDMW 2023). However, as this research is not concerned with the estimation of very frequent flood events, the precision of low flow data is not crucial, hence the quality of this data was considered sufficient. Estimated data readings are signified by code 60, which typically is observed for extrapolated high flow events beyond the limit of the gauged rating (DRDMW 2023). Similarly, the code 255 is used to signify unreported or missing data, including inaccurate data removed through an audit (DRDMW 2023). Aside from equipment faults, instances of estimated and unreported data records have occurred more frequently during extreme, high flow events (DRDMW 2023) and require considered judgment about their inclusion within input data sets for peak flow modelling.

The proportion of each data quality classification over the operational period of the primary station of interest, 143229A, were computed and listed in Table 3.1. The majority of data was considered normal, good, fair or below threshold. 4.6% of the data was considered unverified or not coded while 2.6% was estimated. No recording gaps in the data sequence were observed.

Table 3.1: Distribution of data quality by classification at 143229A station

Quality code	Definition	Occurrence (days)	Occurrence (%)
9	CITEC Normal	5343	43.3
10	Good	3323	26.9
15	Below threshold	2433	19.7
20	Fair	358	2.9
40	Unverified	245	2.0
60	Estimate	323	2.6
130	Not coded	318	2.6
255	No reading	0	0

3.3.2.1 Rating curve and control evaluation for 143229A (Warrego Highway)

Flows through the 143229A gauging station are regulated by a two metre concrete crump weir as pictured in Figure 3.6. Low to moderate flow depth readings through the section are reliable

because the fixed concrete hydraulic control section has remained constant during its period of operation. During higher flow events, the water level rises above the weir and the section incorporates the surrounding earth banks, which have a non-uniform soil profile and feature varying levels of vegetation. These areas are more susceptible to change over time from processes such as erosion and sediment deposition. However, the surrounding vegetation would offer some protection of the stream banks, meaning the section has likely remained similar over time. Regardless, the limitation that the section characteristics may not necessarily be maintained between high flow events is worth consideration.



Figure 3.6: 143229A Warrego Highway hydraulic control and surrounding stream section (Department of Regional Development, Manufacturing & Water 2024)

The accuracy of the rating curve at 143229A is crucial towards obtaining useful results from this research given its proximity to the Warrego Highway. The site rating curve, as pictured in Figure 3.7, is well defined by 59 gaugings across an adequate range of flows. The maximum recorded flow has only exceeded the maximum gauging of approximately 980 m³/s on four separate days, with flows of this magnitude still considered fairly reliable by DRDMW Department of Regional Development, Manufacturing and Water (2024). The main limitation with the streamflow record of 143229A is its relatively short operational period, having only opened in November 1991. However, the record obtained is very reliable, as supported by Barton et al. (2015). The streamflow record from 143229A is fundamental to the modelling undertaken in this report.

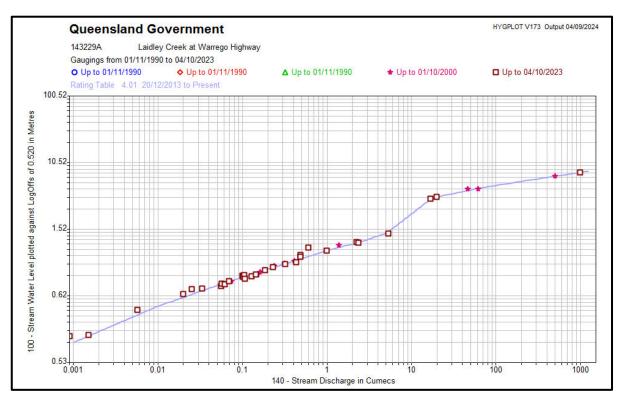


Figure 3.7: Well defined site rating curve from individual gaugings for 143229A station (Department of Regional Development, Manufacturing & Water 2024)

3.3.2.2 Rating curve and control evaluation for 143209B Laidley Creek at Mulgowie

The second gauging station along Laidley Creek is located at Mulgowie and is the outlet of an upper catchment area of 167 km². An earth-rock mound is used to regulate flows through the 143209B gauging station, as pictured in Figure 3.8.



Figure 3.8: 143209B Mulgowie hydraulic control and surrounding stream section (Department of Regional Development, Manufacturing & Water 2024)

The reliability of the streamflow data is compromised by this type of hydraulic control, as the section is vulnerable to movement over time. Cross section instability likely explains the significant variance (scatter) observed between flow gaugings of similar magnitude, as illustrated in Figure 3.9.

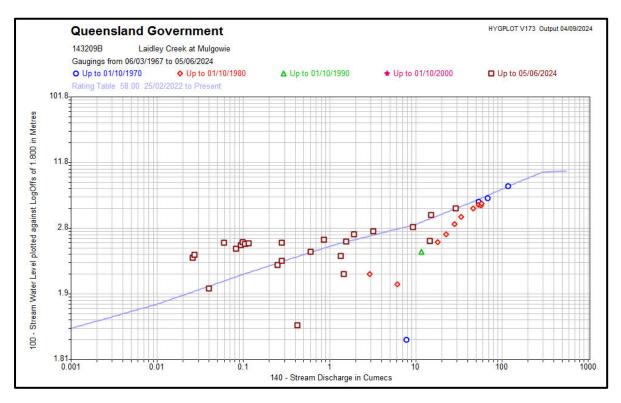


Figure 3.9: Site rating curve with significantly varied gaugings for 143209B station (Department of Regional Development, Manufacturing & Water 2024)

Since the site was opened in January 1968, 224 gaugings have been undertaken, yet the maximum gauged discharge of approximately 121 m³/s is significantly less than the maximum recorded flow of 348.9 m³/s (Department of Regional Development, Manufacturing & Water 2024). The site rating curve is poorly defined, with flows above 10.9 m³/s considered an estimate by DRDMW (2024). Peak recorded flow levels above 9.0 metres are considered 'suspect', indicating the readings are very unreliable. Flows levels of similar magnitude have very different discharges, as illustrated in Figure 3.9. Several limitations associated with the quality of the rating curve at the 143209B gauge, especially at high flows experienced during peak flood events, mean the streamflow datasets from this station are too compromised to be useful for any modelling in this report. All modelling will instead utilise reliable data from the 143229A station, which was the preferred station regardless because of its proximity to the Warrego Highway.

3.3.3 Observed rainfall station data

Recorded rainfall data is required to calibrate the catchment parameters of a runoff routing model against the characteristics of previously observed events. The two identified DRDMW operated streamflow gauging stations situated within the Laidley Creek catchment have an insufficient rainfall record to of any use, having only been upgraded to capture rainfall data in 2024. Therefore, alternative sources of rainfall data at daily and hourly intervals were required.

3.3.3.1 Daily observed rainfall totals

Daily observed rainfall data was sourced from 22 stations operated by BOM through its Climate Data Online portal (2024). Eleven of these stations were located within the delineated Laidley Creek catchment while the remaining eleven were located less than ten kilometres outside of the catchment. Rainfall data observed within the ten kilometre external buffer external was still considered potentially useful because of the rainfall decorrelation distance introduced in the literature review of ten kilometres as described by Podger (2004).

The BOM rainfall station provide a reasonable level of spatial and time coverage across the catchment. An initial viability assessment of the usefulness of each station was undertaken by determining which station datasets contained records during the peak flow events listed in Table 2.1. As shown in Figure 3.10, eight stations had full recorded coverage of the peak flow events while another five had a partial record. The remaining nine were not operational during any of the events of interest. Only three of the thirteen sites were situated within the catchment, however, these stations were well distributed along its north-south axis. The majority of the stations were located within the northern extent of the region and are closely clustered around the population centres of Gatton and Plainland. The recorded rainfall data in this area was able to be corroborated between stations to determine the reliability of each dataset. Additionally, where gaps existed in a station record during an event of interest, a secondary backup site was easily adopted with an adjustment correction factor applied to relate the datasets. In the more remote sections of the catchment, rainfall station coverage was much more limited. Three stations provided some level coverage of the peak flow events listed in Table 2.1 in the central catchment region, while a singular station at Townson offered coverage of the upper-most reaches. Catchment coverage was considered suboptimal through the 20 kilometre section between Mulgowie and Townson. However, in the absence of alternative rainfall data, the maximum isolation distance between any location and a rainfall dataset of approximately ten kilometres meant coverage was deemed sufficient, with limitations about the poor coverage and potential data decorrelation in certain portions of the catchment noted.

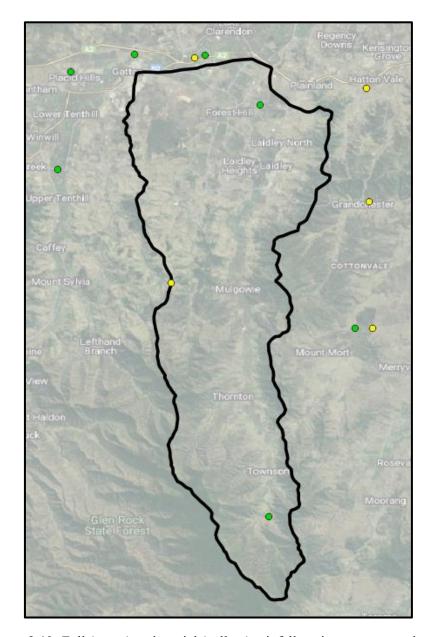


Figure 3.10: Full (green) and partial (yellow) rainfall station event record status

The reliability of the data recorded at each station was evaluated against a baseline average computed for the Forest Hill station. This station was adopted because of its location within the Laidley Creek catchment and its 130 year ongoing record to 30 April 2024. As detailed in Appendix B, the average daily rainfall at the stations in close proximity to Forest Hill conformed closely when averaged over a common operational period. This confirmed the Forest Hill station, with an average daily rainfall depth of 2.13 mm/day provided an ideal representation of the lower catchment area. Higher common period average rainfall depths were observed in the upper catchment reaches, which were not considered unreliable station readings, rather an indication of the spatial variability of rainfall through the mountainous terrain surrounding Laidley Creek.

3.3.3.2 Hourly observed rainfall

Hourly rainfall recordings were required to define the specific temporal distribution patterns of previously observed events. However, none of the rainfall stations within the catchment contained a sub-daily rainfall dataset. Two DRDMW operated streamflow gauging stations in adjacent catchments, namely Tenthill Creek and Adams Bridge (Bremer River), each had an ongoing hourly rainfall record spanning over 30 years (Department of Regional Development, Manufacturing & Water 2024). A third 30 minute rainfall depth dataset at the UQ Gatton station was sourced from BOM (2024c) however did attract a student-use retrieval and licensing fee. This dataset was reformatted at an hourly interval to correspond with the resolution of the other two datasets. The details of each hourly dataset are included in Table B2, refer to Appendix B.

The UQ Gatton and Tenthill Creek stations provide a reasonable level of sub-daily record coverage for the northern portion of the catchment. The relevancy of coverage provided by Adam's Bridge site over the eastern and southern reaches is limited as the station is located further than 10 kilometres from the catchment boundary. However, because no viable alternative dataset exists, the Adams Bridge record will be utilised with caution, noting potential limitations regarding relevancy and event decorrelation.

3.3.4 Design rainfall event parameters

Design rainfall parameters are required for discharge estimation following calibration of the runoff routing model.

3.3.4.1 LIMB-BOM IFD enveloped dataset

To support an accurate and conservative modelling approach, the LIMB-BOM IFD envelope was selected as the design rainfall dataset. The LIMB-BOM dataset was extracted from the ARR Datahub using the input catchment centroid coordinates 27.727 °S 152.355 °E. Compared to the BOM IFDs, the design rainfalls which increased in the LIMB set were less than one hour in duration. All design rainfalls of a duration of more than one hour originated from the BOM IFD set. The full enveloped IFD dataset is presented in Table C1, attached to Appendix C.

3.4.4.2 Projected future climate IFD datasets

The LIMB enveloped design rainfalls were revised to reflect the localised impacts of climate change. The projected rainfall depths for both the RCP 4.5 and 8.5 emissions scenarios by 2030, 2050 and 2090 were calculated according to Equations 2.15 and 2.16, and the values of α and ΔT according to Tables 2.3 and 2.4 respectively, as well as the ARR Datahub. The complete scaled IFD datasets are presented in Tables C4 to C9, attached to Appendix C.

3.3.4.3 ARR Datahub parameters

Design event parameters including storm losses, ARF, areal temporal patterns and preburst depths were extracted from the ARR Datahub in a text file, which is attached in Appendix D.

3.3.5 Finalisation of the design methodology scope

Considering the limited availability of some data as described above, the design methodology scope was finalised. The streamgauge record of Laidley Creek at the Warrego Highway was deemed too short to utilise FFA as the preferred method of peak flow estimation, while the limitations of regression methods were deemed too significant. Hence a runoff routing approach using the RORB package was selected and other methods including FFA, RFFE and P&W were utilised for comparative purposes, as recommended by Jakeman et al. (2018). The streamgauge rating at Mulgowie was considered too unreliable for use as a secondary gauge within the model, therefore catchment calibration and design flow simulations were based solely on the Warrego Highway site. The limitations associated with the daily rainfall record coverage in the upper Laidley Creek catchment, and hourly rainfall record coverage across the entire catchment, were recognised. However, in the absence of alternate data, the BOM and DRDMW sourced datasets were used, knowing that some uncertainty would be introduced to the model. The scaled IFD envelopes encompass three design periods: short, medium and longterm; two potential emissions scenarios: moderate and high. This arrangement provides adequate coverage of plausible future climate conditions, however utilising the median rate of change α instead of a range of values also contributes to the uncertainty of the results.

3.4 Semi-distributed runoff routing model development

Processing and management of the collected spatial, rainfall and gauged runoff data is required to develop the standard format RORB files. Firstly, the catchment spatial features were converted to the equivalent semi-distributed catchment representation for efficient file development in the RORB graphical editor. Then the recorded rainfall and runoff data was analysed to determine the RORB storm file inputs. The catchment and storm files were analysed within the RORB package to calibrate and verify the model parameter values against previously observed events of varying magnitude. Finally, the design event simulation procedure was described in preparation for the detailed results presented within Chapter 4.

3.4.1 Catchment model generation from spatial data

Preparatory spatial modelling within QGIS was required to convert the identified catchment and stream network into a format compatible with a node-reach runoff routing catchment file.

Therefore, the Laidley Creek catchment shapefile was partitioned into 20 subcatchments corresponding to the identified major flow paths and non-uniform sections of terrain. Subcatchments within the southern, mountainous sections were typically smaller than those delineated for the expansive, open downstream sections. This was because defined creak streams form between the steep ridgelines separating the upper subcatchments, while the lower subcatchments generally only contribute overland flows to either Laidley or Sandy Creek as upstream flows are conveyed towards the outlet. As determined from the subcatchment layer attribute table, the maximum subcatchment area was 56.59 km² while minimum was 3.83 km².

As introduced in the literature review, a semi-distributed model assumes uniform rainfall-excess is produced within a subcatchment and enters the network at or adjacent to the centroid of each subcatchment (Laurenson, Mein & Nathan 2010). Therefore, the subcatchment centroid positions were computed in QGIS, represented by the black nodal layer in Figure 3.11. These points were then snapped to the surrounding stream geometry, represented by the orange nodal layer, to represent the network entrance assumptions of Laurenson, Mein and Nathan (2010).

The displacement between the two point layers at 18 of the 20 subcatchments was 500 metres or less, while subcatchment 13 and 19 required additional flow reaches to connect isolated centroids to the stream network. The nature of a semi-distributed model also rendered the stream networks redundant upstream of the snapped centroid of each subcatchment. Therefore, the streams were trimmed at the subcatchment centroids as also shown in Figure 3.11.

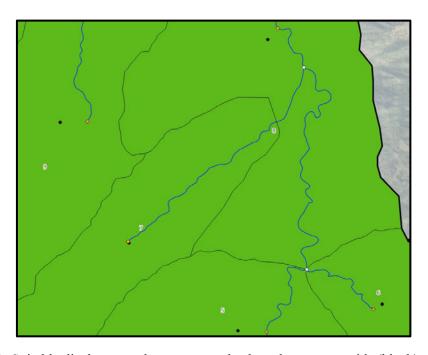


Figure 3.11: Suitable displacement between actual subcatchment centroids (black) and snapped centroid nodes to adjacent stream network (orange)

The reach length of each stream section between the centroidal and junction nodes was extracted from the shapefile attribute table. The partitioned subcatchments, configured centroids and adjusted stream network features are illustrated in Figure 3.12.

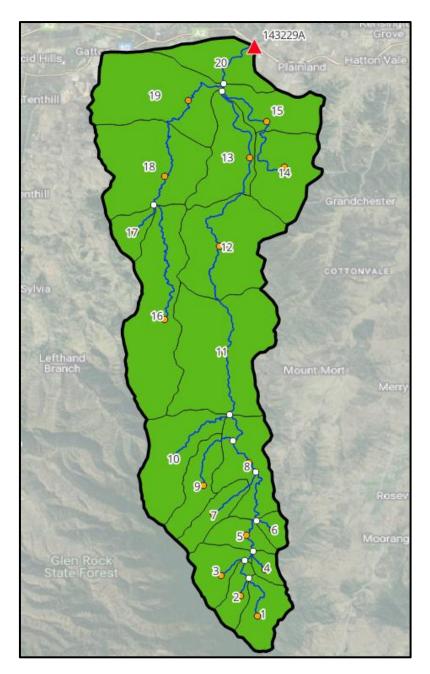


Figure 3.12: Partitioned subcatchments and adjusted stream reaches for RORB catchment file development.

The node-reach configuration of Figure 3.12 was formatted such that a RORB catchment file could be easily developed in the graphical editor package. This approach facilitated the graphical generation of a semi-distributed catchment file compatible with the RORB runoff routing modelling program. The network configuration, consisting of subcatchment areas,

reach lengths and junctions, was replicated in the graphical editor. The slope of each reach was considered as the average gradient between the bounding upper and lower node elevations.

The 6.9 GL capacity Lake Dyer is situated within subcatchment 13. While RORB has functionalities to represent special storage reservoirs as nodes, the upstream catchment that supplies Lake Dyer is only 3 km² (SEQWater 2024). Hence, the reservoir capacity itself and the tributary area were both considered insignificant when compared to Laidley Creek catchment area of 462 km². Similarly, the proportion of impervious land coverage, concentrated about the towns of Laidley and Forest Hill, as well as road surfaces located within the catchment, was deemed insignificant such that all subcatchments were considered to be entirely pervious. It is acknowledged that these simplifications have the potential to introduce some minor uncertainty to the model outputs.

A singular nodal outlet was established downstream of subcatchment 20 to represent the 143229A gauge. For the calibration phase of modelling, the print hydrograph code 7.1 was utilised to compare gauged flows. However, during the design simulation phase, the code 7 was used to generate the simulated ensemble of design discharges. Once imported to the RORB modelling package, the graphical catchment file was converted to a regular catchment file, attached to Appendix E.

3.4.2 Generation of calibration storm files from previously observed events

Rainfall and streamflow data was collected to develop the storm event files for calibration of the RORB model parameters. Storm files were developed for eight peak flow events, including the events listed in Table 2.1 where inundation of the Warrego Highway was confirmed. To ensure the calibration accurately represented as many flow scenarios as possible, events of various magnitude and timing were analysed.

The collected hourly rainfall data had two purposes, firstly, to define the event temporal distributions. Each subcatchment temporal pattern was defined by the nearest available hourly rainfall station during an event. In some instances, when the hourly rainfall station was non-operational, the second closest site was adopted. Secondly, storm event timings were dictated from the hourly rainfall distributions, including components which were considered pre-bursts and hence excluded from the events per Ladson (2016).

Similarly, the gauged streamflow record had two purposes. The event hydrographs would ultimately function as the calibration reference, while the recession timing indicated the event

conclusion. In all circumstances, the event duration was extended beyond rainfall ceasing to account for catchment transmission delays. The eight modelled extreme rainfall events are listed in Table 3.2. The number of rainfall datasets data available for analysis and the type of rainfall burst experienced are specified. The desired spread in peak flow magnitude between events is also demonstrated.

Table 3.2: Extreme rainfall events used to calibrate RORB model parameters

Start date and time of	Rainfall	Active rainfall	Rainfall	Maximum
rainfall event	duration	station datasets	burst	discharge at
	(hours)	(daily/hourly)	classification	143229A
30 April 1996 17:00	138	10 / 2	Dual	496.6
18 November 2008 10:00	48	13 / 3	Single	255.9
5 January 2011 0:00	168	9/3	Dual	1387.1
25 January 2013 20:00	62	10 / 2	Single	1041.5
30 March 2017 0:00	23	10 / 3	Single	248.7
24 February 2022 09:00	96	6/2	Single	1097.7
11 May 2022 09:00	76	8 / 2	Single	521.1
28 January 2024 0:00	65	8/3	Single	366.7

The total rainfall received at the centroid of each subcatchment during the individual events listed in Table 3.2 was approximated by developing an event-specific isohyetal distribution. From the total point rainfall recorded at each available daily rainfall station, a raster distribution was generated in QGIS by the inverse distance weighted (IDW) interpolation operation, as illustrated in Figure 3.13 for the May 2022 event. Then, the total rainfall at each of the 20 subcatchment centroids was assigned from the identify raster attribute function in QGIS. This process was undertaken for each burst of the dual-burst events.

This method of areal rainfall approximation has several associated strengths and limitations. The influence of a singular point rainfall observation when compared to another station reduces with distance using the IDW interpolation. Therefore, a potentially inaccurate and undetected event rainfall measurement is contained within a localised proximity to the questionable station. The dominant (and assumed more accurate) range of measurements sourced from surrounding stations then are interpolated through the remaining majority of the catchment.

Additional rainfall datasets originating from locations just outside the catchment decorrelation distance of 10 kilometres (but within the extent of the definitive mountain ranges illustrated in

Figure 2.1) were unable to be sourced. Consequently, the IDW interpolation produced unrealistic circular isohyets, predominantly centred about the Townson station. However, the difference between isohyets overlaid on the southern catchment reaches was minimal, especially within the catchment boundary itself. The largest misrepresentation of rainfall distribution was observed to the south of the catchment (beneath the catchment boundary in Figure 3.13). Additionally, considering the centroidal rainfall depth as the subcatchment average is prone to some uncertainty, especially in the larger subcatchment areas where rainfall varies significantly within the singular subcatchment.

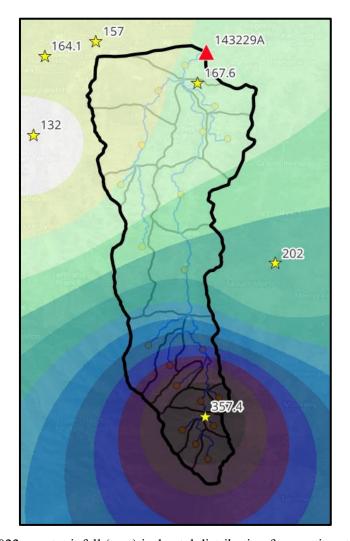


Figure 3.13: May 2022 event rainfall (mm) isohyetal distribution from active station datasets (stars)

The formatted storm event files for calibration of the RORB parameters contained hourly rainfall and runoff observations from available stations; timing definitions of these observations; and event total subcatchment rainfall depths. The storm files and associated graphical isohyetal distributions are attached within Appendix F.

3.4.3 Calibration of the RORB model parameters

The eight storm event files were separately loaded into RORB to calibrate the model parameter k_c for the Laidley Creek catchment file. Calibration of each event was based upon comparing the model calculated runoff at the catchment file outlet against the observed streamflow record from 143229A. The values of k_c and IL depth were adjusted using a trial and error approach until the calculated and observed runoff hydrographs were as aligned as possible. The m parameter value was fixed as 0.8 per the recommendations of Laurenson, Mein and Nathan (2010). The RORB FIT specification was used for calibration, such that CL values were automatically computed by the program for a given IL to produce a flow volume as equal as possible to the observed data. The IL depth considered for each event varied to represent the spread of diverse antecedent conditions captured. For events consisting of two separate bursts, two IL depths were assigned, generally the second was minimal given the preceding burst.

The suitability of the correlation achieved by the calculated hydrograph was evaluated against a set of qualitative and quantitative criteria. Aspects of the hydrograph shape were evaluated for alignment, including:

- The slope and curvature of rising and falling limbs
- The number of discrete local maxima (peaks)
- The initial commencement timing of runoff generation
- The lateral displacement between the observed and calculated hydrographs curves

Quantitative measures of correlation were determined directly within the RORB statistics panel, including the percentage difference in:

- Peak discharge
- Duration of time to peak discharge
- Duration of time to flow volume centroid

Considering multiple events during calibration of the model was necessary to ensure the selected parameters, especially k_c , were an as viably accurate representation of the physical processes that occur in the Laidley Creek catchment as possible. For a singular event modelled in the RORB environment, many different combinations of k_c and IL produce hydrographs with similar attributes and appearances. However, modelling additional events typically demonstrates which parameters are suitably applicable to a variety of events with different attributes including flow duration, peak magnitude and number of flow peaks. Then the singular value corresponding to a model parameter was determined as the measure of central

tendency of the spread of all events, excluding any notable outliers. In this instance, the mean and median values of k_c were considered.

Calibration commenced for each event by adopting the regionalised value of $k_c = 22.48$ as automatically calculated by RORB for a Queensland catchment with an area of 452 km². Then the IL depth was adjusted as much as possible for each event in an attempt to closely align the calculated and actual hydrographs. However, it was evident that the regionalised value of k_c was too low because the upper limit of IL (based upon the total rainfall received) was reached and the flood hydrograph peak was significantly greater than observed, as illustrated in Figure 3.14 for the February 2022 event and Figure 3.15 for the May 2022 event.

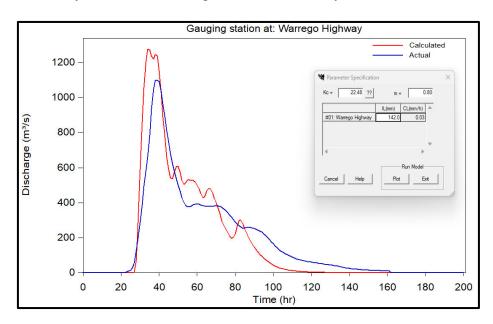


Figure 3.14: Calculated hydrograph (red) for February 2022 event using regionalised kc = 22.48

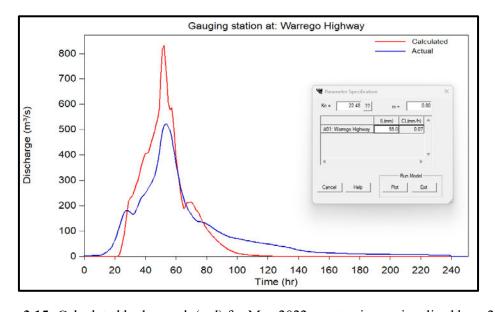


Figure 3.15: Calculated hydrograph (red) for May 2022 event using regionalised kc = 22.48

Therefore, the value of k_c was adjusted as required for each event, as documented in Appendix F and is further discussed below. To ensure adjustments were not completely 'random', the maximum IL such that the CL remained zero or greater was identified. While not the defining parameter for every event, it was a useful indicator to narrow the range suitable k_c values.

It is recognised that the order the events were considered had the potential to somewhat affect the calibration outcomes. This is an accepted limitation of the manual calibration method provided in RORB, which does not have the capacity to automatically compute optimal parameter values for a single event, let alone across multiple events considered simultaneously.

The lateral displacement between the calculated and observed hydrographs was another notable aspect of the calibration. In some instances, the observed hydrograph shape was closely replicated by the calculated hydrograph, but the timing difference between them was significant, as exemplified in Figure 3.16 for the January 2013 event.

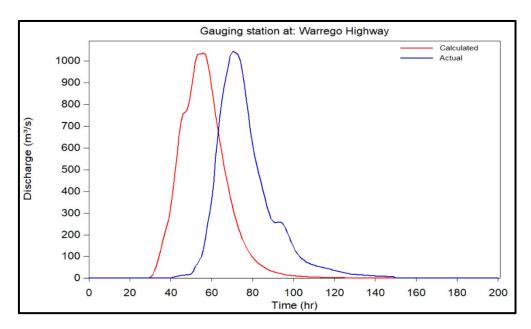


Figure 3.16: Lateral separation between calculated and observed hydrographs for January 2013 event

The potential for timing issues in catchments with very flat lower reaches was addressed in section 7.2.6 of the RORB User Manual by Laurenson, Mein and Nathan (2010), who recommended the insertion of event specific translations into the model to resolve such issues. For the events that experienced timing issues, a specified time delay was incorporated along a final arbitrary stream reach of zero length in the catchment file, which resulted in minimal separation between the hydrographs.

The majority of the events analysed for calibration were deemed suitable and the spread of parameter values was narrow. The calibration obtained for each event is expanded on below.

3.4.3.1 Calibrated events with closely replicated peak flows

a) November 2008: $k_c = 34 \& IL = 77 \text{ mm}$ (with 4 hour delay translation)

The singular discharge peak magnitude, shape and timing was closely correlated to observed data. The peak discharge was the second lowest analysed, potential for event k_c to be less representative of higher magnitude flows (rainfall was atypically low in upper catchment).

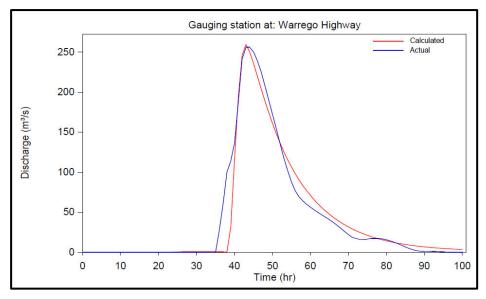


Figure 3.17: Calibrated hydrograph for November 2008 event

b) January 2011: $k_c = 26 \& IL = 9 mm$

The final runoff peak closely aligned with the observed data. The magnitude and shape of the preceding peaks was generally replicated (except the peak immediately prior to the event maxima) however timing was approximately 10 hours earlier than observed.

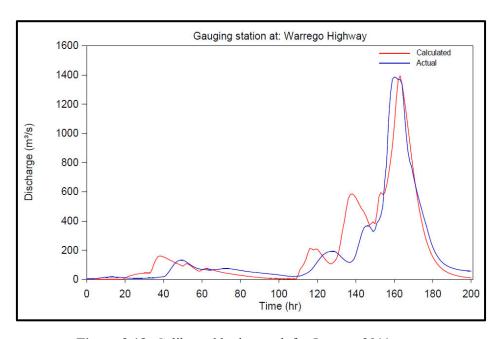


Figure 3.18: Calibrated hydrograph for January 2011 event

c) January 2013: $k_c = 30 \& IL = 60 \text{ mm}$ (with 20 hour delay translation)

The calibration achieved for the January 2013 event was the closest of the three significant events (peak flows greater than 1000 m³/s). The hydrograph shape, especially the slopes of the rising and falling limbs, corresponded very closely to the observed data, however the timing of the singular peak was slightly delayed.

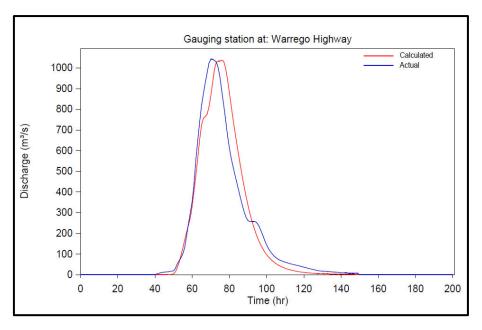


Figure 3.19: Calibrated hydrograph for January 2013 event

d) March 2017: $k_c = 23 \& IL = 107 \text{ mm}$ (with 4 hour delay translation)

This event had a very similar magnitude to the November 2008 peak, and the hydrograph peak magnitude, shape and timing was also very closely correlated to the observed data. Rainfall in the lower subcatchments was low. Potentially k_c is also less representative of higher flows.

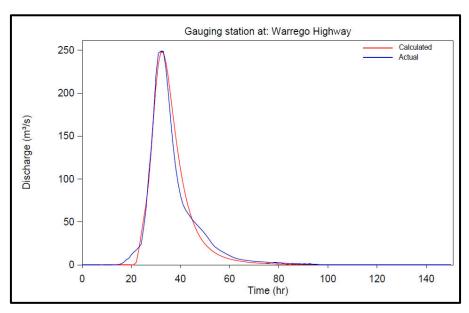


Figure 3.20: Calibrated hydrograph for March 2017 event

e) February 2022: $k_c = 29 \& IL = 142 mm$

The timing and magnitude of the peak discharge was well calibrated. The shape of the 100 hour recession curve was somewhat replicated. The calculated hydrograph overestimated several observed local maxima followed by an underestimation of the final falling limb.

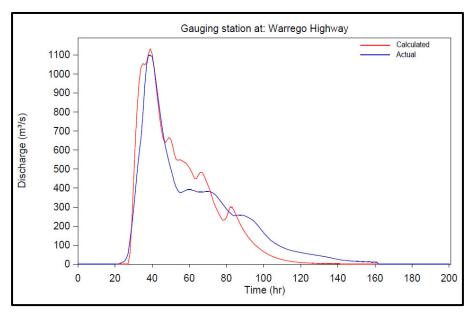


Figure 3.21: Calibrated hydrograph for February 2022 event

f) January 2024: $k_c = 28 \& IL = 39 \text{ mm}$ (with 12 hour delay translation)

The magnitude and shape of the larger hydrograph peak was suitably approximated by the calculated hydrograph. The timing of both distinct peaks was also replicated. The smaller peak was slightly overestimated while the first recession curve was subsequently delayed. A dual burst rainfall model may have improved the calibration, however it was satisfactory regardless.

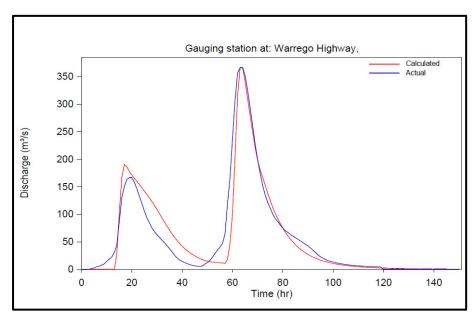


Figure 3.22: Calibrated hydrograph for January 2024 event

3.4.3.2 Problematic or uncertain calibration events

a) May 1996: $k_c = 26 \& IL = 100 \text{ mm}$ (with 15 hour delay translation)

The oldest event analysed had two distinct flood peaks of similar magnitude that remained effectively stable for 16 to 18 hours each as a result of two separate rainfall bursts. The extent of the observed peaks were unable to be replicated by any combination of k_c and IL. While the magnitude of the first peak was somewhat similar, the second calculated peak was significantly higher when the remainer of the hydrograph was relatively well calibrated. The error between the calculated and observed peak discharge resulting from the second rainfall burst was 26.5%.

Upon inspection of the hourly streamgauge record, the peak flow readings were graded as code 60 'estimates' which indicated the readings were unreliable. This explanation is reasonable given the flows experienced during the May 1996 event were the highest observed since the site had opened in 1991. Hence no previous gaugings existed of the maximum flows experienced in May 1996 and the rating curve was poorly defined. It is plausible the actual streamflow hydrograph replicated the calculated hydrograph however this cannot be proven or demonstrated.

Otherwise, the complex hydrograph shape and timing encompassing two significant peaks and multiple local maxima was rather well replicated and the k_c value of 26 will be considered as reasonable with caution.

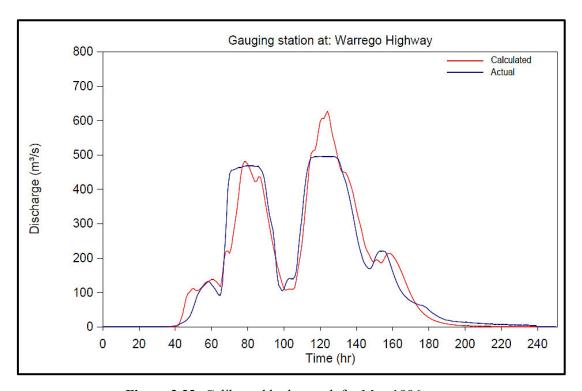


Figure 3.23: Calibrated hydrograph for May 1996 event

b) May 2022: $k_c = 41 \& IL = 0 mm$

Compared to every other event, the May 2022 event required a significantly higher value of k_c to obtain a satisfactory calibration. A lengthy recessional tail period was observed in the hydrograph as continuous yet insignificant rainfall was experienced in the days following the maximum recorded discharge.

For a constant k_c value, as the IL depth was reduced, the calculated hydrograph peak also reduced and was more closely aligned towards the observed data. However, for k_c values considered for other events (between 25-35), the calculated maximum discharge still significantly exceeded the observed record when IL was minimised (set to zero) and the hydrograph shape was inaccurate. Therefore, the value of k_c was increased until both the maximum discharge and hydrograph shape were closely aligned to the observed data. This meant a k_c value of 41 was adopted for this event, which when applied to the remaining seven events, resulted in multiple unsatisfactory calibrations.

Hence this event was considered unrepresentative of the general catchment response to extreme rainfal, and the value of $k_c = 41$ was deemed an outlier.

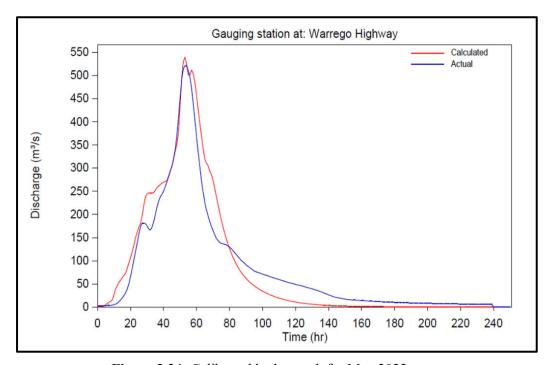


Figure 3.24: Calibrated hydrograph for May 2022 event

Because eight events covering an adequate range of durations and flow magnitudes were analysed, if one or two events were deemed unsuitable, a sufficient number of events were still calibrated to verify the accuracy of the chosen parameter values.

The selected event parameter values are summarised in Table 3.3. The mean and median value of k_c was calculated as 27. Because the same value was obtained for both measures of central tendency, $k_c = 27$ was settled upon as the calibrated value for the Laidley Creek catchment.

Table 3.3: Best-fit and average parameters for RORB calibration events

Event	k_c	IL (mm)	CL (mm/h)
			[computed]
April - May 1996	26	100	2.66
November 2008	34	77	9.78
January 2011	26	9	3.18
January 2013	30	60	0.94
March 2017	23	107	0.65
February 2022	29	142	0.03
May 2022 *	41	0	1.09
January 2024	34	39	5.52
Mean	27	* Indicates omi	tted outlier
Median	27		

3.4.4 Design discharge simulations for current and future climate conditions

Following successful calibration of the catchment k_c parameter, the design discharge estimates were obtained using the runoff routing simulation method. The RORB DESIGN specification was used to estimate design event discharge quantities in conjunction with several design storm files. These files consisted of the ARR Datahub .txt file, the areal temporal pattern .csv file and an IFD .csv file for each climate scenario, as attached in Appendix D, E and C respectively.

Another limitation of the RORB package was that a singular set of IL and CL values had to be specified for the design simulation, despite Table 3.3 demonstrating a wide range of values were used to calibrate the catchment response to different antecedent conditions. Although a stochastic distribution about an assigned IL value can be enabled using the Monte Carlo simulation method, a singular known IL depth is still required to begin this process. The regionalised median IL and CL depths of 25 mm and 1.40 mm/hr respectively, as specified in the ARR Datahub file, were automatically assigned as the RORB simulation losses.

As the catchment area was larger than 75 km², an areal temporal pattern file was specified in the RORB input. The patterns were sorted by catchment area at discrete intervals. In this instance, RORB considered the 500 km² areal temporal patterns as closest to the catchment area of 452 km². Because areal patterns were used, the number of modelled events for each AEP was reduced from 25 IFD durations to ten events between 12 and 168 hours. Two alternate simulation methods of design discharge estimation were evaluated: Ensemble simulation and Monte Carlo (MC) simulation. The advantages and limitations of each approach are discussed below:

3.4.4.1 Ensemble simulation method

The Ensemble simulation method separately routed catchment runoff from ten temporal distribution patterns for each design event AEP between 50% and 1%, and duration between 12 and 168 hours. The set of ten simulations for the 5% AEP, 48 hour event are presented as an example in Figure 3.25. The median peak discharge from the set of ten is considered the peak flow estimate. Once repeated for all event durations, the maximum-median peak flow across all durations was considered the critical discharge estimate for a specific AEP.

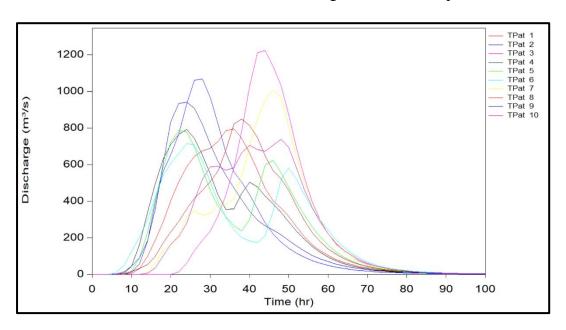


Figure 3.25: 5% AEP, 48 hour Ensemble simulation event hydrographs

The main advantage of the Ensemble method is the ability to easily demonstrate the range and extent of uncertainty in the results (through a quartile box and whisker plot). Operationally, the method had near-instantaneous simulation processing times and the generated results were able to be reproduced on separate occasions. However, the ensemble method is restricted to considering only the ten current temporal patterns and a singular IL depth.

3.4.4.2 MC simulation method

The RORB MC simulation method uses stratified sampling to derive a runoff frequency curve across a defined range of probabilities for each storm duration. Compared to the Ensemble method, the MC method maintains probability neutrality of the inputs by considering a stochastic distribution of the naturally variable processes that influence runoff generation. Therefore, the MC method was selected as the preferred simulation method for this project.

The frequency range was uniformly divided into intervals from which stochastically sampled rainfall depths were routed through the catchment file to generate a set of peak discharge values (Laurenson, Mein & Nathan 2010). The simulated results of each interval were statistically analysed to estimate the probabilistic runoff frequency curve. This procedure was repeated for each specified event duration to constitute one simulation, as illustrated in Figure 3.26. The event corresponding to the maximum discharge for a given AEP was considered the critical duration event.

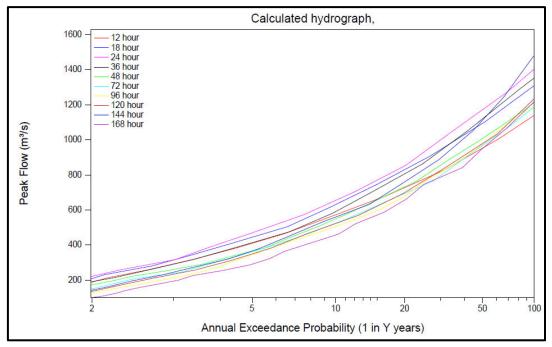


Figure 3.26: Sample set of runoff frequency curves generated from MC simulation

The model default settings of 20 stochastically sampled rainfall and IL depths simulated across 50 event frequency intervals to produce 20 flows per interval were retained as recommended in section 8.5 of the RORB User Manual by Laurenson, Mein and Nathan (2010). The nature of an MC simulation means that the results of each execution are unique. Therefore, to assess the variability of the outputs, 10 trial simulations were executed for each of the 7 climate scenarios (current and 6 future) to determine the median peak flow associated with the design

AEPs. Each simulation took between 30 and 60 seconds to execute, so modelling was restricted to 70 simulations as a balance between output result accuracy and available resources. The results are presented in chapter 4 of this report.

3.5 Bayesian Flood Frequency Analysis

Technical guidelines published by DTMR stipulate that where a streamgauge is situated in close proximity to state controlled road infrastructure, "design discharges should be calibrated (with) FFA at the gauge location if sufficient recorded data exists" (2024c). The Laidley Creek at Warrego Highway (143229A) gauge has a 33 year streamflow record, of which the AM discharge series is plotted in Figure 3.27. The relatively short gauge operating period was identified as a significant limitation of the results of an FFA.

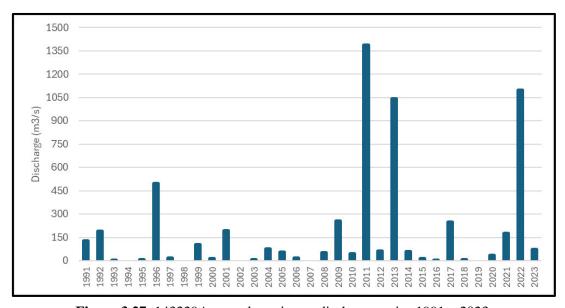


Figure 3.27: 143229A annual maximum discharge series 1991 – 2023.

The RMC-BestFit (RMC) Bayesian estimation and distribution fitting software package developed by the US Army Corps of Engineers and Institute for Water Resources (2024) was used to undertake an FFA at 143229A. While the TUFLOW FLIKE package is more commonly used to determine flood frequencies from historic data, including throughout Book 3 of ARR 2019, Ball et al. (2019) clarified that alternative software packages were acceptable for use, if applied appropriately to the modelling scenario. An annual licensing fee is associated with the TUFLOW package, therefore RMC was chosen for this study as a free option while still possessing sufficient modelling capabilities. The RMC package utilises an interactive platform to conduct a three-stage modelling approach. The software quickly processes large quantities of input data to produce both graphical and tabulated outputs. RMC is an ideal approach to efficiently undertake a FFA for this research project.

3.5.1 Initial plotting position from annual maximum series

The maximum discharge for each water year was required for analysis. The commencement of a water year corresponds to the calendar month with the lowest average flow across the 33 year record. For the 143229A gauge, August had the lowest average monthly flow of 0.1985 m³/s. Therefore, the preceding August to the current July was considered one water year. For example, the period spanning August 2008 to July 2009 was considered the 2009 water year.

Once the AM series was inputted to the RMC package, the Cunnane plotting position of each discharge was automatically computed. Alternative plotting position parameters can also be adopted. The Multiple Grubbs-Beck low outlier test is an optional measure to identify and exclude potentially influential low flows (PILFs), as recommended in Book 3, Section 2.8.6 of ARR 2019 (Ball et al. 2019). The test is required to enable the selection of some distributions that are not compatible with PILF gaugings. The rationality of this results yielded by executing this test will be evaluated given the relatively low operational period of the streamgauge.

The RMC input data interface and the subsequently generated plotting position graph of discharge against AEP for gauge 143229A are shown in Figure 3.28.

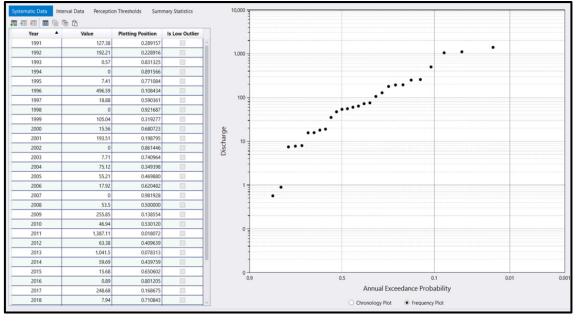


Figure 3.28: 143229A Laidley Creek at Warrego Highway AM series input and plotting position graph generated in RMC-BestFit

3.5.2 Fitted distribution functions

The second component of the RMC analysis procedure involves fitting up to 13 distribution functions to the plotted input data by the maximum likelihood estimation (MLE) method. The location, shape and scale factors, the statistical moments, and the yield value for various AEPs

corresponding to each function are presented in an output results table. The closeness of each function along the AEP domain was qualitatively inspected through a visual comparison of the plotted distributions against the input data. RMC computed three objective functions as a quantitative measurement of fit for each distribution, namely:

- Akaike Information Criteria (AIC)
- Bayesian Information Criteria (BIC)
- Root-Mean Squared Error (RMSE)

The lowest value for each objective function represents the distribution with the best fit to the input data (US Army Corps of Engineers & Institute for Water Resources 2024). Considering these metrics, in combination with a qualitative judgement of the closeness of the fit, the most suitable distribution function is identifiable. In this instance, the LP3 distribution produced the lowest RMSE of 109.40, the fifth lowest AIC and BIC ratings, and by visual inspection, appeared to effectively represent the full spread of gaugings. As such, the findings of Rahman et al. (2013) regarding the performance of the LP3 distribution are validated for the Laidley Creek catchment. The Weibull and Gamma distributions also perform well, ranking highly in terms of AIC and BIC rating; and second and third respectively in terms of RSME. Other distributions failed to perform as highly against all metrics. Because the Weibull and Gamma distributions have comparable performance to the LP3 distribution, as illustrated in Figure 3.29, all three distributions will be individually modelled using Bayesian parameter estimation.

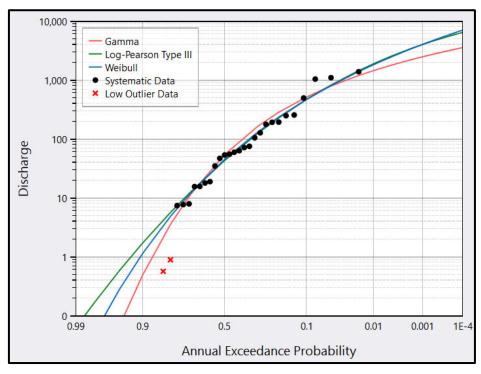


Figure 3.29: LP3, Gamma and Weibull distributions fitted to AM series with outlier flows identified

3.5.3 Bayesian parameter estimation

The final phase in the RMC modelling procedure is the Bayesian parameter estimation analysis, which is undertaken through a reiterative Markov Chain MC simulation (US Army Corps of Engineers & Institute for Water Resources 2024). The input data from step one, the three distributions from step two (once-at-a-time), and the desired parameter distribution were considered by RMC to generate 10,000 distribution parameter sets in approximately eight seconds, which converge to parametric distributions for mean, skew and standard deviation instead of a singular value of each in step two. The parametric and quantile distributions can be refined based on regionalised flood data (Smith & Doughty 2020), however, in the absence of such data, these options were kept within the default domains of the software package.

The parameters with the most frequent recurrence were identified by RMC as the posterior mode parameters, which form the posterior mode function (in the same arrangement as the parent distribution function). The uncertainty in the generated parameter sets is characterised by the 95% confidence intervals about the posterior function, as shown in Figure 3.30. It was observed that as AEP increased beyond the interpolated zone of gauged data, the confidence intervals increased significantly, emphasising the uncertainty of the estimations.

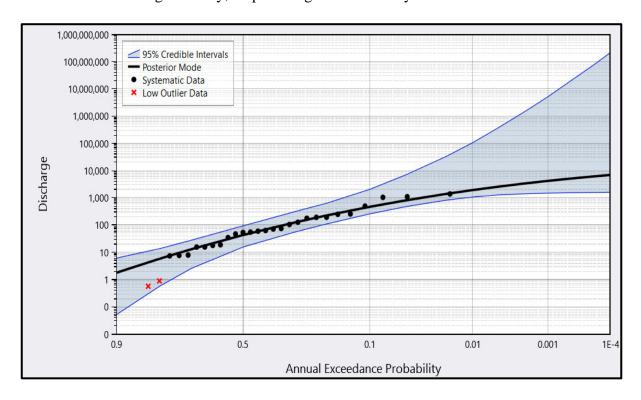


Figure 3.30: Posterior mode and confidence intervals of the Bayesian parametric estimation method for the LP3 distribution

3.5.4 FFA design discharge results

The discharges corresponding to the posterior mode frequency curve for each fitted distribution are presented in Table 3.4, while full outputs from RMC are included in Appendix I. It is intended that the simulated runoff routed discharges for the current climate scenario will be compared to the FFA results.

Table 3.4: Posterior mode FFA discharges (m³/s) for suitable distributions

AEP (%)	LP3	Gamma	Weibull
50	42.91	54.40	45.04
20	223.55	282.02	233.65
10	465.77	514.93	470.48
5	802.04	775.95	786.87
2	1386.27	1147.58	1325.57
1	1925.76	1442.20	1823.29

It is likely the Gamma function estimates are most representative of the true catchment characteristics, however this prediction will be assessed in the simulation sanity verification. It was intended for the FFA discharges to be compared to the current climate discharges estimated from the RORB simulations. However, the accuracy of the FFA method is limited by the 33 year streamgauge record, which inherently is reflective of recent events rather than the extended site flood history. A longer spanning AM series record would provide a more accurate representation of likely flood quantiles – a 33 year record has limited accuracy when considering a 1 in 100 year event. These inaccuracies are demonstrated by the significant discrepancies between the design discharges of the three functions.

3.6 Regression methods of peak flow estimation

The RFFE and P&W methods introduced in the literature review are simple peak flow estimation techniques which provide comparisons to the simulated routed discharges. The application of these methods to the Laidley Creek catchment are described below.

3.6.1 RFFE method

The RFFE method is incorporated into an online modelling tool titled RFFE Model – 2016 Release Version (Rahman et al. 2015). The RFFE model was used to generate discharge estimates for the 50%, 20%, 10%, 5%, 2% and 1% AEPs. The following user stipulated catchment attributes were required for the model to generate the estimates:

- Site name or ID
- Outlet latitude and longitude (in decimal degrees)
- Centroid latitude and longitude (in decimal degrees)
- Area (in km²)

An initial RFFE computation was executed using the corresponding catchment parameters previously determined. The centroid location was determined in QGIS from the polygon centroid execution. RFFE approximated the catchment shape as an ellipse configured about the centroid and outlet locations, as shown on an interactive map reproduced in Figure 3.31. The blue shading indicated that the shape factor of the catchment is appropriate (rather than being too narrow or irregularly shaped) such that the method is within its limits of accuracy.

The model generated discharge estimates for the six design AEPs based upon regionalised LP3 distribution parameters (Rahman et al. 2015). However, the regionalised parameters are likely unrepresentative of the Laidley Creek catchment because of its atypical topographical attributes (compared to the surrounding East Coast region) that cause significant variance in the distribution and quantity of rainfall received. To overcome this issue, the RFFE 'Nearby' spreadsheet containing details of surrounding gauged catchments was downloaded and inspected. The distances between the user inputted catchment (Figure 3.31) and surrounding gauged catchments used to originally formulate the RFFE method are included in this spreadsheet. The Laidley Creek catchment at the 431229A station was one such catchment used as indicated by the nearest gauge distance of 0.21 km.

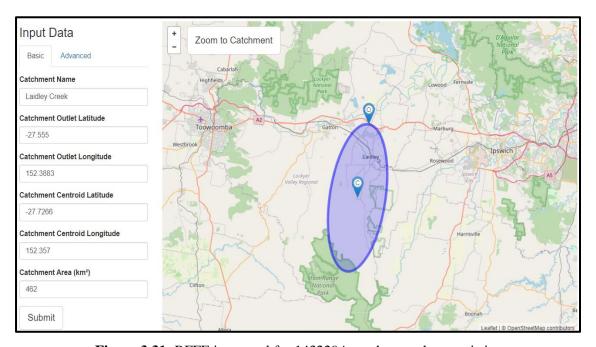


Figure 3.31: RFFE input tool for 143229A catchment characteristics

Discharge data sourced from a gauged catchment and adopted for the development of the RFFE method was subjected to multiple-stage processing and validation measures (Rahman et al. 2015). Therefore, this data is considered significantly more reliable than the equivalent regionalised model output. On the basis of these considerations, the estimated design discharges for the catchment were extracted directly from the existing gauged dataset and are listed in Table 3.5

Table 3.5: 143229A discharges estimates from RFFE method

AEP (%)	Discharge (m ³ /s)	
50	47.15	
20	209.72	
10	402.95	
5	661.58	
2	1110.63	
1	1535.72	

3.6.2 P&W method

The P&W method of discharge estimation for rural Queensland catchments reduced the number of input parameters required (in comparison to RFFE) to only the catchment area and the 2% AEP, 72-hour duration design rainfall intensity at the catchment centroid. The peak discharge estimates for six ARIs ranging from 2 to 100 years at the 143229A gauge for the current climate scenario were calculated from the P&W equations listed as Equation 3.1-3.6:

$$Q_2 = 0.122 \times A^{0.757} \times i72h50y^{1.588}$$
 (3.1)

$$Q_5 = 0.664 \times A^{0.709} \times i72h50y^{1.301}$$
 (3.2)

$$Q_{10} = 1.419 \times A^{0.682} \times i72h50y^{1.174}$$
 (3.3)

$$Q_{20} = 2.547 \times A^{0.673} \times i72h50y^{1.074}$$
 (3.4)

$$Q_{50} = 4.731 \times A^{0.656} \times i72h50y^{0.968}$$
 (3.5)

$$Q_{100} = 7.031 \times A^{0.644} \times i72h50y^{0.899}$$
 (3.6)

Where the catchment area $A = 462 \text{ km}^2$ and the current 72 hour, 2% AEP design rainfall intensity i72h50y at the Laidley Creek catchment centroid = 4.733 mm/hr (equivalent event rainfall depth from BOM-LIMB IFD envelope of 340.8 mm).

Table 3.6: 143229A discharge estimates from P&W method for current climate

ARI (years)	AEP (%)	Discharge (m ³ /s)
2	39.25	149.83
5	18.13	388.82
10	10	577.93
20	5	840.29
50	2	1192.59
100	1	1479.09

The numeric form of Equations 3.1-3.6 meant that the increased rainfall intensities for the future climate scenarios could be directly substituted to predict the future design discharges. The updated i72h50y intensities ranged between 5.207 and 6.485 mm/hr. This approach assumed that the distributions and relationships on which the P&W were originally derived from remain applicable for the increased rainfall intensities of future scenarios, however, this assumption has yet to be validated. Therefore, the results obtained were considered as limited provisional indications. The future climate scenario discharge estimates from the P&W equations are listed in Tables 3.7 and 3.8, and as projected, reflected significant increases.

Table 3.7: 143229A discharge estimates (m³/s) from P&W equations - RCP 4.5 scenarios

AEP (%)	2030	2050	2090
39.25	174.35	184.50	200.20
18.13	440.22	461.11	493.02
10	646.45	674.07	716.02
5	930.99	967.31	1022.23
2	1308.01	1353.91	1423.01
1	1611.59	1664.05	1742.79

Table 3.8: 143229A discharge estimates (m³/s) from P&W equations - RCP 8.5 scenarios

AEP (%)	2030	2050	2090
39.25	176.87	194.86	247.05
18.13	445.43	482.22	585.69
10	653.34	701.86	836.43
5	940.06	1003.72	1178.43
2	1319.49	1399.77	1617.59
1	1624.72	1716.33	1963.07

3.7 Research methodology evaluation

The finalised research methodology represents a culmination of several refinements as each modelling aspect was developed, whilst ensuring the scope remained viable despite time and resource constraints. The availability and applicability of sourced data was delicately evaluated against modelling recommendations introduced in the literature review, including ARR 2019, DTMR guidelines and the RORB User Manual, to ensure the finalised model represented the physical hydrologic properties of the Laidley Creek catchment as accurately as possible. Any necessary assumptions and simplifications of these complex characteristics were justified. Alternative modelling approaches were investigated at several stages throughout the methodology, including the automatic catchment delineation method, the continuous catchment simulation method of runoff estimation, and the Ensemble event-based method. In each instance, justifications to support the selection of the finalised method were provided.

Many aspects of the methodology are typical of best practices currently used in industry. This project considered the most recent climate change projections and guidance for hydrologic modelling provided by ARR 2019 and DTMR. Specifically, the IFD adjustment factors for the different RCP scenarios (Appendix C) and median surface temperature projections were released on the ARR Datahub on 27 August 2024 to coincide with the significantly revised ARR 2019 Book 1: Chapter 6 titled *Climate Change Considerations*. Similarly, the CCNHRA and associated engineering policies were published by DTMR in June 2024. With continuing advancements in research and understand, it is practically certain these methods, guidelines and projections will be revised again in the future.

3.8 Summary

This research methodology presented a five-stage approach to quantify the implications of climate change scenarios on peak discharge hydrology within the Laidley Creek catchment. The reviewed background literature was synthesised to correspond with the project objectives and industry best-modelling practices to guide model development. Input data was obtained from relevant sources to formulate the catchment and storm event files. Eight historic peak rainfall runoff events were analysed to calibrate the RORB routing model parameter $k_c = 27$ for the Laidley Creek catchment. Discharge estimation methods were documented, including independent analytical and statistical methods, while the Ensemble and MC simulation methods were compared. The MC method was selected to generate the median design discharge estimates for the current and future climate scenarios, as presented in Chapter 4.

4. Results

The design simulation results for the current and future climate IFDs are discussed in this chapter. A sensitivity analysis of the simulated model parameter values was also completed.

4.1 Design discharge simulation for current climate conditions

The design routing simulation outlined in the previous chapter was first undertaken for current climate conditions by specifying the LIMB-BOM IFD envelope as the input design rainfall set. The calibrated catchment parameters $k_c = 27$, m = 0.8 (assumed fixed) and regionalised losses IL = 25 mm and CL = 1.4 mm/hr were specified. 10 sample simulations were executed, of which the median, minimum and maximum peak discharges for the six design AEPs between 50% and 1% are listed in Table 4.1. The median flows were adopted as the design discharge estimates. The full sample simulation outputs are attached in Appendix G.

Table 4.1: Simulated design discharges for current climate scenario (2020 envelope)

	Discharge (m ³ /s)		
AEP (%)	<u>Median</u>	Minimum	Maximum
50	223.0	210.4	233.7
20	467.1	457.6	482.2
10	661.9	649.4	676.5
5	868.4	833.7	888.0
2	1164.5	1149.4	1178.0
1	1411.4	1381.3	1479.9

The critical duration of each design AEP across the ten samples was the 24 hour event, except the 1% AEP, where for one sample simulation, the 144 hour event corresponded with the critical flow. In most sample instances the 144 hour event produced the second highest peak discharges.

As an indication of the uncertainty of the results, the inter-quartile and full ranges of simulated peak discharges for each AEP were represented as a proportion of the median flow. The full range varied between 2.4% for the 2% AEP event and 10.4% for the 50% AEP event. Similarly, the inter-quartile range varied between 1.3% for the 2% AEP event and 3.1% for the 50% AEP event. The extent of uncertainty associated with the 10 sample simulations for each design AEP is illustrated in Figure 4.1.

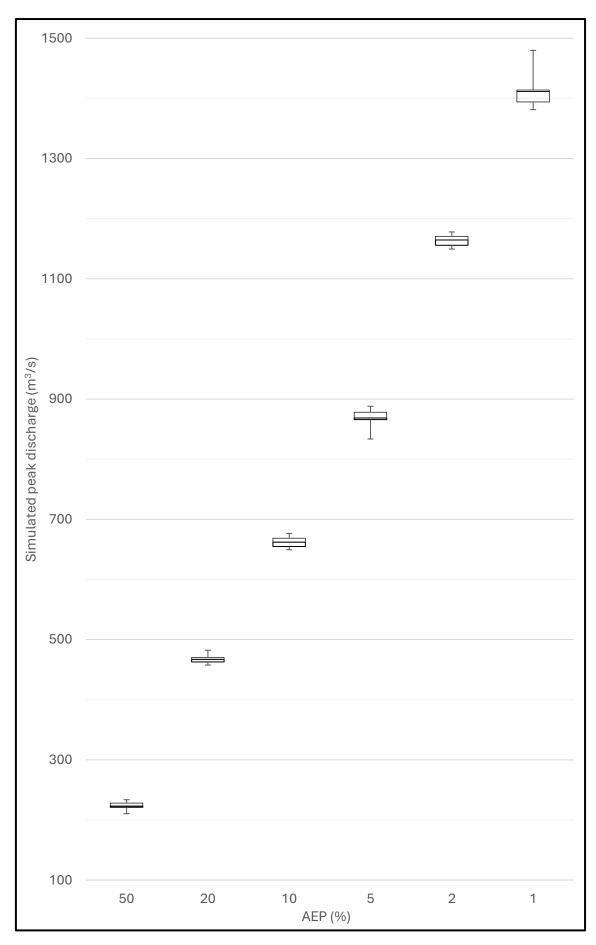


Figure 4.1: Distribution of MC simulated peak discharges for current climate scenario

4.2 Design discharge simulation for future climate scenarios

The design routing simulations were repeated for the six future climate projections each corresponding to a revised IFD file. The median peak discharges estimates for the RCP 4.5 scenarios are listed in Table 4.2 while estimates for the RCP 8.5 scenarios are listed in Table 4.3. Full simulation outputs are attached in Appendix G. The percentage increase compared to the corresponding median discharge for current conditions per Table 4.1 is also indicated.

<u>Table 4.2: Simulated median discharge estimates for RCP 4.5 scenarios and percentage increase compared to current climate simulated median</u>

AEP	Discharge	Increase	Discharge	Increase	Discharge	Increase
(%)	2030 (m ³ /s)	(%)	2050 (m ³ /s)	(%)	2090 (m ³ /s)	(%)
50	281.1	26	296.2	33	330.0	48
20	547.5	17	580.1	24	630.4	35
10	767.4	16	799.5	21	865.8	31
5	992.7	14	1044.5	20	1116.7	29
2	1315.1	13	1383.0	19	1483.7	27
1	1587.9	12	1659.5	18	1772.3	26

For the RCP 4.5 scenarios, the 24 hour event was the dominant critical duration. However, six of the 30 samples indicated that the critical duration of the 1% AEP event was 144 hours.

<u>Table 4.3: Simulated median discharge estimates for RCP 8.5 scenarios and percentage</u> increase compared to current climate simulated median

AEP	Discharge	Increase	Discharge	Increase	Discharge	Increase
(%)	2030 (m ³ /s)	(%)	2050 (m ³ /s)	(%)	2090 (m ³ /s)	(%)
50	280.3	26	320.0	43	431.5	94
20	555.5	19	614.9	32	780.9	67
10	779.5	18	842.5	27	1056.5	60
5	1010.4	16	1086.9	25	1341.2	54
2	1340.7	15	1446.6	24	1762.8	51
1	1612.4	14	1736.0	23	2096.3	49

For the RCP 8.5 scenarios, the 24 hour event was also the dominant critical duration. However, the critical duration of two of the 30 samples for the 1% AEP event was 144 hours. The 18 hour event was deemed critical for 16 of the 90 sample discharges across the 50%, 20% and 10% AEP events, which was progressively more prevalent by 2090.

Several trends were observed in the percentage increase in the median peak discharge of the Laidley Creek catchment compared to the current (2020 envelope) simulations. This percentage increase was larger across the RCP 8.5 events, which ranged between 14% and 94%, compared to the RCP 4.5 events, which ranged between 12% and 48%.

Rarer events (lower AEP) were observed to have a reduced percentage increase compared to more frequent events. The 50% AEP events experienced a significantly larger percentage increase compared to all other design events. Events of an earlier projection year had a reduced percentage increase compared to the longer term projections. In summary, the observed trends generally align with the relationships between the individual IFD adjustment factors included in Appendix B.

The uncertainty in the results was also considered at the design event level across the current and future climate simulations. The increased simulated peak discharge for the 1% AEP event until 2090 as a result of the RCP 8.5 emissions scenario is illustrated in Figure 4.2. The interquartile and full ranges about the median discharge of each year are also shown to demonstrate the uncertainty of the sample outputs. Similar figures for each design AEP and RCP scenario were produced and are attached in Appendix G.

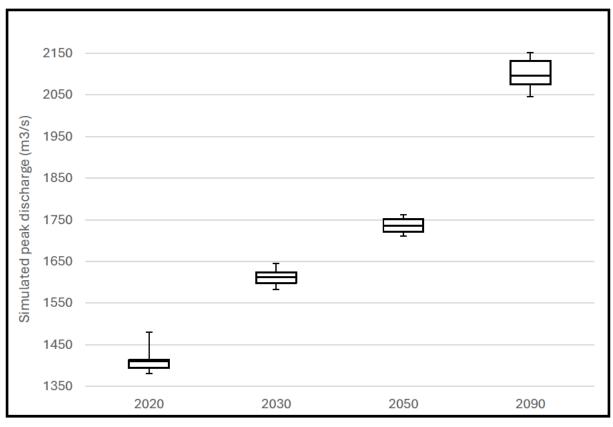


Figure 4.2: Distribution of sample 1% AEP event discharges for current IFDs and RCP 8.5 scenarios

A comparison of the RCP 4.5 and 8.5 emissions scenarios was also undertaken for each design event frequency. The increased peak discharges for the 1% AEP event until 2090 from both the RCP 4.5 and 8.5 scenarios (assuming a linear rate of change between simulated years) are shown in Figure 4.3. The full range of each sample set of outputs constituting a simulation is also shown. Similar figures for each design AEP were also produced and are attached in Appendix G.

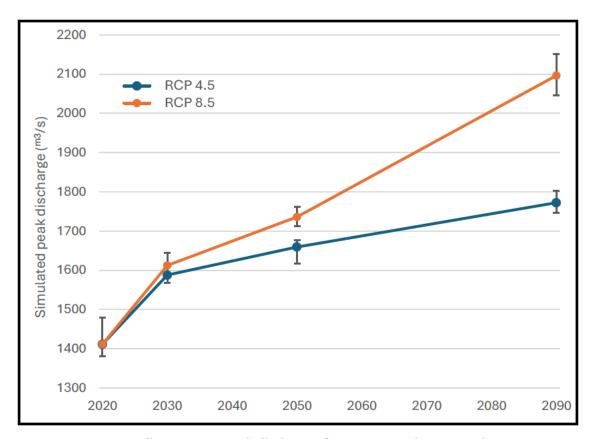


Figure 4.3: Median 1% AEP peak discharges for RCP 4.5 and 8.5 scenarios to 2090

Figure 4.3 shows that a similar rate of increase is projected until 2030 for the 1% AEP event under both emissions scenarios. However, following 2030, peak discharges simulated for the RCP 8.5 scenario increased significantly more than those of the RCP 4.5 scenario. A similar trend was observed for the 2% AEP event.

No apparent trend was noted in the extent of sample uncertainty across Figures 4.1, 4.2 or 4.3.

4.3 Sensitivity analysis of the model parameters

An analysis of the model sensitivity was performed through a once-at-a-time adjustment of the input parameters. A parameter value was separately increased or decreased while the original values of the remaining parameters were preserved. The m parameter was assumed a fixed value of 0.8, therefore six separate adjustments were made. Five sample simulations were

executed for each adjustment, resulting in a total of 30 additional sample simulations being completed. The resultant changes in the median peak discharge of the design AEPs for the current climate model (Table 4.1) were evaluated in Table 4.4. The full outputs of the sensitivity analysis are included in Appendix H.

Table 4.4: Sensitivity analysis of model parameters – current climate simulation

RORB input parameter	Value for simulation	Once-at-a- time adjustment	Evaluation of model response to individual parameter adjustment
k_c	27	± 2.7 [10%]	 +10%: Average median discharge decreased by 7.4%. 1% AEP critical duration increased to 144 hours in four samples, 36 hrs in one sample. 24 hours for other AEPs. - 10%: Average median discharge increased by 7.9%. 50% AEP critical duration decreased to 18 hours in one sample. 24 hours for all other AEPs.
m	0.8	Assumed fix	xed value for modelling of natural streams and channels per Laurenson, Mein and Nathan (2010).
IL	25 mm	± 5 mm [20%]	 +20%: Average median discharge decreased by 5.8%. 1% AEP critical duration increased to 144 hours in one sample, 24 hours for all other AEPs. - 20%: Average median discharge increased by 3.8%. 20% AEP critical duration decreased to 18 hours in one
			sample. 24 hours for all other AEPs.
CL	1.4 mm/hr	± 0.28 mm	+20%: Average median discharge decreased by 4.6%. 1% AEP critical duration increased to 144 hours in one sample, 24 hours for all other AEPs.
		[20%]	- 20%: Average median discharge increased by 5.1%. 1% AEP critical duration increased to 144 hours in one sample, 24 hours for all other AEPs.

The sensitivity analysis demonstrated that the catchment storage parameter k_c was the most influential parameter on the results simulated, with the largest changes in output median discharge experienced when k_c was adjusted. The IL and CL depths had less impact on the model output despite being adjusted by a greater proportion of the originally simulated value. Adjustments of the parameters also varied the critical duration of some samples however the 24 hour event remained the dominant duration.

5. Discussion

This chapter examines the simulated design discharges in greater detail and considers the outputs in terms of the physical hydrologic processes of the Laidley Creek catchment. The validity of the simulated results was evaluated through comparisons to independent estimation techniques. The limitations associated with the model inputs, assumptions and capabilities were referenced to discuss the output limitations. Finally, potential improvements are suggested before the outputs are incorporated within a future 2D hydraulic investigation of the Warrego Highway bridge crossing of Laidley Creek.

5.1 Independent sanity validation of the results

The independent methods of discharge estimation introduced in the literature review and applied to the Laidley Creek catchment in the methodology were used as a comparison for the current and future simulated design discharges. It was recognised that results obtained from these methods are simply estimates, and only provide an indicative representation of catchment hydrologic characteristics.

5.1.1 Comparison of current design discharge simulation

FFA, RFFE and P&W methods were three independent approaches used to compare the simulation results of Table 4.1 for the current climate IFDs. The corresponding estimates obtained by each independent method and the percentage difference compared to the median simulated design discharges are listed in Table 5.1.

Table 5.1: Comparison of current climate simulation with independent estimates (m³/s)

AEP	Median Discharge	FFA m	ethod	RFFE m	ethod	P&W m	ethod
(%)	(Table 4.1)	Discharge	% Diff.	Discharge	% Diff.	Discharge	% Diff.
50	223.0	54.4	75.6	47.2	-78.9	Estimat	es not
20	467.1	282.0	-39.6	209.7	-55.1	provided by	y method
10	661.9	514.9	-22.2	403.0	-39.1	577.9	-12.7
5	868.4	776.0	-10.6	661.6	-23.8	840.3	-3.2
2	1164.5	1147.6	-1.5	1110.6	-4.6	1192.6	2.4
1	1411.4	1442.2	2.2	1535.7	8.8	1479.1	4.8

The comparisons of Table 5.1 indicated that the less frequent events were aligned closer to the independent methods, improving the confidence in these simulations. In particular, the 2% and

1% AEP simulated discharges were within 5% of all three independent method estimates except for the 1% RFFE estimate, which differed by 8.8%. Confidence in the results of these events was identified as the most crucial aspect of the simulation, because the 2% and 1% AEP design discharges are the recommended flood immunity threshold flows for the Warrego Highway crossing of Laidley Creek, as introduced in the literature review.

For the more frequent events, the difference between the simulations and estimates increased significantly, especially for the 50%, 20%, and to a lesser extent, the 10% AEP events. Generally, the simulated discharges had the closest alignment to the P&W estimates, followed by the FFA gamma distribution estimates, which also yielded the closest correlation to the 2% and 1% AEP events. The RFFE estimates had the largest difference for all design events.

5.1.2 Comparison of design discharges for future climate scenarios

Of the three independent techniques, only the P&W method had the capability to estimate future discharges from the revised IFD sets. The primary limitation of this approach is that the P&W method cannot estimate the 50% and 20% AEP event discharges, which had the highest percentage differences identified in Table 5.1. These discharges cannot be estimated because the P&W equations were developed for the 1 in 2 and 1 in 5 ARIs, which do not correspond to the aforementioned AEPs per Figure 2.11. Therefore the 50% and 20% AEP discharges for future climate scenarios cannot be validated and should be treated with extreme caution.

The remaining simulated discharge events were compared with the estimates obtained from the P&W equations for the two RCP scenarios in 2030, 2050 and 2090 (per Tables 3.7 and 3.8) by calculating the corresponding percentage difference, as listed in Tables 5.2 and 5.3.

Table 5.2: Comparison of RCP 4.5 scenario simulations with P&W estimates (m³/s)

AEP	Med	ian Discl	narge	P&W m	ethod revi	ised for R	CP 4.5 sce	enario (T	able 3.7)
(%)	(Table 4.2	2)	2030	% Diff.	2050	% Diff.	2090	% Diff.
50	281.1	296.2	330.0		D.:		. 1 11	.1 1	
20	547.5	580.1	630.4		Estimat	es not pro	vided by n	nethod	
10	767.4	799.5	865.8	646.5	-15.8	674.1	-15.7	716.0	-17.3
5	992.7	1044.5	1116.7	931.0	-6.2	967.3	-7.4	1022.2	-8.5
2	1315.1	1383.0	1483.7	1308.0	-0.5	1353.9	-2.1	1423.0	-4.1
1	1587.9	1659.5	1772.3	1611.6	1.5	1664.1	0.3	1742.8	-1.7

Table 5.3: Comparison of RCP 8.5 scenario simulations with P&W estimates (m³/s)

AEP	Med	ian Discl	harge	P&W m	ethod revi	ised for R	CP 8.5 sce	enario (T	able 3.8)
(%)	(Table 4.3	3)	2030	% Diff.	2050	% Diff.	2090	% Diff.
50	280.3	320.0	431.5		Б.:		'1 11	.1 1	
20	555.5	614.9	780.9		Estimat	tes not pro	vided by n	nethod	
10	779.5	842.5	1056.5	653.3	-16.2	701.9	-16.7	836.4	-20.8
5	1010.4	1086.9	1341.2	940.1	-7.0	1003.7	-7.7	1178.4	-12.1
2	1340.7	1446.6	1762.8	1319.5	-1.6	1399.8	-3.2	1617.6	-8.2
1	1612.4	1736.0	2096.3	1624.7	0.8	1716.3	-1.1	1963.1	-6.4

A similar level of alignment was observed between the RCP scenarios of Tables 5.2 and 5.3 indicating that the P&W equations were not necessarily less applicable for a particular scenario. In a similar manner to Table 5.1, the less frequent event estimations were aligned much closer to the corresponding simulated discharges for both emissions scenarios compared to the 10% and 5% AEP events. For the RCP 8.5 scenario, the estimate correlation decreased for the more distant projections however no definitive trend was observed in the RCP 4.5 projections.

The close correlation between the 2% and 1% AEP future event discharges confirmed the confidence already established in these simulations, enhancing the model suitability for future roadway immunity modelling. The magnitude of the differences between the P&W estimates and the simulated discharges is similar to the differences in the sample event distributions. However, it is reiterated that these estimates are simply a supporting validation of the model, which itself forms the most detailed representation of the Laidley Creek catchment hydrology.

5.2 Limitations of the results

The output validation indicated that the RORB simulation model was suitably calibrated for the less frequent 2% and 1% AEP events across current and future climate projection scenarios. This attribute was crucial in ensuring the results could support highway immunity modelling, as the primary intention of this project. However, the model performance worsened for the higher frequency events. The limitations of the results ultimately originated from the limitations associated with the input data as well as the modelling techniques used.

5.2.1 Data limitations

Several limitations were noted with the various forms of input data used to formulate the model. The limiting factors with the highest impact on the model outputs are discussed below.

5.2.1.1 Uncertainty in the climate projections

The uncertainty of the climate projections is best characterised by the extremely wide range of likely rates of change in increases to extreme rainfall per degree of surface temperature warming, as presented in Table 2.3. The median rates of change value were adopted for this project, rather than a distribution of potential rates of change.

The limitations in the climate projections originate from the uncertainties embedded within the GCMs to accurately forecast anticipated future climate conditions. This uncertainty is represented by the wide range of results obtained between the individual GCMs for different climate metrics, including rises in surface temperature, wind, humidity and evaporation.

In addition, the climate projections are simply that – forecasts of potential future conditions – and can only be verified after the occurrence through observation. These projections are continually revised with updated observations and advancements in literature.

5.2.1.2 Inadequate catchment coverage from rainfall stations

The spatial coverage of the catchment by the rainfall stations varied between calibrated events, but especially at an hourly level, was limited. The two streamflow gauging stations only began recording rainfall measurements in July 2024 while the majority of the BOM operated sites were located to the north of the catchment around the populated areas of Gatton and Plainland.

The rainfall interpolation was significantly influenced by the singular Townson daily rainfall station in the central and southern subcatchments It was difficult to distinguish between naturally occurring, long-term increased average rainfall at the site or an unreliable rainfall gauge. Other influential sites for these subcatchments were situated outside the catchment boundary, often amongst separate mountain ranges with highly variable and potentially unrepresentative rainfall patterns. The central subcatchments were poorly covered during the most recent events, as illustrated in Figure 5.1 following the closure of the Thornton gauge in 2008 and the Mount Berryman gauge in 2020. The coverage of older events, such as May 1996, was far superior.

These limitations were exacerbated when the hourly rainfall data was used to define the temporal patterns of the calibration events. Only three stations were identified and all three stations were located outside the catchment boundary, resulting in a questionable level of coverage. The UQ Gatton and Upper Tenthill stations are positioned close together with minimal unique coverage while the Bremer River site is located over 10 kilometres to the east of the catchment. The available hourly rainfall data was potentially unrepresentative of the

rainfall distributions within the central subcatchments, increasing the uncertainty of the results including magnitude and timing aspects.

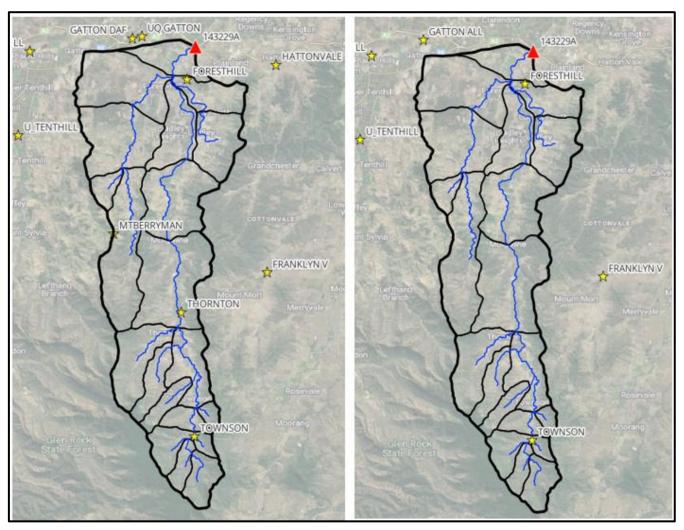


Figure 5.1: Reduction of central catchment daily rainfall station coverage from May 1996 (left) to February 2022 (right)

5.2.1.3 Duration of the gauged streamflow record

Despite the well-defined rating curve of the 143229A gauging station, the 33 gauged streamflow gauged record was considered relatively short ultimately limiting two critical aspects of the model development. The restricted duration significantly affected the reliability of the FFA as a broad range of dry and wet conditions prior to 1991 were not captured in the analysis. Within the most recent 13 years of the entire 33 year period, all three maximum flow events exceeding 1000 m³/s at 143229A occurred. These recorded events had a significant impact on increasing the discharge estimates for all AEPs, which may have resulted in an uncertain or inaccurate representation of the extended site history. This uncertainty is demonstrated by both the large variance between discharge estimates from the three fitted

distribution functions, as well as the large margin of error about the Bayesian posterior mode distribution. Secondly, notable high flow events which occurred prior to the gauge opening 1991 were not captured, hence reducing the number of available events to consider for calibration of the model parameters.

5.2.1.4 Simplification of catchment spatial features

The incorporation of catchment spatial features and attributes into a semi-distributed node-reach model required several simplifications. As previously discussed, the manually delineated catchment had an area of 452 km² while the WMIP 143229A station details page noted the catchment area as 462 km². This difference has the ability to impact the results to some extent.

The stream reach configuration and the subcatchment centroidal flow path arrangement has an impact on the results. The length of each reach was approximated by the vectorized stream path segments. These lengths were altered slightly where the subcatchment centroids did not closely align with the delineated streams. The reach slopes were averaged by considering the change in elevation between the centroid and junction nodes. A sufficient level of accuracy was difficult to obtain from the supplied DEM within the very flat lower reaches, potentially explaining the necessity for event specific translations during the calibration event modelling.

The entire catchment area was assumed as both rural and pervious to simplify modelling. In reality, a portion of the ground surface area consists of impervious surfaces, especially around the communities of Laidley, Forest Hill, and the eastern edges of Gatton. Computing the proportion of pervious land coverings was deemed outside the scope of work for this project. Similarly, the characteristics of urban runoff and storage in these locations were not considered, but would also impact the catchment response during an extreme rainfall event.

5.2.2 Modelling limitations

Limitations associated with the adopted modelling methodology facilitated the uncertainty in the results, especially for the higher frequency events. The basis of this report assumed that the relatively certain relationships, calibrations and simulations of the model remain an accuracte representation of future hydrologic conditions following revisions to the input rainfall datasets. The most notable considerations are discussed below.

5.2.2.1 Calibration uncertainty

Limitations attributed to the calibration of the RORB model parameters originated from the flood immunity modelling objectives of this project. The 2% and 1% AEP recommended

design flow thresholds meant an emphasis was placed on ensuring maximum flow characteristics were as precisely calibrated as possible, compared to the lower flow components of the event hydrographs. This emphasis is illustrated in Figure 5.2 for the January 2011 event, where an accurate representation of timing, shape and magnitude was obtained for the final flood peak, however the preceding fluctuations were less precisely calibrated.

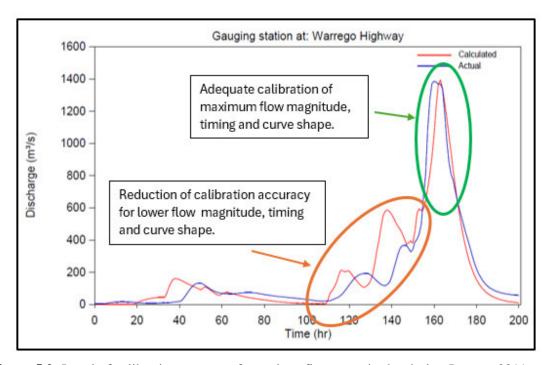


Figure 5.2: Level of calibration accuracy for various flow magnitudes during January 2011 event

Several technical components of the calibration also introduced uncertainties to the model. Foremost, many combinations of different k_c and IL values produced calculated hydrographs of similar appearance, meaning it was difficult to calibrate the most suitable set of parameters for each event. This limitation was overcome by considering the median value of kc across eight unique peak flow events, however this approach also introduced uncertainties to the results. The median kc = 27 simply represented the most appropriate representation for most events, however, this value was not identified as the best fit value for any individual event. The calibrated events were not re-evaluated using the median kc = 27 to adjust the fitted IL depth due to time constraints.

The number of events considered for calibration indicated multiple aspects associated with the uncertainty of the results. One identified outlier event from eight modelled events demonstrated the calibrated catchment parameters may not necessarily be applicable to all past events. Considering only 8 events during the calibration limited the variety of rainfall-runoff characteristics evaluated. Additional event modelling would confirm or disassociate the

calibrated parameter values. The requirement to introduce event specific translations during calibration also indicated a limited ability for the model to represent past peak flow events in the very flat, lower subcatchment reaches.

5.2.2.2 Simulation uncertainty

Limitations of the design simulations originated from the calibration limitations in conjunction with components of the simulation method. The nature of the correlation between the simulated and independent estimates for varying event frequencies was unsurprising, given the calibration was based upon eight extreme rainfall events with an emphasis on the accuracy of the alignment against maximum flow characteristics. However, the differences observed for the 50% and 20% AEP events for the current enveloped IFDs was significant enough to indicate the model had limited capability to accurately represent frequent flow events.

Limiting the simulation of each climate scenario to a set of ten sample simulations limited the accuracy of the results. The margin of error about each set of simulations represented up to 10.4% of the median event discharge (for current conditions) but this varied drastically between AEPs with no apparent trend. Additional samples would facilitate the convergence of results allowing for a more certain median discharge to be extracted.

The critical event duration was a particularly uncertain aspect of the simulation. Despite seven of the eight historic events used to calibrate the model exceeding 48 hours in duration, the critical duration throughout the simulation of current and future climate scenarios was predominantly the 24 hour design event, followed by the 144 hour event. Therefore, the accuracy of the design simulations, especially for current conditions, was limited. Two potential causes of this uncertainty were identified:

The RORB simulation specification considered a uniform spatial distribution of design rainfall across the 20 subcatchment centroidal inputs. However, the rainfall interpolations undertake for each of the eight calibrated events demonstrated that the distribution of rainfall is spatially variable across the catchment. The extent of the variability of rainfall across the calibrated events was observed to a variable factor itself. Therefore, the assumption that the design storm specification was spatially uniform is inaccurate. However, due to the insufficient coverage of long-term rainfall observation throughout several regions of the Laidley Creek catchment, it was impossible to ascertain an average rainfall spatial distribution pattern.

Secondly, the highest IFD adjustment factors for future climate projections corresponded to the shortest duration events, increasing the likelihood these events would dominate in future years.

5.2.2.3 Validation uncertainty

The independent validation methods and the generic comparative approach to evaluate the simulated results have some degree of uncertainty. As discussed, the streamgauge duration considered for FFA was inadequate while the regionalised parameters of the RFFE and P&W methods represent major simplifications of the catchment hydrologic characteristics.

Minimal verification of the discharges corresponding to future climate scenarios was attained. While FFA methods typically provide the most localised representation of the three estimation techniques, the method is fundamentally an analysis of historically gauged data which cannot be applied to forecast evolving hydrologic characteristics.

Only the singular P&W method was able to produce some revised estimates, however, the high frequency 50% and 20% AEP events were unable to be validated by any method. The necessary assumption that the numeric P&W relationships were maintained for future climate scenarios, despite being developed from historically gauged data, further limited the certainty of the verification.

5.2.2.4 Limited consideration of time variable components

As introduced in the literature review, regional catchment hydrology is influenced by several factors which are likely to experience changes over time but were not incorporated within the revised future scenario models. Factors not considered are both dependent and non-dependent on the level of future climate change, as noted below:

- Altered storm event characteristics including temporal distribution patterns and losses as a consequence of evolving climate conditions.
- Urbanisation and development within the Laidley Creek catchment with the potential to modify natural flow regimes and reduce the proportion of pervious surface cover.

5.3 Potential model improvements

Potential improvements were identified to reduce the limitations and uncertainties associated with the results. These improvements were unable to be implemented within the defined project scope due to overall time and resource constraints. The suggested improvements include the completion of:

- Additional model calibrations of past events to verify the suitability of the median catchment storage parameter k_c , with a focus on improving the calibration of medium and high frequency flows. The viability of continuous catchment simulation methods to improve the accuracy of frequent event discharges should be explored.
- Additional design simulations with modified loss values more representative of the wide spread of losses observed during calibration.
- Additional design simulations with an increased number of revised IFD sets incorporating a wider range of the projected surface temperature rises.
- Additional sample simulations for each climate scenario to converge about the median discharge estimates.
- An analysis of the long-term averaged spatial distribution of rainfall (potentially through gridded rainfall products) to improve the accuracy of the design event timing and critical duration.
- The acquisition of remote-sensed Light Detection and Ranging (LiDAR), ground survey, or similarly precise elevation data to improve the accuracy of the catchment spatial model, especially within the very flat lower subcatchments which resulted in the necessity for event specific translations.
- The consideration of additional time variable factors within the future climate dependent discharge models, as discussed.

6. Conclusion

The content presented in this dissertation described the extent and adequacy in which the research project addressed each of the six project objectives. Each chapter represented the progression of the project towards the definition of a set of current and future design discharges for the Laidley Creek catchment at the Warrego Highway crossing site. The set of results offers a practical indication of anticipated future flow characteristics in the short, medium and long-term horizons to inform corridor planning and infrastructure design within DTMR.

The literature review demonstrated an evident lack of knowledge surrounding the exact impacts of climate change on peak discharge hydrology within the Laidley Creek catchment. Several background aspects of the research problem were introduced in this chapter, including the physical and hydrological attributes of the case study catchment. The cross-disciplinary flood immunity design requirements for the state-controlled Warrego Highway corridor were introduced, which indicated that producing accurate design discharges for the 2% and 1% AEP events was crucial towards achieving the objectives of this project. The applications, advantages and limitations of various hydrologic modelling techniques were examined, ultimately leading to the development of an event based, semi-distributed node-reach runoff routing model within the RORB platform. The broad implications of climate change on catchment hydrology were introduced and recently released design approaches of ARR 2019 were examined.

The literature review findings were synthesised into an appropriately scoped design methodology which considered best practice modelling techniques. Spatial and hydrologic data was retrieved from relevant sources and processed to formulate the catchment model and storm event files. The RORB model parameters were calibrated against eight historic high rainfall-runoff events. The MC simulation method was chosen to estimate the design discharges for the current and six future climate scenarios from sets of sample simulations, which are analysed in the results chapter of this report. A sensitivity analysis of the model parameters was undertaken and independent estimation techniques were used to evaluate the likely accuracy and extent of uncertainty in the results. The advantages, assumptions, uncertainties and limitations of the model were discussed throughout the development and analysis of the model.

6.1 Summary of results

The results provided a useful indication of the anticipated peak flow characteristics of the Laidley Creek catchment under a range of differing climate forecasts. Most crucially, the independent estimation techniques implied that the model was accurately calibrated for the National Highway flood immunity threshold flows corresponding to the 2% and 1% AEP events. The uncertainty in the results increased for the higher frequency events, with the alignment between the independent estimates and the simulated flows decreasing.

The results projected that for the 2% AEP design event, peak discharge from the RCP 4.5 emissions scenario will increase from its current magnitude by 13% by 2030, 19% by 2050 and 27% by 2090. For the RCP 8.5 scenario, these projections rise to 15% by 2030, 24% by 2050 and drastically to 51% by 2090.

For the 1% AEP design event, the corresponding increases to peak discharge are slightly less than the 2% AEP projections. Under the RCP 4.5 emissions scenario, the projected rise in 1% AEP discharges are 12% by 2030, 18% by 2050 and 26% by 2090. For the RCP 8.5 scenario, these projections are 14% by 2030, 23% by 2050 and 49% by 2090.

6.2 Recommended future work

Two avenues of additional work are recommended as an extension of this dissertation. These recommendations were formed by considering the broader context of the design problem, specifically flood resilience and immunity modelling of the Warrego Highway at Laidley Creek for DTMR as administrators of the state-controlled network.

Firstly, it is recommended that enhancements are made to the hydrologic model (developed in this project report) by implementing some or all of the suggested list of improvements. These enhancements would reduce the uncertainty of the simulated design discharges, particularly for the higher frequency discharges.

Secondly, undertaking a two dimension hydraulic investigation of the Warrego Highway crossing of Laidley Creek at the Jack Martin Bridge is the next logical avenue of research. The output discharges obtained from this hydrologic simulation were formatted to be directly applicable as inputs for a hydraulic flow model. A potential model could therefore consider the design discharges corresponding to both current and future climate conditions, to evaluate the impact of climate change on the flood immunity of state-controlled roadway infrastructure.

The acquisition of LiDAR and/or bridge section survey data was not viable for this project but would significantly assist the development of a two dimensional hydraulic study.

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Appendix A: Project specification and work plan (topic later refined)

Title: Instantaneous Peak Flow Estimation Considering Climate Change Scenarios in the

Upper Condamine Basin

Name: Hayden Jago Student ID:

Major: Civil Engineering
Supervisor: Dr Sreeni Chadalavada
Enrolment: ENP4111-YL1-2024

Project Specification

Introduction and Background:

Changing climates and weather patterns are having diverse impacts on regions globally. Surface temperatures have risen by approximately 1.5°C since Australian records began in 1910 (Bureau of Meteorology & CSIRO 2022). Surface temperature increases are a consequence of greenhouse gas emissions from energy, industry, transport, and agriculture sectors (United Nations Sustainable Development Goals Summit 2023). As a result, total annual rainfall in regional Queensland has declined since 1970 (Bureau of Meteorology & CSIRO 2022). Climate forecasts have predicted even greater rises in surface temperatures, particularly in Southern Queensland. Per degree of future warming, heavy rainfall intensity is predicted to increase by 2-15% (Guerreiro et al. 2018; Bates et al. 2019; Bureau of Meteorology & CSIRO 2022). Further surface warming could increase the intensity and frequency of short, heavy precipitation "bursts" causing flood events in regional catchments (Visser et al. 2023).

The Upper Condamine River Basin is located in regional southeastern Queensland and encompasses the headwaters and tributaries of the Condamine River. The basin features varied topography including the elevated western faces of the Great Dividing Range and the flat open plains of the Darling Downs and Western Downs, see Figure 1. The catchment is the main water source for many irrigation and domestic uses (Dafny & Silburn 2013). The diverse elevation of locations within the catchment contribute to dynamic rainfall distribution patterns (Dafny & Silburn 2013). Consequently, the generated hydrologic runoff systems are highly complex. This research will focus on quantifying the parameters that affect the hydrologic cycle in the Upper Condamine Basin to generate a rainfall-discharge hydrograph representative of the conditions expected in future decades.

Modelling allows engineers and policymakers to comprehend hydrologic processes and enable measures that best manage water resources distribution during periods of drought or flooding. Queensland hydrologic modelling is directed by design guidelines published in Books 1 to 5 of Engineers Australia's *Australian Rainfall and Runoff: A Guide to Flood Estimation* (referred to as ARR 2019 (Ball et al. 2019; Ball, Weinmann & Boyd 2019; Bates et al. 2019; Jordan, Seed & Nathan 2019; Nathan et al. 2019). Book 1 Chapter 5 provides generalised technical advice for considering future climate change scenarios in a project design life. ARR 2019 uses flood hydrograph generation as a method to estimate instantaneous peak flow. This technique considers catchment runoff generation as the conversion of rainfall through losses into a downstream flood hydrograph (Bates et al. 2019).

The major sources of raw data for this project are easily accessible online. Design rainfall IFDs for locations across the Upper Condamine Basin are obtainable from the Bureau of Meteorology (2016). Additionally, streamflow discharge data is available at several Queensland Government gauge locations along the Condamine River (Department of Regional Development, Manufacturing & Water 2023).

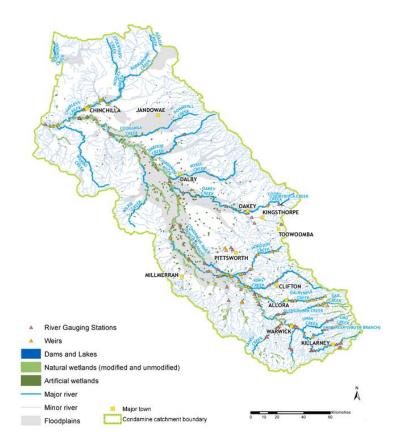


Figure 1: The Upper Condamine River Basin (Department of Agriculture, Water & the Environment (Geological & Bioregional Assessments) 2019)

Objectives and Aims:

This research project aims to fulfil gaps in knowledge about the impacts of climate change scenarios on hydrologic processes at locations of various elevation in the Upper Condamine River Basin.

Specific Objectives:

- Calibrate and modify existing design rainfall models to accurately represent dynamic future rainfall events impacted by climate change.
- Develop future rainfall-runoff hydrographs for a number of high- and low-elevation case study locations within the Upper Condamine Basin.
- Compare the findings of future flow hydrograph modelling with current flow characteristics.

Expected Outcomes:

- Clear identification of changes observed to rainfall models and discharge hydrographs in future decades to understand the quantifiable impacts of climate change scenarios on peak discharge across sites in the catchment.
- Enhanced understanding of high elevation hydrology in Queensland.
- Possible starting point for further research into integrated watershed management, catchment flood simulations, or flood damage probability assessment.

The knowledge from this research would have the potential to facilitate the development of flood management strategies and engineering solutions to mitigate the impacts of future weather events on local communities in the Upper Condamine Basin. Additionally, this research may assist sustainable water resource allocation and distribution in the region.

Work Plan

Programme: Version 2 - February 2024

Month 1: Project Commencement and Proposal

- Liaise with supervisor to refine scope of research.
- Develop project specification and formulate detailed programme schedule for proposal submission.

Months 2-3: Literature Review

- Conduct a comprehensive literature review of future climate scenarios and their impact on fundamental hydrologic design parameters.
- Review key model parameters including rainfall/flow data sources relevant to the case study region. Identify parameters that require further investigation.
- Review existing hydrologic technical guidelines and publications to begin developing the research methodology.

Month 4: Research Methodology Development

- Collate findings from literature review to identify appropriate modelling processes and critical locations suitable to become case study sites for hydrologic analysis (including sites of high and low elevations)
- Finalise the project methodology through a clearly defined hydrologic modelling procedure, step-by-step.

Months 5-7: Hydrologic Modelling

- Conduct baseline modelling for flow at existing streamgauge sites.
- Verify current rainfall models correspond to current streamflow models.
- Extend current rainfall models to develop rainfall-runoff hydrographs at each study location with no available streamgauge data.
- Formulate future rainfall models by considering the calibrated rainfall models in conjunction with the identified impacts of climate change scenarios on rainfall patterns.
- Apply future rainfall models to the study locations to develop future rainfall-runoff hydrographs.

Months 8-9: Hydrologic Analysis, Model Refinements and Report Writing

- Compare the current and future rainfall-runoff hydrographs to identify impacts of climate change factors on peak flow estimation in the case study catchment.
- Make any model refinements to improve accuracy and usefulness of findings.
- Compose findings into a technical discussion in the dissertation report.

Month 10: Project Finalisation

- Complete, review, proof, and submit the report.
- Prepare a visual presentation summarising the project development, model findings and outcomes. Submit a personal reflection about the project.

See attached proposed schedule for detailed breakdown of individual project components below.

Timeline:

YEAR LONG STUDY PERIOD [ENP4111] WEEK	1	2	3	4 5	9	7	8	6	10	11	12	13	14 1	15 1	16 1	17 18	18 19	9 20	21	22	23	24	25	26	27	28	29	30	31	32	33	34	35 3	36 3	37 38	38 36	39 40	0 41	1 42	43	44	45	46
	₽Z-nel-22	42-nsl-92	1S-Feb-24 05-Feb-24	13-Feb-24	26-Feb-24	04-Mar-24	11-Mar-24	18-Mar-24	25-Mar-24	PS-1qA-10	42-1qA-80	15-Apr-24	42-7pk-22	42-7qA-62 42-ysM-30	#2-Y5IVI-00	PZ-YBIVI-CT	₽Z-γsM-7Z	42-nul-€0	\$2-nul-0£	₽S-nul-\t	42-nul-42	PS-Int-to	42-lut-80	pz-Int-st	22-Jul-24	₽2-Iul-62	₽ Δ-8υΑ-20	₽2-guA-Σ£	42-3uA-€£	42-8uA-82	%Z-dəS-Z0	₽Z-dəS-60	\$2-də2-97	30-Sep-24	30-Sep-24	74-DC-54	77-DG-57	28-04-24	42-voN-40	42-voN-11	₽2-voN-81	PS-voN-SS	02-Dec-24
TASKS:						TRIMESTER 1	STER	1						BREAK	×						TRIME	IESTER	2						BREAK	AK						TRIN	rrimester	R 3					
Task 1: Project preparation	1	2	3 '	4 5	9	7	8	6	10	11	12	13	14	1 .	2	1 2	3	4	2	9	7	8	6	10	11	12	13	14	1	2	1	2	3	4	5 6	6 7	7 8	6 8	10	11	12	13	14
1A Proposal & justification approval																																											
1B Report paper drafting																																					H						
Task 2: Literature review and methodology	1	2	3	4 5	9	7	8	6	10	11	12	13	14	1	2	-7	3	4	2	9	7	8	6	10	11	12	13	14	1	2		2	3	4	2 6	6 7	7 8	6	11	11	12	13	14
2A Background literature review																																											
2B Methodology - hydrologic method formulation	ō																																										
2C Methodology - site selection																																											
Task 3: Hydrologic model development	1	2	3	4 5	9	7	8	6	10	11	12	13	14	1 .	2	1 2	2 3	4	2	9	7	8	6	10	11	12	13	14	1	2	1	2	3	4 5	5 6	6 7	7 8	6	10	11	12	13	14
3A Existing streamgauge site modelling																																											
3B Current rainfall model calibration=													_		-																												
3C Future (Climate impacted) rainfall modelling																					호	HOLIDAY TRIP	TRIP																				
3D Future peak flow modelling																																					-						
Task 4: Hydrologic model verification	1	7	3	4 5	9	7	8	6	91	11	12	13	14	1	2	-	3	4	S	9	7	∞	6	10	#	12	13	14	1	2		2	3	4	5 6	6 7	7 8	6	10	11	12	13	14
4A Results analysis																																											
4B Model refinements & recommendations			\dashv	\dashv				j	Ī	T	d	\exists			\dashv	\dashv	-		_																	\dashv	_	_	_	_	_	_	_
Task 5: Project Finalisation	1	2	3	4 5	9	7	∞	6	10	#	12	13	14	1	2	7	3	4	2	9	7	∞	6	10	11	12	13	14	1	2		2	8	4	5 6	6 7	7 8	6	2	11	12	13	14
5A Research presentation and reflection																																	\exists		TRP	d.		_					
58 Report paper completion			\dashv	$-\ $	4								\dashv	\dashv	\dashv	\dashv	\dashv	-	4	_	_									7										_	_	4	_

Resources required:

o **Equipment:**

- Personal computer with internet access.
- Dual monitor setup for convenient, simultaneous access to different aspects of project (i.e. modelling calculations and report paper).

Software:

- Prepare dissertation paper and presentation slides using Word and PowerPoint, which are included in the university edition of the Microsoft Office suite.
- Perform hydrologic model calculations and obtain graphical outputs in Excel.
- Manage references using university-provided EndNote citation tool.
- Manage dissertation storage and backup using OneDrive cloud.

Access:

- Access to university library resources and online scholarly search engines for literature review and methodology development.
- Retrieve streamgauge observations and rainfall metrics from official government databases.

Appendix B: Rainfall station datasets coverage

Table B1: Daily rainfall stations with dated averages to Forest Hill reference baseline

Station Name	Station	Record	Record	Average daily	F.H. time equivalent	Difference to baseline		ion event (Table 3.2)
Station Ivalle	ID	start date	finish date	rainfall (mm/d)	baseline (mm/d)	(%)	Full record	Partial record
Forest Hill	40079	1/02/1894	Open	2.132	Catchme	nt baseline	V	
Mulgowie	40570	1/01/1998	8/12/2004	2.053	1.784	15.07		
Laidley	40716	27/01/2009	30/05/2011	9.782	2.556	282.75		
Laidley PO	40114	1/02/1894	31/12/1993	2.202	2.132	3.27		
Laidley Ck	40011	1/10/1964	31/12/1979	2.564	2.539	0.98		
Mulgowie TM	40835	2/06/2010	6/01/2012	0.711	2.943	-75.83		
Mt Berryman	40310	1/07/1961	26/01/2020	2.413	2.173	11.03		Ø
Grandchester	40091	12/02/1894	31/03/2014	2.406	2.122	13.37		$\overline{\mathbf{Q}}$
Upper Tenthill	40388	1/01/1959	Open	2.108	2.228	-5.39	Ø	
Mt Sylvia	40384	1/07/1953	31/10/2002	2.085	2.245	-7.13		
Gatton DAFF	40436	1/07/1968	31/05/2014	2.194	2.181	0.62		Ø
UQ Gatton	40082	1/08/1897	30/04/2024	2.094	2.133	-1.85	\square	
Thornton	40751	1/11/1993	31/03/2008	2.189	1.787	22.49		
Franklyn Vale	40374	1/02/1894	Open	2.407	2.132	12.88	Ø	
Rock View	40605	1/03/1919	31/12/1969	2.748	2.102	30.73		
Franklyn Vale Alert	40912	30/11/2000	Open	2.220	2.128	4.32		Ø
Gatton Allan St	40083	1/02/1894	Open	2.130	2.132	-0.07	Ø	
Hattonvale	40095	1/08/1909	11/03/2020	2.115	2.120	-0.22		Ø
Thornton BVRT	40202	1/08/1915	30/06/1983	2.647	2.166	22.23		
Placid Hills	40449	1/01/1970	Open	2.183	2.206	-1.07	V	
Townson East	40392	1/01/1958	31/08/1978	3.250	2.439	33.23		
Townson	40675	1/12/1977	9/09/2023	3.023	2.117	42.85		V

High priority daily rainfall stations

- Forest Hill (baseline for evaluation of poor sites)
- Upper Tenthill
- UQ Gatton
- Franklyn Vale (eastern coverage)
- Gatton Allan St
- Placid Hills
- Townson (southern coverage until 2023)
- Mt Berryman (south-western coverage until 2020)
- Grandchester (north-eastern coverage until 2014)
- Hattonvale (north-eastern coverage until 2014)
- Gatton DAFF (northern sites covered during offline periods until 2014)

Table B2: Hourly rainfall stations

Station name	Station	Station	Record	Record		tion event e (Table 3.2)
	ID	operator	start date	finish date	Full record	Partial record
UQ Gatton	40082	BOM	16/07/2002	Open		Ø
Tenthill Creek	143212A	DRDMW	10/02/1993	Open	Ø	
Bremer River at Adams Bridge	143110A	DRDMW	13/11/1992	Open	Ø	

Appendix C: Design rainfall IFDs at Laidley Creek catchment centroid

	Table C1: LI	MB 2020 – BOM	IFD envelope	(current climate)
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Coordinates	-27.	727	152	.355		ENVE	LOPED)
AEP (%) / Duration	63.23	50	20	10	5	2	1	0.5
5 min	9.5	10.5	13.7	16.1	18.5	21.9	24.7	28.3
10 min	15.5	17.3	22.5	25.9	29.4	34.8	38.9	44.6
15 min	19.5	21.9	28.5	32.5	37.3	44.1	49.4	56.5
20 min	22.4	25.2	32.9	37.5	43.1	51.2	57.3	65.5
25 min	24.6	27.7	36.4	41.3	47.8	56.7	63.6	72.8
30 min	26.4	29.8	39.2	44.5	51.6	61.3	68.8	78.7
45 min	30.2	34.2	45.2	51.5	59.9	71.4	80.2	91.8
1 hr	32.9	37.3	49.5	56.5	65.6	78.3	88.2	101
1.5 hr	36.6	41.5	55.5	63.7	73.6	88	99.3	114
2 hr	39.3	44.6	59.9	69.2	79.3	94.9	107	123
3 hr	43.3	49.2	66.6	77.6	87.8	105	119	136
4.5 hr	47.8	54.3	74.4	87.4	99.8	117	132	151
6 hr	51.3	58.5	80.6	95.4	109.7	128.3	143	163
9 hr	56.9	65.1	90.9	108.7	126.1	149.7	167.9	191.9
12 hr	61.3	70.3	99.3	119.6	139.8	167.6	189.4	217.2
18 hr	68.9	78.7	112.9	137.5	162.5	197.4	225.5	258.4
24 hr	75.9	86	123.8	151.9	181	222.1	255.6	292.1
30 hr	81.7	92.8	132.9	164.1	196.7	243.3	281.6	321.6
36 hr	86.5	98.5	140.7	174.7	210.5	261.9	304.6	349.1
48 hr	94.3	108	153.6	192.3	233.5	293.4	343.7	394.7
72 hr	105	120	173	218.5	268	340.8	403.2	464.2
96 hr	111	128	186.9	237.6	292.9	375.1	446.1	513.3
120 hr	116	133	198.3	252.5	311.9	400.8	477.9	548.4
144 hr	119	136	208	264.7	326.9	420.4	501.6	576.1
168 hr	121	141.8	216.5	275.1	338.9	435.6	519.3	599.5

<u>Table C2: IFD adjustment factors – RCP 4.5 scenarios</u>

Storm Duration / Year	<1 hr	1.5 hr	2 hr	3 hr	4.5 hr	6 hr	9 hr	12 hr	18 hr	>24 hr
2030	1.18	1.17	1.16	1.14	1.13	1.12	1.12	1.11	1.1	1.1
2050	1.27	1.24	1.23	1.21	1.19	1.18	1.17	1.16	1.15	1.14
2090	1.4	1.36	1.34	1.31	1.28	1.26	1.24	1.23	1.21	1.2

<u>Table C3: IFD adjustment factors – RCP 8.5 scenarios</u>

Storm Duration / Year	<1 hr	1.5 hr	2 hr	3 hr	4.5 hr	6 hr	9 hr	12 hr	18 hr	>24 hr
2030	1.2	1.18	1.17	1.16	1.14	1.13	1.13	1.12	1.11	1.11
2050	1.34	1.31	1.29	1.26	1.24	1.23	1.21	1.2	1.18	1.18
2090	1.77	1.69	1.64	1.58	1.52	1.49	1.45	1.42	1.39	1.37

Table C4: RCP 4.5 2030 IFDs

Coordinates	-27.	727	152.355		RCP 4.5 2030				
AEP (%) / Duration	63.23	50	20	10	5	2	1	0.5	
5 min	11.2	12.4	16.2	19.0	21.8	25.8	29.1	33.4	
10 min	18.3	20.4	26.6	30.6	34.7	41.1	45.9	52.6	
15 min	23.0	25.8	33.6	38.4	44.0	52.0	58.3	66.7	
20 min	26.4	29.7	38.8	44.3	50.9	60.4	67.6	77.3	
25 min	29.0	32.7	43.0	48.7	56.4	66.9	75.0	85.9	
30 min	31.2	35.2	46.3	52.5	60.9	72.3	81.2	92.9	
45 min	35.6	40.4	53.3	60.8	70.7	84.3	94.6	108.3	
1 hr	38.8	44.0	58.4	66.7	77.4	92.4	104.1	119.2	
1.5 hr	42.8	48.6	64.9	74.5	86.1	103.0	116.2	133.4	
2 hr	45.6	51.7	69.5	80.3	92.0	110.1	124.1	142.7	
3 hr	49.4	56.1	75.9	88.5	100.1	119.7	135.7	155.0	
4.5 hr	54.0	61.4	84.1	98.8	112.8	132.2	149.2	170.6	
6 hr	57.5	65.5	90.3	106.8	122.9	143.7	160.2	182.6	
9 hr	63.7	72.9	101.8	121.7	141.2	167.7	188.0	214.9	
12 hr	68.0	78.0	110.2	132.8	155.2	186.0	210.2	241.1	
18 hr	75.8	86.6	124.2	151.3	178.8	217.1	248.1	284.2	
24 hr	83.5	94.6	136.2	167.1	199.1	244.3	281.2	321.3	
30 hr	89.9	102.1	146.2	180.5	216.4	267.6	309.8	353.8	
36 hr	95.2	108.4	154.8	192.2	231.6	288.1	335.1	384.0	
48 hr	103.7	118.8	169.0	211.5	256.9	322.7	378.1	434.2	
72 hr	115.5	132.0	190.3	240.4	294.8	374.9	443.5	510.6	
96 hr	122.1	140.8	205.6	261.4	322.2	412.6	490.7	564.6	
120 hr	127.6	146.3	218.1	277.8	343.1	440.9	525.7	603.2	
144 hr	130.9	149.6	228.8	291.2	359.6	462.4	551.8	633.7	
168 hr	133.1	156.0	238.2	302.6	372.8	479.2	571.2	659.5	

Table C5: RCP 4.5 2050 IFDs

Coordinates	-27.727		152.355		RCP 4.5 2050				
AEP (%) / Duration	63.23	50	20	10	5	2	1	0.5	
5 min	12.1	13.3	17.4	20.4	23.5	27.8	31.4	35.9	
10 min	19.7	22.0	28.6	32.9	37.3	44.2	49.4	56.6	
15 min	24.8	27.8	36.2	41.3	47.4	56.0	62.7	71.8	
20 min	28.4	32.0	41.8	47.6	54.7	65.0	72.8	83.2	
25 min	31.2	35.2	46.2	52.5	60.7	72.0	80.8	92.5	
30 min	33.5	37.8	49.8	56.5	65.5	77.9	87.4	99.9	
45 min	38.4	43.4	57.4	65.4	76.1	90.7	101.9	116.6	
1 hr	41.8	47.4	62.9	71.8	83.3	99.4	112.0	128.3	
1.5 hr	45.4	51.5	68.8	79.0	91.3	109.1	123.1	141.4	
2 hr	48.3	54.9	73.7	85.1	97.5	116.7	131.6	151.3	
3 hr	52.4	59.5	80.6	93.9	106.2	127.1	144.0	164.6	
4.5 hr	56.9	64.6	88.5	104.0	118.8	139.2	157.1	179.7	
6 hr	60.5	69.0	95.1	112.6	129.4	151.4	168.7	192.3	
9 hr	66.6	76.2	106.4	127.2	147.5	175.1	196.4	224.5	
12 hr	71.1	81.5	115.2	138.7	162.2	194.4	219.7	252.0	
18 hr	79.2	90.5	129.8	158.1	186.9	227.0	259.3	297.2	
24 hr	86.5	98.0	141.1	173.2	206.3	253.2	291.4	333.0	
30 hr	93.1	105.8	151.5	187.1	224.2	277.4	321.0	366.6	
36 hr	98.6	112.3	160.4	199.2	240.0	298.6	347.2	398.0	
48 hr	107.5	123.1	175.1	219.2	266.2	334.5	391.8	450.0	
72 hr	119.7	136.8	197.2	249.1	305.5	388.5	459.6	529.2	
96 hr	126.5	145.9	213.1	270.9	333.9	427.6	508.6	585.2	
120 hr	132.2	151.6	226.1	287.9	355.6	456.9	544.8	625.2	
144 hr	135.7	155.0	237.1	301.8	372.7	479.3	571.8	656.8	
168 hr	137.9	161.7	246.8	313.6	386.3	496.6	592.0	683.4	

Table C6: RCP 4.5 2090 IFDs

Coordinates	-27.	727	152.355		RCP 4.5 2090				
AEP (%) / Duration	63.23	50	20	10	5	2	1	0.5	
5 min	13.3	14.7	19.2	22.5	25.9	30.7	34.6	39.6	
10 min	21.7	24.2	31.5	36.3	41.2	48.7	54.5	62.4	
15 min	27.3	30.7	39.9	45.5	52.2	61.7	69.2	79.1	
20 min	31.4	35.3	46.1	52.5	60.3	71.7	80.2	91.7	
25 min	34.4	38.8	51.0	57.8	66.9	79.4	89.0	101.9	
30 min	37.0	41.7	54.9	62.3	72.2	85.8	96.3	110.2	
45 min	42.3	47.9	63.3	72.1	83.9	100.0	112.3	128.5	
1 hr	46.1	52.2	69.3	79.1	91.8	109.6	123.5	141.4	
1.5 hr	49.8	56.4	75.5	86.6	100.1	119.7	135.0	155.0	
2 hr	52.7	59.8	80.3	92.7	106.3	127.2	143.4	164.8	
3 hr	56.7	64.5	87.2	101.7	115.0	137.6	155.9	178.2	
4.5 hr	61.2	69.5	95.2	111.9	127.7	149.8	169.0	193.3	
6 hr	64.6	73.7	101.6	120.2	138.2	161.7	180.2	205.4	
9 hr	70.6	80.7	112.7	134.8	156.4	185.6	208.2	238.0	
12 hr	75.4	86.5	122.1	147.1	172.0	206.1	233.0	267.2	
18 hr	83.4	95.2	136.6	166.4	196.6	238.9	272.9	312.7	
24 hr	91.1	103.2	148.6	182.3	217.2	266.5	306.7	350.5	
30 hr	98.0	111.4	159.5	196.9	236.0	292.0	337.9	385.9	
36 hr	103.8	118.2	168.8	209.6	252.6	314.3	365.5	418.9	
48 hr	113.2	129.6	184.3	230.8	280.2	352.1	412.4	473.6	
72 hr	126.0	144.0	207.6	262.2	321.6	409.0	483.8	557.0	
96 hr	133.2	153.6	224.3	285.1	351.5	450.1	535.3	616.0	
120 hr	139.2	159.6	238.0	303.0	374.3	481.0	573.5	658.1	
144 hr	142.8	163.2	249.6	317.6	392.3	504.5	601.9	691.3	
168 hr	145.2	170.2	259.8	330.1	406.7	522.7	623.2	719.4	

Table C7: RCP 8.5 2030 IFDs

Coordinates	-27.	727	152	.355		RCP 8	3.5 2030	
AEP (%) / Duration	63.23	50	20	10	5	2	1	0.5
5 min	11.4	12.6	16.4	19.3	22.2	26.3	29.6	34.0
10 min	18.6	20.8	27.0	31.1	35.3	41.8	46.7	53.5
15 min	23.4	26.3	34.2	39.0	44.8	52.9	59.3	67.8
20 min	26.9	30.2	39.5	45.0	51.7	61.4	68.8	78.6
25 min	29.5	33.2	43.7	49.6	57.4	68.0	76.3	87.4
30 min	31.7	35.8	47.0	53.4	61.9	73.6	82.6	94.4
45 min	36.2	41.0	54.2	61.8	71.9	85.7	96.2	110.2
1 hr	39.5	44.8	59.4	67.8	78.7	94.0	105.8	121.2
1.5 hr	43.2	49.0	65.5	75.2	86.8	103.8	117.2	134.5
2 hr	46.0	52.2	70.1	81.0	92.8	111.0	125.2	143.9
3 hr	50.2	57.1	77.3	90.0	101.8	121.8	138.0	157.8
4.5 hr	54.5	61.9	84.8	99.6	113.8	133.4	150.5	172.1
6 hr	58.0	66.1	91.1	107.8	124.0	145.0	161.6	184.2
9 hr	64.3	73.6	102.7	122.8	142.5	169.2	189.7	216.8
12 hr	68.7	78.7	111.2	134.0	156.6	187.7	212.1	243.3
18 hr	76.5	87.4	125.3	152.6	180.4	219.1	250.3	286.8
24 hr	84.2	95.5	137.4	168.6	200.9	246.5	283.7	324.2
30 hr	90.7	103.0	147.5	182.2	218.3	270.1	312.6	357.0
36 hr	96.0	109.3	156.2	193.9	233.7	290.7	338.1	387.5
48 hr	104.7	119.9	170.5	213.5	259.2	325.7	381.5	438.1
72 hr	116.6	133.2	192.0	242.5	297.5	378.3	447.6	515.3
96 hr	123.2	142.1	207.5	263.7	325.1	416.4	495.2	569.8
120 hr	128.8	147.6	220.1	280.3	346.2	444.9	530.5	608.7
144 hr	132.1	151.0	230.9	293.8	362.9	466.6	556.8	639.5
168 hr	134.3	157.4	240.3	305.4	376.2	483.5	576.4	665.4

Table C8: RCP 8.5 2050 IFDs

Coordinates	-27.	727	152	.355		RCP 8	3.5 2070	
AEP (%) / Duration	63.23	50	20	10	5	2	1	0.5
5 min	12.7	14.1	18.4	21.6	24.8	29.3	33.1	37.9
10 min	20.8	23.2	30.2	34.7	39.4	46.6	52.1	59.8
15 min	26.1	29.3	38.2	43.6	50.0	59.1	66.2	75.7
20 min	30.0	33.8	44.1	50.3	57.8	68.6	76.8	87.8
25 min	33.0	37.1	48.8	55.3	64.1	76.0	85.2	97.6
30 min	35.4	39.9	52.5	59.6	69.1	82.1	92.2	105.5
45 min	40.5	45.8	60.6	69.0	80.3	95.7	107.5	123.0
1 hr	44.1	50.0	66.3	75.7	87.9	104.9	118.2	135.3
1.5 hr	47.9	54.4	72.7	83.4	96.4	115.3	130.1	149.3
2 hr	50.7	57.5	77.3	89.3	102.3	122.4	138.0	158.7
3 hr	54.6	62.0	83.9	97.8	110.6	132.3	149.9	171.4
4.5 hr	59.3	67.3	92.3	108.4	123.8	145.1	163.7	187.2
6 hr	63.1	72.0	99.1	117.3	134.9 157.8		175.9	200.5
9 hr	68.8	78.8	110.0	131.5	152.6	181.1	203.2	232.2
12 hr	73.6	84.4	119.2	143.5	167.8	201.1	227.3	260.6
18 hr	81.3	92.9	133.2	162.3	191.8	232.9	266.1	304.9
24 hr	89.6	101.5	146.1	179.2	213.6	262.1	301.6	344.7
30 hr	96.4	109.5	156.8	193.6	232.1	287.1	332.3	379.5
36 hr	102.1	116.2	166.0	206.1	248.4	309.0	359.4	411.9
48 hr	111.3	127.4	181.2	226.9	275.5	346.2	405.6	465.7
72 hr	123.9	141.6	204.1	257.8	316.2	402.1	475.8	547.8
96 hr	131.0	151.0	220.5	280.4	345.6	442.6	526.4	605.7
120 hr	136.9	156.9	234.0	298.0	368.0	472.9	563.9	647.1
144 hr	140.4	160.5	245.4	312.3	385.7	496.1	591.9	679.8
168 hr	142.8	167.3	255.5	324.6	399.9	514.0	612.8	707.4

Table C9: RCP 8.5 2090 IFDs

Coordinates	-27.	727	152	.355		RCP 8	3.5 2090	
AEP (%) / Duration	63.23	50	20	10	5	2	1	0.5
5 min	16.8	18.6	24.2	28.5	32.7	38.8	43.7	50.1
10 min	27.4	30.6	39.8	45.8	52.0	61.6	68.9	78.9
15 min	34.5	38.8	50.4	57.5	66.0	78.1	87.4	100.0
20 min	39.6	44.6	58.2	66.4	76.3	90.6	101.4	115.9
25 min	43.5	49.0	64.4	73.1	84.6	100.4	112.6	128.9
30 min	46.7	52.7	69.4	78.8	91.3	108.5	121.8	139.3
45 min	53.5	60.5	80.0	91.2	106.0	126.4	142.0	162.5
1 hr	58.2	66.0	87.6	100.0	116.1	138.6	156.1	178.8
1.5 hr	61.9	70.1	93.8	107.7	124.4	148.7	167.8	192.7
2 hr	64.5	73.1	98.2	113.5	130.1	155.6	175.5	201.7
3 hr	68.4	77.7	105.2	122.6	138.7	165.9	188.0	214.9
4.5 hr	72.7	82.5	113.1	132.8	151.7	177.8	200.6	229.5
6 hr	76.4	87.2	120.1	142.1	163.5	191.2	213.1	242.9
9 hr	82.5	94.4	131.8	157.6	182.8	217.1	243.5	278.3
12 hr	87.0	99.8	141.0	169.8	198.5	238.0	268.9	308.4
18 hr	95.8	109.4	156.9	191.1	225.9	274.4	313.4	359.2
24 hr	104.0	117.8	169.6	208.1	248.0	304.3	350.2	400.2
30 hr	111.9	127.1	182.1	224.8	269.5	333.3	385.8	440.6
36 hr	118.5	134.9	192.8	239.3	288.4	358.8	417.3	478.3
48 hr	129.2	148.0	210.4	263.5	319.9	402.0	470.9	540.7
72 hr	143.9	164.4	237.0	299.3	367.2	466.9	552.4	636.0
96 hr	152.1	175.4	256.1	325.5	401.3	513.9	611.2	703.2
120 hr	158.9	182.2	271.7	345.9	427.3	549.1	654.7	751.3
144 hr	163.0	186.3	285.0	362.6	447.9	575.9	687.2	789.3
168 hr	165.8	194.3	296.6	376.9	464.3	596.8	711.4	821.3

Appendix D: ARR Datahub files - Laidley Creek catchment

```
Results - ARR Data Hub
[STARTTXT]
Input Data Information
[INPUTDATA]
Latitude -27.727
Longitude 152.355
[END_INPUTDATA]
River Region
[RIVREG]
Division North East Coast
River Number 43
River Name Brisbane River
[RIVREG_META]
Time Accessed 08 September 2024 10:12AM
Version 2016_v1
[END_RIVREG]
ARF Parameters
[LONGARF]
Zone East Coast North
      0.327
      0.241
      0.448
       0.36
       0.00096
```

```
0.48
       -0.21
g
       0.012
       -0.0013
[LONGARF_META]
Time Accessed 08 September 2024 10:12AM
Version 2016_v1
[END_LONGARF]
Storm Losses
[LOSSES]
   12400
Storm Initial Losses (mm) 25
Storm Continuing Losses (mm/h) 1.4
[LOSSES_META]
Time Accessed 08 September 2024 10:12AM
Version 2016_v1
[END_LOSSES]
Temporal Patterns
[TP]
code ECnorth
Label East Coast North
[TP_META]
Time Accessed 08 September 2024 10:12AM
Version 2016_v2
[END_TP]
```

```
Areal Temporal Patterns
[ATP]
code
        ECnorth
arealabel
                East Coast North
[ATP_META]
Time Accessed 08 September 2024 10:12AM
Version 2016_v2
[END_ATP]
Median Preburst Depths and Ratios
[PREBURST]
min (h)\AEP(%) 50
                        20
                                10
                                        5
                                                 2
                                                         1
                                                 2.3 (0.041)
60 (1.0)
                0.2 (0.005)
                                1.4 (0.031)
                                                                 3.1 (0.047)
6.2 (0.079)
                8.5 (0.097)
90 (1.5)
                0.1 (0.002)
                                0.7 (0.013)
                                                 1.1 (0.017)
                                                                 1.4 (0.020)
6.7 (0.076)
                10.6 (0.107)
120 (2.0)
                0.0 (0.000)
                                0.7 (0.012)
                                                 1.1 (0.016)
                                                                 1.5 (0.019)
                14.7 (0.137)
9.0 (0.095)
                0.0 (0.000)
                                0.8 (0.013)
                                                 1.3 (0.018)
                                                                 1.8 (0.021)
180 (3.0)
13.5 (0.128)
                22.2 (0.187)
                                                                 3.2 (0.030)
360 (6.0)
                0.2 (0.004)
                                1.5 (0.020)
                                                 2.3 (0.026)
10.6 (0.084)
                16.2 (0.114)
720 (12.0)
                0.5 (0.007)
                                6.3 (0.067)
                                                10.1 (0.090)
                                                                 13.8 (0.106)
18.2 (0.116)
                21.4 (0.121)
1080 (18.0)
                0.0 (0.000)
                                7.0 (0.065)
                                                11.6 (0.090)
                                                                 16.1 (0.107)
20.5 (0.114)
                23.9 (0.116)
                                                 6.6 (0.046)
                                                                 9.1 (0.055)
1440 (24.0)
                0.0 (0.000)
                                4.0 (0.033)
13.5 (0.067)
                16.8 (0.073)
2160 (36.0)
                0.0 (0.000)
                                1.4 (0.010)
                                                 2.4 (0.014)
                                                                 3.3 (0.017)
6.7 (0.028)
                9.3 (0.034)
                                0.5 (0.003)
                                                 0.9 (0.005)
                                                                 1.2 (0.005)
2880 (48.0)
                0.0 (0.000)
3.2 (0.012)
                4.7 (0.015)
                                0.0 (0.000)
                                                 0.0 (0.000)
                                                                 0.0 (0.000)
4320 (72.0)
                0.0 (0.000)
0.9 (0.003)
                1.5 (0.004)
[PREBURST_META]
Time Accessed 08 September 2024 10:12AM
Version 2018_v1
```

Note Preburst interpolation methods for catchment wide preburst has been slightly altered. Point values remain unchanged.

[END_PREBURST]From preburst class

10% Preburst Depths

[PREBURST10]

min (h)\AEP(%)	50 20	10 5	2 1	
60 (1.0) 0.0 (0.000)	0.0 (0.000) 0.0 (0.000)	0.0 (0.000)	0.0 (0.000)	0.0 (0.000)
90 (1.5) 0.0 (0.000)	0.0 (0.000) 0.0 (0.000)	0.0 (0.000)	0.0 (0.000)	0.0 (0.000)
120 (2.0) 0.0 (0.000)	0.0 (0.000) 0.0 (0.000)	0.0 (0.000)	0.0 (0.000)	0.0 (0.000)
180 (3.0) 0.0 (0.000)	0.0 (0.000) 0.0 (0.000)	0.0 (0.000)	0.0 (0.000)	0.0 (0.000)
360 (6.0) 0.0 (0.000)	0.0 (0.000) 0.0 (0.000)	0.0 (0.000)	0.0 (0.000)	0.0 (0.000)
720 (12.0) 0.0 (0.000)	0.0 (0.000) 0.0 (0.000)	0.0 (0.000)	0.0 (0.000)	0.0 (0.000)
1080 (18.0) 0.0 (0.000)	0.0 (0.000) 0.0 (0.000)	0.0 (0.000)	0.0 (0.000)	0.0 (0.000)
1440 (24.0) 0.0 (0.000)	0.0 (0.000)	0.0 (0.000)	0.0 (0.000)	0.0 (0.000)
2160 (36.0)	0.0 (0.000) 0.0 (0.000)	0.0 (0.000)	0.0 (0.000)	0.0 (0.000)
0.0 (0.000) 2880 (48.0)	0.0 (0.000) 0.0 (0.000)	0.0 (0.000)	0.0 (0.000)	0.0 (0.000)
0.0 (0.000) 4320 (72.0) 0.0 (0.000)	0.0 (0.000) 0.0 (0.000) 0.0 (0.000)	0.0 (0.000)	0.0 (0.000)	0.0 (0.000)
[PREBURST10_MET	7			

Time Accessed 08 September 2024 10:12AM

Version 2018_v1

Note Preburst interpolation methods for catchment wide preburst has been slightly altered. Point values remain unchanged.

[END_PREBURST10]From preburst class

25% Preburst Depths

[PREBURST25]

min (h)\AEP(%) 50 20 10 5 2 1

60 (1.0)	0.0 (0.000)	0.0 (0.000)	0.0 (0.000)	0.0 (0.000)
0.3 (0.004) 90 (1.5)	0.5 (0.006) 0.0 (0.000)	0.0 (0.000)	0.0 (0.000)	0.0 (0.000)
0.3 (0.003)	0.5 (0.005)	(01000)	(01000)	(01000)
120 (2.0)	0.0 (0.000)	0.0 (0.000)	0.0 (0.000)	0.0 (0.000)
0.1 (0.001) 180 (3.0)	0.2 (0.002) 0.0 (0.000)	0.0 (0.000)	0.0 (0.000)	0.0 (0.000)
0.2 (0.002)	0.4 (0.003)			
360 (6.0) 0.2 (0.001)	0.0 (0.000) 0.3 (0.002)	0.0 (0.000)	0.0 (0.000)	0.0 (0.000)
720 (12.0)	0.0 (0.000)	0.0 (0.000)	0.0 (0.000)	0.0 (0.000)
1.5 (0.010)	2.7 (0.015)	0.0.40.000	0.3 (0.003)	0.5 (0.003)
1080 (18.0) 1.9 (0.010)	0.0 (0.000) 2.9 (0.014)	0.2 (0.002)	0.3 (0.003)	0.5 (0.003)
1440 (24.0)	0.0 (0.000)	0.0 (0.000)	0.0 (0.000)	0.0 (0.000)
0.0 (0.000) 2160 (36.0)	0.0 (0.000) 0.0 (0.000)	0.0 (0.000)	0.0 (0.000)	0.0 (0.000)
0.0 (0.000)	0.0 (0.000)	0.0 (0.000)	0.0 (0.000)	0.0 (0.000)
2880 (48.0)	0.0 (0.000)	0.0 (0.000)	0.0 (0.000)	0.0 (0.000)
0.0 (0.000) 4320 (72.0)	0.0 (0.000) 0.0 (0.000)	0.0 (0.000)	0.0 (0.000)	0.0 (0.000)
0.0 (0.000)	7	(31000)	0.0 (0.000)	0.0 (0.000)
[PREBURST25_ME	TA]			

Time Accessed 08 September 2024 10:12AM

Version 2018_v1

Note Preburst interpolation methods for catchment wide preburst has been slightly altered. Point values remain unchanged.

[END_PREBURST25]From preburst class

75% Preburst Depths

[PREBURST75]

min (h)\AEP(%)	50	20	10	5	2	1		
60 (1.0)	4.8 (0.		18.1 (0	.389)	27.0 (0.480)	35.4	(0.540)
32.8 (0.419) 90 (1.5)	30.8 (0	100)	12.7 (0	. 244)	18.7 (9.298)	24.5	(0.333)
46.1 (0.524) 120 (2.0)	62.3 (0	064)	11.2 (0	.199)	16.9 (0.250)	22.4	(0.283)
64.7 (0.681) 180 (3.0)	96.3 (0 3.0 (0.	068)	22.4 (0	.359)	35.3 (6	0.470)	47.6	(0.542)
71.9 (0.684) 360 (6.0)	90.2 (0 14.8 (0		29.6 (0	.393)	39.5 (6	0.437)	48.9	(0.465)
65.5 (0.520) 720 (12.0)	78.0 (0 15.1 (0	,	28.7 (0	.307)	37.7 (6	3.338)	46.4	(0.357)
72.3 (0.463) 1080 (18.0)	91.8 (0 16.0 (0		33.0 (0	.307)	44.2 (6	0.344)	55.0	(0.367)

```
65.4 (0.362)
               73.3 (0.356)
1440 (24.0)
               10.5 (0.122)
                               21.0 (0.176)
                                             28.0 (0.196)
                                                             34.6 (0.207)
1440 (24.0) 10.5 (0.122)
58.7 (0.290) 76.7 (0.332)
             10.9 (0.111)
2160 (36.0)
                              16.4 (0.119)
                                              20.0 (0.120)
                                                              23.4 (0.120)
31.2 (0.131) 37.0 (0.135)
             2.1 (0.019)
2880 (48.0)
                               8.2 (0.054)
                                              12.2 (0.066)
                                                              16.0 (0.073)
30.5 (0.114)
              41.4 (0.134)
               0.0 (0.000)
                              4.9 (0.029)
                                              8.2 (0.039)
                                                              11.3 (0.045)
4320 (72.0)
             26.5 (0.073)
20.0 (0.064)
[PREBURST75_META]
```

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Version 2018_v1

Note Preburst interpolation methods for catchment wide preburst has been slightly altered. Point values remain unchanged.

[END_PREBURST75]From preburst class

90% Preburst Depths

[PREBURST90]

min (h)\AEP(%)	50 20	10 5	2 1	
60 (1.0) 105.2 (1.343)	15.7 (0.481) 96.0 (1.089)	60.0 (1.289)	89.4 (1.592)	117.5 (1.791)
90 (1.5) 151.7 (1.724)	19.9 (0.540) 180.5 (1.818)	60.6 (1.160)	87.5 (1.390)	113.3 (1.539)
120 (2.0) 158.2 (1.667)	27.1 (0.679) 213.0 (1.989)	52.3 (0.929)	69.0 (1.018)	85.1 (1.072)
180 (3.0) 139.1 (1.323)	27.2 (0.611) 143.3 (1.207)	73.4 (1.175)	104.0 (1.385)	133.4 (1.519)
360 (6.0) 164.0 (1.301)	39.6 (0.728) 198.1 (1.389)	74.0 (0.982)	96.8 (1.072)	118.6 (1.127)
720 (12.0) 156.6 (1.002)	35.7 (0.526) 195.8 (1.105)	65.5 (0.701)	85.3 (0.763)	104.2 (0.801)
1080 (18.0) 119.0 (0.658)	31.5 (0.404) 129.6 (0.629)	63.4 (0.590)	84.6 (0.658)	104.8 (0.699)
1440 (24.0) 114.9 (0.568)	34.9 (0.406) 133.8 (0.579)	58.8 (0.493)	74.6 (0.522)	89.7 (0.537)
2160 (36.0) 90.2 (0.378)	37.1 (0.376) 118.2 (0.431)	43.9 (0.318)	48.4 (0.291)	52.7 (0.270)
2880 (48.0) 86.6 (0.323)	27.1 (0.252) 111.8 (0.361)	38.4 (0.252)	45.8 (0.248)	53.0 (0.243)
4320 (72.0) 65.1 (0.208)	9.5 (0.079) 79.0 (0.218)	25.7 (0.149)	36.3 (0.172)	46.6 (0.184)
[PREBURST90_MET	Ά]			

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Appendix E: RORB catchment file

C Laidley	Creek										
C #FILE C	OMMENTS										
C											
C #SUB-AR C 0 C	EA AREA COMMENT	S									
C #IMPERV	IOUS FRACTION C	OMM	ENTS								
CTF	OUND IMAGE										
C #NODES											
c 1	95.000 13.918520		95.000 0.000000	0	0	0	1.000	1	0	3	Subcatchment
C 36 C 1 1 C 2 C 2	92.891 7.407192		53.309 0.000000	0	0	0	1.000	1	0	3	Subcatchment
0.000000	90.484	0	0 0 60.233				1.000	0	0	4	
C 4 0.000000	88.047 0.000000	0	0 0 60.204				1.000	0	0	6	
C 5 3 C	90.239 12.661318		46.882 0.000000	0	0	0	1.000	1	0	4	Subcatchment
C 6	86.814 0.000000	0	0 0 60.197				1.000	0	0	8	
C 7	89.535 3.893288		87.993 0.000000	0	0	0	1.000	1	0	6	Subcatchment
C 7 4 C 8 5	85.356 8.606579		60.283 0.000000	0	0	0	1.000	1	0	9	Subcatchment
0.000000	83.915 0.000000	0	0 0 60.283				1.000	0	0	11	
C C 10 6	84.906 4.189867		80.382	0	0	0	1.000	1	0	9	Subcatchment
C 11 0.000000	76.262 0.000000	0	0 0 60.167				1.000	0	0	12	
C C 12 8	73.371 23.696005		60.198 0.000000	0	0	0	1.000	1	0	14	Subcatchment
C 13 7 C 14 0.000000	86.367 9.419337		16.621 0.000000	0	0	0	1.000	1	0	11	Subcatchment
0.000000	0.000000	0	0 0 60.284				1.000	0	0	16	
C C 15 9 C	77.910 14.119680		23.860 0.000000	0	0	0	1.000	1	0	14	Subcatchment
C 16	65.765 0.000000	0	0 0 60.370				1.000	0	0	18	
C C 17	74.307 28.255606		13.842	0	0	0	1.000	1	0	16	Subcatchment
10 C C 18	57.186 56.592533		60.082 0.000000	0	0	0	1.000	1	0	19	Subcatchment
C 19	51.063		59.987				1.000	0	0	20	

0.000000 C Mulgowie	0.000000	0	0	0							
C 20	42.602 41.276543		0	60.072	0	0	0	1.000	1 0	21	Subcatchment
0.000000	31.585 0.000000	0	0	60.486				1.000	0 0	22	
C C 22 13	25.249 23.390030		0	60.385	0	0	0	1.000	10	23	Subcatchment
C 23 0.000000	16.911 0.000000	0	0	60.495				1.000	0 0	24	
C 24 0.000000	14.714 0.000000	0	0	60.375				1.000	0 0	28	
C C 25	48.329 25.320984		0	92.470 .000000	0	0	0	1.000	1 0	26	Subcatchment
C 26 0.000000	33.685 0.000000	0	0	92.513				1.000	0 0	27	
C C 27	27.167 18.146849		0	92.641 .000000	0	0	0	1.000	1 0	23	Subcatchment
C C 28 20 C	10.724 17.716366		0	60.169 .000000	0	0	0	1.000	1 0	36	Subcatchment
C 29	70.865 48.529217		0	5.000	0	0	0	1.000	1 0	31	Subcatchment
C C 30 17 C	49.582 7.546566		0	13.393 .000000	0	0	0	1.000	1 0	31	Subcatchment
0.000000 C	41.564 0.000000	0	0	25.328 0				1.000	0 0	32	
C 32 18 C	32.689 43.470638		0	25.617 .000000	0	0	0	1.000	1 0	33	Subcatchment
C 33 0.000000 C	0.000000	0	0	33.806				1.000	0 0	34	
C 34 19 C	18.898 44.153233		0	46.223 .000000	0	0	0	1.000	1 0	24	Subcatchment
C 35 19	3.558 44.153233		0	60.252	70	0	0	1.000	0 1	0	Subcatchment
C 36 0.000000 C	4.507 0.000000	0	0	67.084				1.000	0 0	35	
C #REACHES	5										
C 1 2.550	1 0 92.378			1	3				0 1	0	4.078
C C C 2 4.793	75.406			2	3				0 1	0	1.064
C C C 3	91.490 57.147			3	4				0 1	0	1.780
C	1 0 89.161 60.233			.T.					(SILE)		857.5.2.7.7
C 4 2.853	1 0			5	4				0 1	0	2.243
C C	89.284										

C 5	52.754	4	6	0 1 0	0.950
C 5 1.368 C C 6 2.534 C 7 0.576	1 0 87.452				
C 6	60.233	7	6	0 1 0	2.013
C C	88.523 76.680			0.1.0	1 563
C 7 0.576 C	1 0 86.194	6	8	0 1 0	1.563
C C C C 8	60.197	8	9	0 1 0	1.866
C C	84.669 60.240				1 205
C 9 3.142 C	1 0 84.542	10	9	0 1 0	1.305
C C 10	71.321	9	11	0 1 0	5.228
C 9 3.142 C 10 0.574 C 11 0.476 C 11 0.476 C 12 2.507 C 13 0.384 C 14 2.191 C 15 0.277 C 15	81.512 60.152		12	0.1.0	1 050
C 11 0.476 C	1 0 74.259	11	12	0 1 0	1.050
C 12	60.112	13	11	0 1 0	4.746
C C	80.393 39.813	12	14	0.1.0	2 205
C 13 0.384 C	1 0 70.959	12	14	0 1 0	3.385
C 14	60.284	15	14	0 1 0	5.797
C C	74.073 39.620	14	16	0.1.0	2 800
C 15 0.277 C	1 0 67.256	14	16	0 1 0	2.890
C C 16 1,275	60.284	17	16	0 1 0	6.668
C C C 17	70.639 33.714	16	18	0 1 0	5.480
0.383 C	1 0 60.353	10	10	0.10	3.400
C 17 0.383 C 18 0.147 C 19 0.369	1 0	18	19	0 1 0	4.756
C C C 19	54.012 60.029	19	20	0 1 0	5.150
0.369 C	1 0 46.456	23			
C C 20 0.280	59.987	20	21	0 1 0	6.420
C C C 21	35.248 60.085	21	22	0 1 0	3.160
0.222 C	1 0 28.097	21	22	010	3.100
C 20 20 0.280 C 21 0.222 C C 22 0.193	60.343	22	23	0 1 0	6.732
	774 N.S.				

	2300 2W0847						
C C C 23 0.000	20.417						
C 23	60.471	23	24	0	1 (0	0.364
0.000	1 0	23	24		- '		0.304
C C C 24 0.351	15.812						
C 34	60.334	25	26				F 000
C 24	1 0	25	26	0	1 (5.990
C	41.201						
C	92.469	26	27				0 005
C 25 0.578 C 26 0.078 C 27 0.432 C 27 0.432 C 28 0.834 C 29 1.021 C 30 0.548 C 29 1.021 C 30 0.548 C 31 0.299 C 32 0.300	1 0	26	27	0	1 (0	0.865
C	30.379						
C	92.638						
0 078	1 0	27	23	0	1 (0	5.109
C.078	22.677						
c	78.904						
C 27	1 0	24	28	0	1 (0	1.622
0.432	1 0 13.123						
c	60.210						
C 28		29	31	0	1 (0 1	11.985
0.834	1 0						
č	53.782 17.230						
C 29		30	31	0	1 (0	3.623
1.021	1 0						
c	45.880 18.715						
C 30		31	32	0	1 (0	2.918
0.548	1 0						
C	36.942 25.328						
C 31	23.320	32	33	0	1 (0	5.685
0.299	1 0						
C	27.587 29.371						
C 32	29.3/1	33	34	0	1 (0	3.328
0.300	1 0			55		5.6	
C	21.296						
0.300 C C C 33 0.147 C C C 0.000	39.499	34	24	0	1 (0	4.090
0.147	1 0	31			_		4.050
C	16.462						
C 34	54.267	28	36	0	1 /	0	2 211
0.000	1 0	20	30	U	1 (,	3.311
C	7.615						
C 35	63.627	36	25	0	1 /	0	0 000
C 35	1 0	36	35	U	1 (9	0.000
C	4.032						
C	63.668						
C C #STORAG C O C C #INFLOW C O C C END ROR	ES						
C #STORAG	123						
C	(facement as						
C #INFLOW	/OUTFLOW						
C							
C END ROR	B_GE						
C 1							
1, 4.078,	-99			,Reach 1	no	de 1	
Sub-area	Subcatchment 1, Reac	h - Gene	erate r	ainfall exces	s h	'graph and	route
downstrea							
Store run	ning hydrograph						
1. 1.064.	-99			,Reach 2	no	de 2	
Sub-area	Subcatchment 2, Reac	h - Gene	erate r	ainfall exces	s h	'graph and	route

```
downstream
Add running h'graph to last stored h'graph
5, 1.780, -99
Reach - Route running h'graph downstream
                                                                           ,Reach 3
Store running hydrograph
1, 2.243, -99
Sub-area Subcatchment 3, Reach - Generate rainfall excess h'graph and route
downstream
Add running h'graph to last stored h'graph
5, .950, -99
Reach - Route running h'graph downstream
                                                                           ,Reach 5
Store running hydrograph
1, 2.013, -99
Sub-area Subcatchment 4, Reach - Generate rainfall excess h'graph and route
downstream
Add running h'graph to last stored h'graph
5, 1.563, -99
Reach - Route running h'graph downstream
                                                                           .Reach 7
2, 1.866, -99
                                                                            Reach 8 node 8
Sub-area Subcatchment 5, Reach - Generate rainfall excess h'graph, add to running h'graph, and route downstream
Store running hydrograph
1, 1.305, -99 ,Reach 9 node 10
Sub-area Subcatchment 6, Reach - Generate rainfall excess h'graph and route
downstream
Add running h'graph to last stored h'graph 5, 5,228, -99
5, 5.228, -99
Reach - Route running h'graph downstream
                                                                           ,Reach 10
Store running hydrograph
1, 4.746, -99 ,Reach 12 node 13
Sub-area Subcatchment 7, Reach - Generate rainfall excess h'graph and route
downstream
Add running h'graph to last stored h'graph
5, 1.050,
               -99
                                                                           ,Reach 11
Reach - Route running h'graph downstream
2, 3.385, -99 ,Reach 13 node 12
Sub-area Subcatchment 8, Reach - Generate rainfall excess h'graph, add to
running h'graph, and route downstream
Store running hydrograph
1, 5.797, -99 ,Reach 14 node 15
Sub-area Subcatchment 9, Reach - Generate rainfall excess h'graph and route
downstream
Add running h'graph to last stored h'graph
5, 2.890, -99
Reach - Route running h'graph downstream
                                                                           ,Reach 15
Store running hydrograph
1, 6.668, -99 ,Reach 16 node 17
Sub-area Subcatchment 10, Reach - Generate rainfall excess h'graph and
route downstream
Add running h'graph to last stored h'graph
5, 5.480, -99
Reach - Route running h'graph downstream
                                                                          Reach 17
2, 4.756, -99
,Reach 18 node 10
Sub-area Subcatchment 11, Reach - Generate rainfall excess h'graph, add to running h'graph, and route downstream
,Reach 19
,Reach 19
5, 5.150, -99
Reach - Route running h'graph downstream
2, 6.420,
                                                                          ,Reach 20 node 20
```

```
Sub-area Subcatchment 12, Reach - Generate rainfall excess h'graph, add to
running h'graph, and route downstream 5, 3.160, -99
                                                                     ,Reach 21
Reach - Route running h'graph downstream
2, 6.732, -99 ,Reach 22 node 22
Sub-area Subcatchment 13, Reach - Generate rainfall excess h'graph, add to
running h'graph, and route downstream
Store running hydrograph
1, 5.990, -99 ,Reach 24 node 25
Sub-area Subcatchment 14, Reach - Generate rainfall excess h'graph and
route downstream
5, .865, -99
                                                                     ,Reach 25
Reach - Route running h'graph downstream
2, 5.109, -99
,Reach 26 node 27
Sub-area Subcatchment 15, Reach - Generate rainfall excess h'graph, add to
Add running h'graph to last stored h'graph
5, .364, -99
Reach - Route running h'graph downstream
                                                                     ,Reach 23
Store running hydrograph
1, 11.985, -99
                                                                      Reach 28 node 29
Sub-area Subcatchment 16, Reach - Generate rainfall excess h'graph and
route downstream
Store running hydrograph
1, 3.623, -99 ,Reach 29 node 30
Sub-area Subcatchment 17, Reach - Generate rainfall excess h'graph and
route downstream
Add running h'graph to last stored h'graph 5, 2.918, -99
5, 2.918, -99
Reach - Route running h'graph downstream
                                                                     ,Reach 30
2, 5.685, -99

Sub-area Subcatchment 18, Reach - Generate rainfall excess h'graph, add to running h'graph, and route downstream

5, 3.328, -99

Reach - Route running h'graph downstream

Reach - Route running h'graph downstream
2, 4.090, -99
,Reach 33 node 34
Sub-area Subcatchment 19, Reach - Generate rainfall excess h'graph, add to running h'graph, and route downstream
Add running h'graph to last stored h'graph 5, 1.622, -99
Reach - Route running h'graph downstream
                                                                     ,Reach 27
2, 3.311, -99 ,Reach 34 node 28
Sub-area Subcatchment 20, Reach - Generate rainfall excess h'graph, add to running h'graph, and route downstream ,Reach 35
Reach - Route running h'graph downstream
                     7.1 in FIT run
PRINT
C Sub Area Data
C Areas, km**2, of subareas A,B...
                                        12.66132,
23.69600,
     13.91852,
                         7.40719,
                                                              3.89329,
                                                                                 8.60658,
      4.18987,
                         9.41934,
                                                             14.11968,
                                                                                28.25561,
     56.59253,
                        41.27654,
                                           23.39003,
                                                             25.32098,
                                                                                18.14685,
     48.52922.
                         7.54657.
                                           43.47064.
                                                             44.15323.
                                                                                17.71637.
 -99
C Impervious Fraction Data
                                                                  ,No impervious areas in
system
```

Appendix F: Calibration storm event files

F1: May 1996 event

Calibration summary

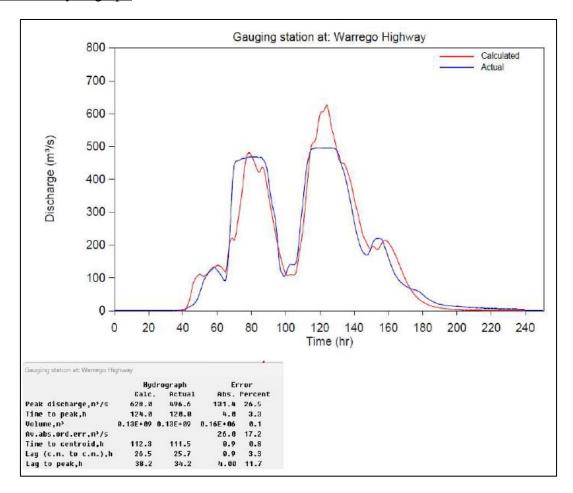
Event ID	May-96
Start Date	30/04/1996 17:00
End Date	6/05/1996 10:00
Duration (hours)	0:00
Number of peaks	2
Maximum gauged	
discharge at 143229A	
(m3/s)	496.592
Supplementary peak	
discharge (m3/s)	470.6

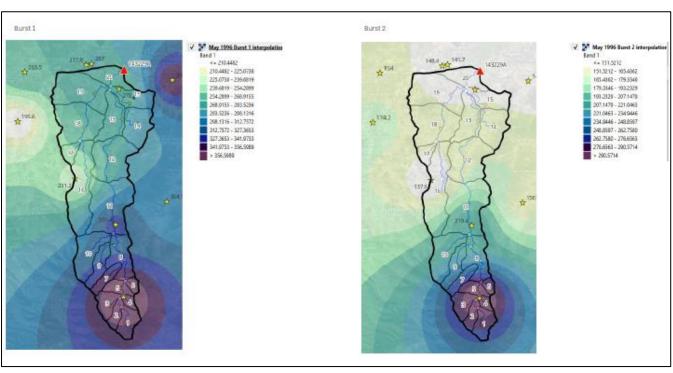
	Tota	l event rainfall record		
Station name	Station ID	Recorded rainfall Burst 1 (mm)	Recorded rainfall Burst 2 (mm)	Notes
Forest Hill	040079	270	151.1	
Upper Tenthill	040388	195.6	174.2	
UQ Gatton	040082	287	141.7	
Franklyn Vale	040374	304.5	156	
Placid Hills	040449	265.5	154	
Townson	040675	371.4	304.8	
Gatton DAFF	040436	277.8	148.4	
Hattonyale	040095	355	143,4	
Mt Berryman	040310	231.2	137.6	
Thornton	040751	331.4	219.4	
Average proportion		0.636	0.364	

Subcatchment ID	Rainfall depth ocatchment ID Burst 1 (mm)	
1	405.9	232.
2	415.7	238.4
3	416.9	239.
4	426.4	244.5
5	423.0	242.5
6	415.4	238.
7	390,6	224.0
8	351.8	201.
9	360.7	206.8
10	337.1	193.3
11	325.9	186.
12	275.8	158.
13	277.2	159.0
14	282.7	162.
15	274.6	157.
16	253.0	145.
17	256.8	147.2
18	270.9	155.4
19	274.2	157.3
20	272.3	156.

	RORB gauged data statistics							
Peak flow (m3/s)	Total flow volume (m3)	Time to peak (hrs)	Time to centroid (hrs)					
496.6	1.30E+08	120	111.5					

1	Error (%)	Time to centroid (hrs)	Error (%)	Time to peak (hrs)	Error (%)	Total flow volume (m3)	Error (%)	Peak flow (m3/s)	CL_2	IL_2	CL_1	IL_1	kc	m
3	-12.3	97.8	-9.2	109	0.1	1.30E+08	24.5	618.3	0.07	0	2.66	100	27	0.8
9	-11.9	98.2	-9.2	109	0.1	1.30E+08	22.5	609	0.07	0	2.66	100	28	0.8
3	-13.3	96.6	-9.2	109	0.1	1.30E+08	26.4	627.7	0.07	0	2.98	80	26	0.8
9	-11.9	98.1	-9.2	109	0.1	1.30E+08	26.5	628.4	0.07	0	2.22	120	26	0.8
5	-12.5	97.5	-9.2	109	0.1	1.30E+08	27.3	632.2	0.02	2	2.66	100	26	0.8
7 Best s	-12.7	97.3	-9.2	109	0.1	1.30E+08	26.5	628	0.07	0	2.66	100	26	8.0
8 15 hr	0.8	112.3	3.3	124	0.1	1.60E+05	26.5	628	0.07	0	2.66	100	26	0.8





```
30 Apr - 7 May 1996
  FIT
                      time inc = 1 hr, calcs for 250 incs, 2 bursts
 C 2 pluvios, non uniform pattern
1,250,2,2,1,-99,
0,84,85,138, rain from 0 - 138 time incs,
 2,2,3,3,2,3,1,3,3,1,3,2,1,1,3,2,2,6,1,3,3,8,8,4,4,1,5,8,5,7,11,3,7,3,4,6,5,8,5,1,0,1,3,0,0,0,1,2,0,4,6,21,9,9,6,2,8,13,11,14,7,12,5,6,4,11,6,10,8,1,1,0,0,0,1,0,2,0,0,0,1,4,0,2,2,0,1,1,1,0,1,1,0,1,1,3,4,2,3,2,3,6,7,4,4,6,4,1,1,0,0,8,10,5,1,1,0,2,1,0,2,0,0,2,1,0,0,0,1,1,1,3,1,2,15,1,-99,mm/hr

Tenthill
 Tenthill
2,2,1,1,2,2,1,1,3,2,7,3,1,2,3,1,2,0,1,5,4,6,3,1,3,1,6,2,3,5,3,2,3,4,1,2,3,3,0,2,0,1,0,0,0,0,0,0,0,0,11,2,2,0,2,5,4,3,5,5,4,2,2,0,1,0,1,1,1,4,3,1,0,2,0,2,0,1,0,0,1,0,1,0,1,0,1,2,3,10,3,6,9,12,7,14,11,6,7,8,11,11,7,6,5,9,2,0,0,3,1,1,0,0,1,0,1,1,2,0,0,0,2,0,0,0,1,3,1,1,2,5,1,0,0,-99,mm/hr

c sub catchment rainfalls from isohyetal map
405.9,415.7,416.9,426.4,423,415.4,390.6,351.8,360.7,337.1,325.9,275.8,277.2,2
82.7,274.6,253,256.8,270.9,274.2,272.3,-99, Sub area event total rainfalls B1
232.7,238.4,239.1,244.5,242.5,238.1,224.0,201.7,206.8,193.2,186.8,158.1,159.0,162.1,157.4,145.0,147.3,155.4,157.2,156.1,-99 Sub area event total rainfalls B2

c pluivo reference numbers
 C Hydrograph data,
  ,15.721,15.128,14.668,14.217,13.766,13.372,12.981,12.589,12.194,11.795,11.396,10.997,10.671,10.35,10.029,9.713,9.404,9.094,8.785,8.52,8.259,7.999,7.761,7.525,7.289,7.053,6.833,6.653,6.472,6.292,6.112,5.936,5.764,5.592,5.419,5.247,5.112,4.978,4.844,4.733,4.623,4.514,4.405,4.3,4.21,-99,mm/hr
       Proportional volumes
  50555,75859,-99, any units
```

F2: November 2008 event

Event calibration summary

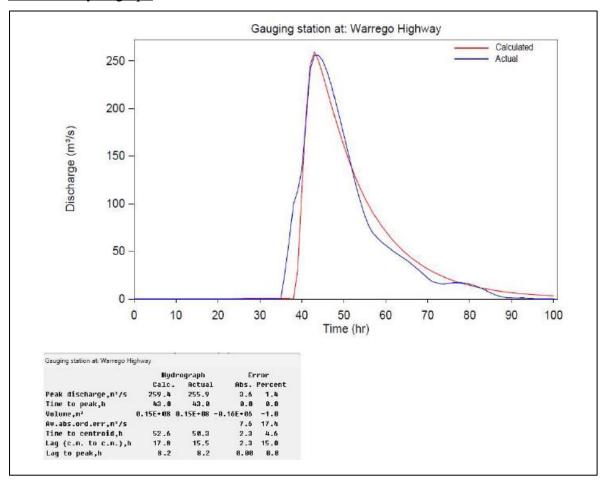
Event ID	Nov-08
Start Date	18/11/2008 10:00
End Date	20/11/2008 9:00
Duration (hours)	48
Number of peaks	1
Maximum gauged	
discharge at 143229A	
(m3/s)	255.9
Supplementary peak	
discharge (m3/e)	2/2

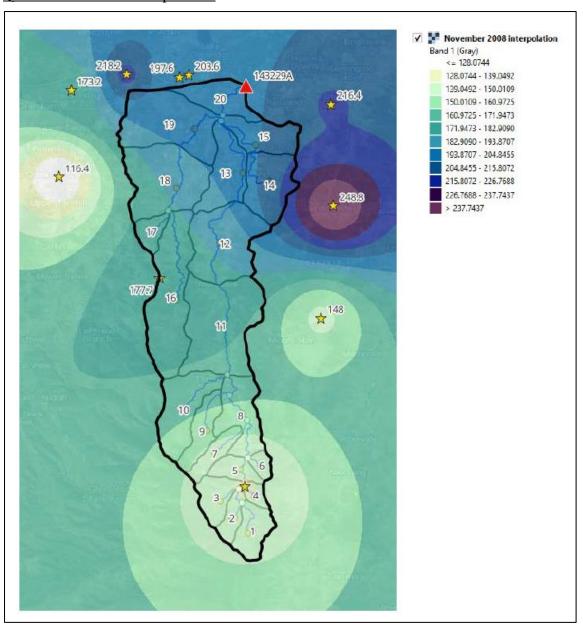
Station name	Station ID	Recorded rainfall (mm)	Notes	
Forest Hill	040079	129.9		
Mt Berryman	040310	177.7		
Upper Tenthill	040388	116.4		
UQ Gatton	040082	203.6		
Franklyn Vale	040374	148		
Franklyn Vale Alert	040912	171		
Gatton Allan St	040083	218.2		
Placid Hills	040449	173.2		
Townson	040675	138.6		
Bremer River	WMIP	108		
Gatton DAFF	040436	197.6		
Grandchester	040091	248.8		
Hattonvale	040095	216.4		

Subcatchment centroid event rainfall depths from QGIS interpolation				
Subcatchment ID	Rainfall depth (mm)			
1	144.3			
2	141.8			
3	141.4			
4	139.2			
5	139.5			
6	141.5			
7	146.8			
8	157.4			
9	154.8			
10	163.1			
11	174.5			
12	185.			
13	200.5			
14	210.			
15	204.5			
16	177.6			
17	177.4			
18	185.8			
19	197.8			
20	200.5			

	RORB gauged data statistics						
Peak flow (m3/s)	Total flow volume (m3)	Time to peak (hrs)	Time to centroid (hrs)				
255.9	1.50E+07	43	50.3				

m	kc	,IL :	CL	Peak flow (m3/s)	Error (%)	Total flow volume (m3)	Error(%)	Time to peak (hrs)	Error (%)	Time to centroid (hrs)	Error (%)
8.0	39	100	7.43	257	0.4	1.50E+07	-1.4	40	-7	49.1	-2.3
0.8	35	100	7.43	286.1	11.8	1.49E+07	-1.1	39	-9.3	47.8	-5
0.8	35	75	9.95	246.5	-3.7	1.49E+07	-0.9	39	-9.3	49.1	-2.2
0.8	35	77	9.78	250.8	-2	1.49E+07	-0.9	39	-9.3	49.1	-2.2
0.8	34	77	9.78	259.4	1.4	1.49E+07	-0.7	39	-9.3	48.7	-3
0.8	34	77	9.78	259.4	1.4	1.48E+07	-1	43	0	52.6	4.6





F3: January 2011 event

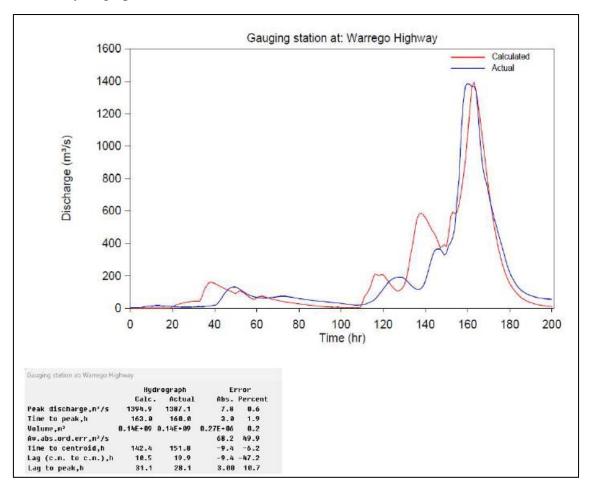
Calibration summary

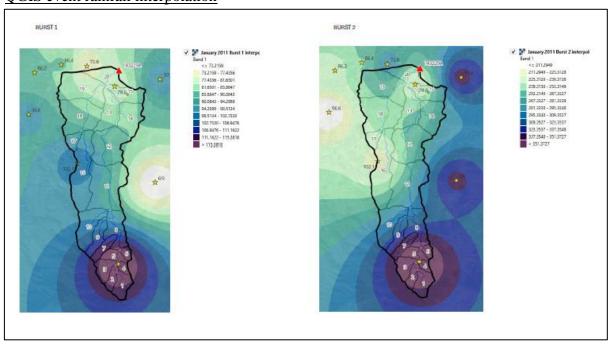
Total event rainfall record								
Station name	Station ID	Recorded rainfall Burst 1 (mm)	Recorded rainfall Burst 2 (mm)	Notes				
Forest Hill	040079	78.6	223.6					
UQ Gatton	040082	N/A	N/A	Record stops mid event				
Mt Berryman	040310	102.1	219.6	The control of the co				
Hattonvale	040095	97.4	343.2					
Placid Hills	040449	86.2	231.9					
Gatton Allan St	040083	88.4	253					
Upper Tenthill	040388	96.6	197					
Franklyn Vale	040374	69	341					
Townson	040675	119.6	365.6					
Gatton DAFF	040436	72.8	293.4					
Average proportion		0.252	0.748					

Subcatchment ID	Rainfall depth Burst 1 (mm)	Rainfall depth Burs 2 (mm)	
1	115.4	354.4	
2	117.2	359.3	
3	117.6	359.9	
4	119	364.2	
5	118.6	363	
6	117.2	360	
7	113.2	348	
8	102.5	326.1	
9	107	329.1	
10	100.7	302.2	
11	91.3	273	
12	91.1	254.2	
13	83.6	253.6	
14	85.1	267.1	
15	81.4	243.8	
16	100	228.5	
17	95.8	234.1	
18	87.8	248.2	
19	80.5	253	
20	79.3	241	

RORB gauged data statistics							
Peak flow (m3/s)	Total flow volume (m3)	Time to peak (hrs)	Time to centroid (hrs)				
1387.1	1.40E+08	160	151.8				

RORB calibrated parameters and statistics												
m	kc	IL1	CL1	IL 2	Peakflow (m3/s)	Error (%)	Total flow volume (m3)	Error (%)	Time to peak (hrs)	Error (%)	Time to centroid (hrs)	Error (%)
0.8	28	9	3.18	9	1328.6	4.2	1.40E+08	0.2	163	1.9	143.3	-5.6
0.8	27	9	3.18	9	1361.4	-1.9	1.40E+08	0.2	163	1.9	142.8	-5.9
0.8	27	9	3.18	15	1376	-0.8	1.40E+08	0.2	163	1.9	143.3	-5.6
0.8	28	9	3.18	9	1394.9	0.6	1.40E+08	0.2	163	1.9	142.4	-6.2
0.8	26	9	3.18	5	1385.1	-0.1	1.40E+08	0.2	163	1.9	142.1	-6.4





```
5-12 January 2011
C time inc = 1 hr, calcs for 250 incs, 2 bursts 2 pluvios, non uniform pattern 1,290,2,2,1,-99, 0,106,107,168, rain from 0 - 168 time incs Upper Tenthill
FIT
Bremer River
11.735,11.478,11.221,10.965,10.783,10.66,10.536,10.413,10.25,-99,m3/s C proportional flow volumes
16309,119743,-99, any units
```

F4: January 2013 event

Calibration summary

Jan-13 EventID Rainfall event characteristics Start date & time 25/01/2013 20:00 28/01/2013 9:00 End date & time Rainfall duration (hours) 61:00:00 Number of bursts

Runoff characteristics
Hydrograph duration 153-pos-153:00:00 1 Maximum gauged

discharge at 143229A

(m3/s) 1041.502

Supplementary peak

N/A discharge (m3/s)

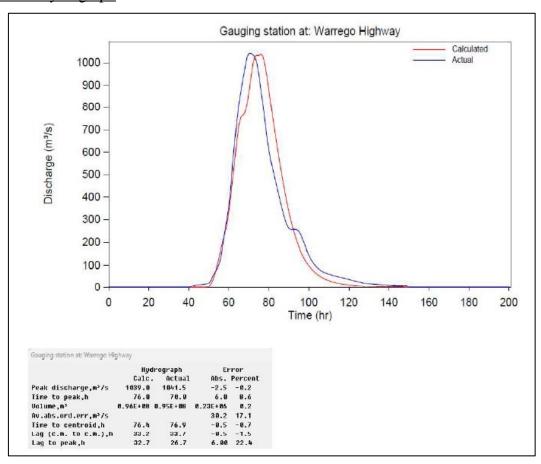
Station name	Station ID	rainfall (mm)	Notes
Forest Hill	040079	162.8	
Mt Berryman	040310	278	
Upper Tenthill	040388	200.2	
UQ Gatton	040082	227	
Franklyn Vale	040374	279	
Franklyn Vale Alert	040912	322.6	
Gatton Allan St	040083	n/a	
Placid Hills	040449	309.2	
Townson	040675	546.8	
Gatton DAFF	040436	220	
Grandchester	040091	308.8	

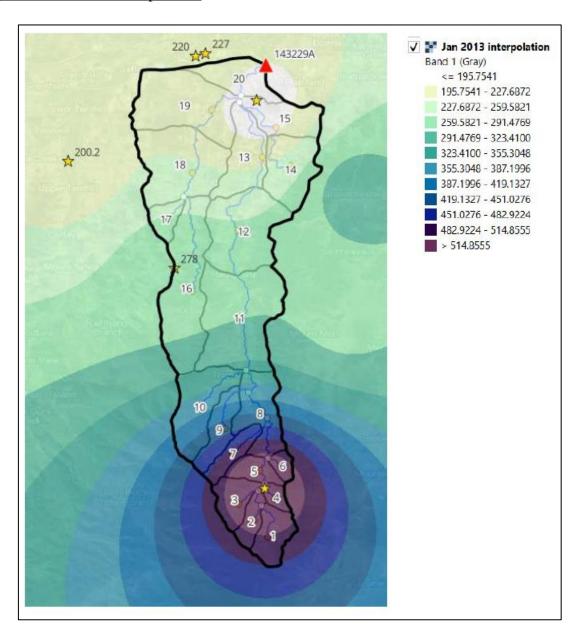
Subcatchment centroid event rainfall depths from QGIS interpolation				
Subcatchment ID	Rainfall depth (mm)			
1	524.			
2	534.			
3	535.			
4	544.			
5	541.			
6	536.			
7	510.			
8	453.			
9	465.			
10	395.			
11	293.			
12	253.			
13	197.			
14	205.			
15	179.			
16	277.			
17	258.			
18	225.			
19	202.			
20	18			

RORB gauged data statistics								
Peak flow (m3/s)	Total flow volume (m3)	Time to peak (hrs)	Time to centroid (hrs)					
1041.5	9.50E+07	70	76.9					

m	kc	IL	CL	Peak flow (m3/s)	Error (%)	Total flow volume (m3)	Error(%)	Time to peak (hrs)	Error (%)	Time to centroid (hrs)	Error (%)
0.8	36	80	0.45	1038.9	-0,3	9.54E+07	0.4	57	-18,6	60,2	-21.8
0.8	34	75	0.6	1045.1	0.3	9.53E+07	0.3	57	-18.6	59	-23.3
0.8	32	66	0.81	1036.6	-0,5	9.53E+07	0.3	56	-20	57.7	-25.1
0.8	32	60	0.94	1019.6	-2.1	9.53E+07	0.3	57	-18.6	57.3	-25.5
0.8	30	60	0.94	1039	-0.2	9.52E+07	0.2	56	-20	56.4	-26.6
0.8	30	60	0.94	1039	-0.2	9.52E+07	0.2	76	8.6	76.4	-0.7

slation





F5: March 2017 event

Calibration summary

Event ID Start Date End Date Duration (hours) Mar-17 30/03/20170:00 30/03/2017 23:00 23:00 Number of peaks Maximum gauged discharge at 143229A 1

(m3/s) Supplementary peak discharge (m3/s) 248.682 N/A

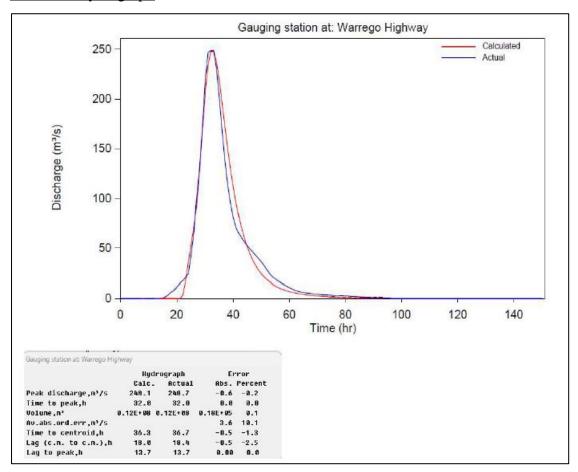
Forest Hill 040079 123 Mt Berryman 040310 123.4 Upper Tenthill 040388 81 UQ Gatton 040082 99 Franklyn Vale 040374 157.5 Franklyn Vale Alert 040912 163.2	
Upper Tenthill 040388 81 UQ Gatton 040082 99 Franklyn Vale 040374 157.5 Franklyn Vale Alert 040912 163.2	
UQ Gatton 040082 99 Franklyn Vale 040374 157.5 Franklyn Vale Alert 040912 163.2	
Franklyn Vale 040374 157.5 Franklyn Vale Alert 040912 163.2	
Franklyn Vale Alert 040912 163.2	
A CONTRACTOR OF THE PARTY OF TH	
Gatton Allan St 040083 95	
Placid Hills 040449 91.9	
Townson 040675 185	
Bremer River WMIP 219	

Subcatchment centroid event rainfall depths from QGIS interpolation				
Subcatchment ID	Rainfall depth (mm)			
1	177.7			
2	180.8			
3	181.3			
4	184			
5	183.2			
6	181.2			
7	173.8			
8	160.7			
9	163.2			
10	149.9			
11	138.1			
12	126.1			
13	120.1			
14	123.4			
15	120.7			
16	124.5			
17	118.3			
18	113			
19	111.3			
20	116.6			

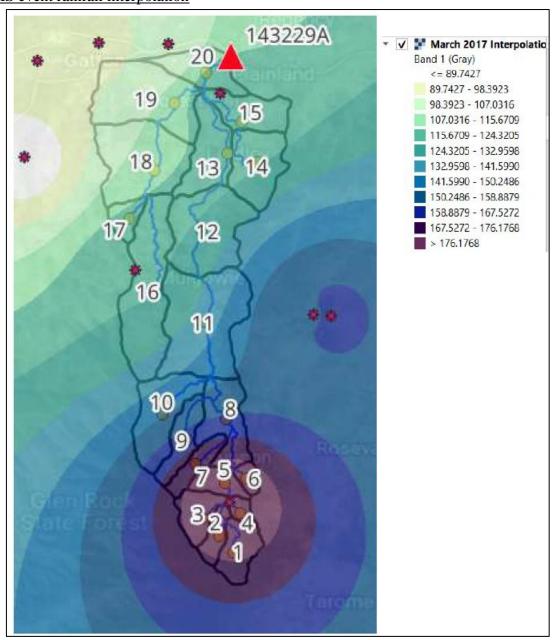
	RORB gauged data statistics								
Peak flow (m3/s)	Total flow volume (m3)	Time to peak (hrs)	Time to centroid (hrs)						
248.7	1.20E+07	32	36.7						

m	kc	IL	CL	Peak flow (m3/s)	Error (%)	Total flow volume (m3)	Error (%)	Time to peak (hrs)	Error (%)	Time to centroid (hrs)	Error (%)
8.0	28	105	0.65	221.7	-10.8	1.21E+07	0.4	29	-9.4	33,5	-8.7
0.8	28	50	7.85	174.7	-29.8	1.20E+07	0.3	28	-12.5	30.1	-18.1
0.8	28	80	4.96	195.3	-21.5	1.20E+07	0.3	29	-9.4	31.6	-13.8
0.8	23	80	4.96	215.4	-13.4	1.20E+07	0.2	27	-15.6	30.3	-17.5
0.8	23	105	0.65	245.9	-1.1	1.20E+07	0,3	28	-12.5	32.1	-12.6
0.8	23	107	0.65	248.1	-0.2	1.20E+07	0.1	32	0	36,3	-1.3

155



```
30-Mar-17
 FIT
time inc = 1 hr, calcs for 150 incs, 1 bursts
C 3 pluvios, non uniform pattern
1,150,1,3,1,-99,
0,21, rain from 0 - 21 time incs
Bremer River
2,8,4,1,6,13,28,13,1,2,0,1,1,12,19,29,21,15,31,4,8,-99, mm/hr
Tenthill
 1,5,10,8,6,4,7,6,3,0,2,1,3,3,3,5,3,4,1,3,3,-99,mm/hr
 UQ Gatton
 1.6,12.2,2.8,3.2,6.2,2.8,11.8,14,0.8,1,0.4,0.6,1.6,5.4,3,12,11.4,4.6,0,1.2,0.
 6,-99,mm/hr
C sub catchment rainfalls from isohyetal map 177.7,180.8,181.3,184,183.2,181.2,173.8,160.7,163.2,149.9,138.1,126.1,120.1,1 23.4,120.7,124.5,118.3,113.3,113.116.6,-99, Sub area event total rainfalls
C pluivo réference numbers
1,1,1,1,1,1,1,1,1,1,1,3,3,3,2,2,2,1,1,-99
 C Hydrograph data
 0,95,-99
 Warrego Highway
Warrego Highway
0.029,0.032,0.035,0.035,0.04,0.041,0.05,0.059,0.181,0.238,0.238,0.218,0.218,0.208,0.249,0.523,1.608,3.349,6.119,7.712,12.106,14.943,17.402,20.526,24.258,42.804,65.299,102.877,132.96,172.763,220.544,246.857,248.682,248.682,230.269,20.923,169.541,139.312,114.542,95.538,80.899,70.582,64.509,60.088,56.059,51.976,48.804,45.872,42.893,39.638,36.215,32.322,28.321,24.258,21.059,19.038,16.932,15.312,13.924,12.407,10.851,9.414,8.231,7.265,6.54,6.018,5.585,5.26,4.942,4.57,4.274,4.025,3.795,3.595,3.412,3.244,3.081,2.933,2.798,2.658,2.521,2.368,2.08,2.046,1.902,1.755,1.617,1.503,1.394,1.307,1.215,1.136,1.046,0.97,0.905,0.839,-99
 0.839.-99. m3/s
```



F6: February 2022 event

Calibration summary

Feb - Mar 22 Rainfall event characteristics

Start date & time 24/02/2022 10:00 End date & time 28/02/2022 9:00

Rainfall duration (hours) Number of bursts Runoff characteristics Hydrograph duration 146 hrs
1, minor recession peaks

1097.7 discharge at 143229A

(m3/s)

Supplementary peak discharge (m3/s)

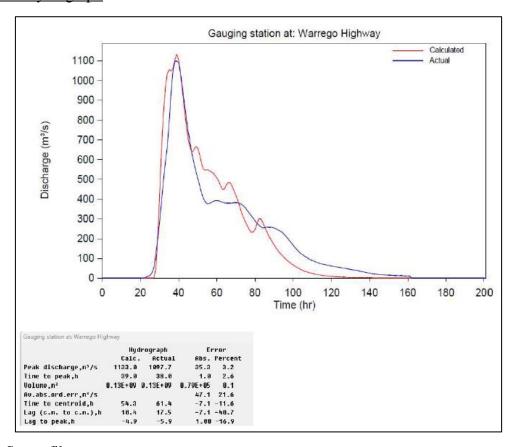
n/a

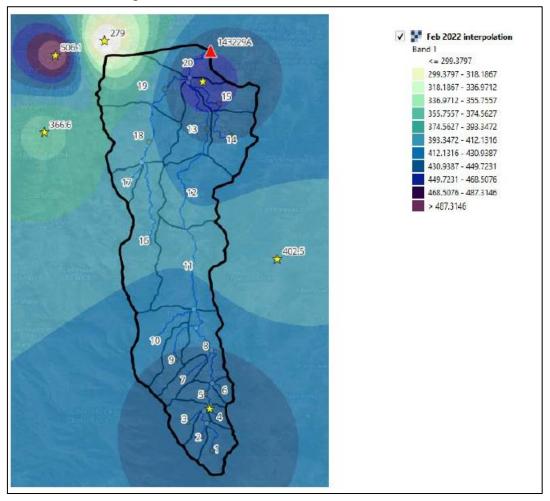
9	Total event rain	nfall record		
Station name	e Station ID Rec		Notes	
Forest Hill	040079	463		
UQ Gatton	040082	n/a		
Placid Hills	040449	506		
Gatton Allan St	040083	279		
Upper Tenthill	040388	366		
Franklyn Vale	040374	402.5		
Townson	040675	443		

Subcatchment centroid event rainfall depths from QGIS interpolation					
Subcatchment ID	Rainfall depth (mm)				
1	440.7				
2	442.1				
3	442.4				
4	443.5				
5	443.1				
6	442.2				
7	439				
8	429.4				
9	433,3				
10	425.7				
11	411.5				
12	413.2				
13	439.2				
14	433.6				
15	453				
16	408.7				
17	401.4				
18	408.8				
19	428.4				
20	453				

	RORB gauged data statistics								
Peak flow (m3/s)	Total flow volume (m3)	Time to peak (hrs)	Time to centroid (hrs)						
1097.7	1.30E+08	38	61.4						

	RORB calibrated parameters and statistics										
m	kc	IL	CL	Peak flow (m3/s)	Error (%)	Total flow volume (m3)	Error (%)	Time to peak (hrs)	Error (%)	Time to centroid (hrs)	Error (%)
0.8	29	140	0.07	1116.6	1.7	1.30E+08	0.1	39	2.6	54.5	-11.3
0.8	28	140	0.07	1140.8	3.9	1.30E+08	0.1	39	2.6	54.1	-11.9
0.8	28	142	0.03	1133	3.2	1.30E+08	0.1	39	2.6	54.3	-11.6
0.8	29	142	0.03	1108.3	1	1.30E+08	0.1	39	2.6	54.7	-11





F7: May 2022 event

Calibration summary

Event ID May-22 11/05/2022 6:00 Start Date End Date 14/05/2022 9:00 Duration (hours) Number of peaks 0:00 1

Maximum gauged discharge at 143229A

(m3/s) Supplementary peak discharge (m3/s) 521.093

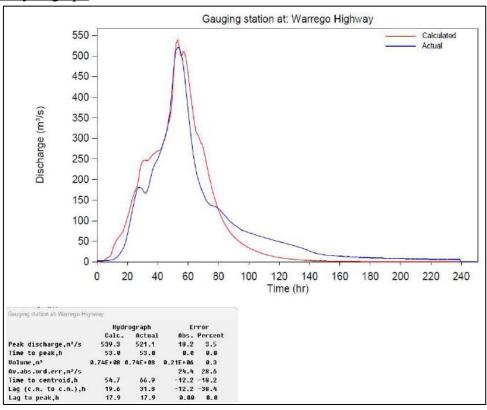
Station ID	Recorded rainfall (mm)	Notes	
040079	167.6		
040388	132		
040082	n/a	Offline	
040374	202	0.039/000	
040912	200.4		
040083	157		
040449	164.1		
040675	357.4		
WMIP	204		
	040079 040388 040082 040374 040912 040083 040449 040675	Station ID rainfall (mm) 040079 167.6 040388 132 040082 n/a 040374 202 040912 200.4 040083 157 040449 164.1 040675 357.4	

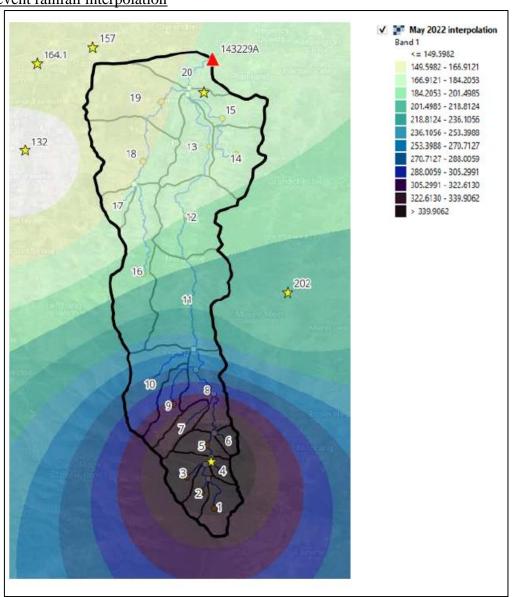
N/A

Subcatchment centroid event rainfall depths from QGIS interpolation					
Subcatchment ID	Rainfall depth (mm)				
1	343				
2	349.5				
3	350.7				
4	355.5				
5	354.2				
6	349.8				
7	335.9				
8	295.3				
9	311.9				
10	277.1				
11	214.1				
12	184.2				
13	170.9				
14	174.8				
15	168.7				
16	196.5				
17	168.1				
18	166.4				
19	166.1				
20	167.4				

RORB gauged data statistics								
Peak flow (m3/s)	Total flow volume (m3) Time to peak (hrs)		Time to centroid (hrs)					
521.1	7.40E+07	53	66.9					

RORB calibrated parameters and statistics											
m	kc	IL	CL	Peak flow (m3/s)	Error (%)	Total flow volume (m3)	Error (%)	Time to peak (hrs)	Error (%)	Time to centroid (hrs)	Error (%)
0.8	35	20	0.81	602	15.5	7.42E+07	0.3	52	-1.9	54.1	-19.1
0.8	38	10	0.96	561.5	7.8	7.42E+07	0.3	53	0	54.3	-18.8
8.0	40	5	1.03	539.3	3.5	7.42E+07	0.3	53	0	54.7	-18.2
8.0	40	0	1.09	532.1	2.1	7.42E+07	0.3	53	0	54.1	-19
8.0	41	0	1.09	525.4	8.0	7.42E+07	0.3	53	0	54.6	-18.3





F8: January 2024 event

Calibration summary

Station name	Station ID	Recorded rainfall (mm)	Notes
Forest HIL	040079	156.1	
Upper Tenthill	040388	151.8	
UQ Gatton	040082	163.8	
Franklyn Vale	040374	180	
Franklyn Vale Alert	040912	188.8	
Gatton Allan St	040083	138.7	
Placid Hills	040449	173.8	
Townson	040675	n/a	
Bremer River	WMIP	164	

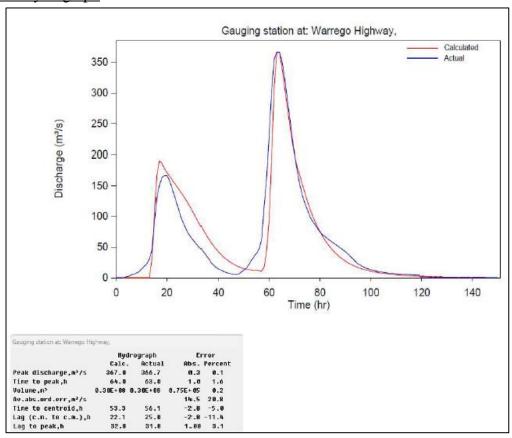
	id event rainfall depths interpolation
Subcatchment ID	Rainfall depth (mm)
1	174
2	174
3	173.2
4	175.1
5	175.2
6	176.2
7	174.4
8	177
9	174.3
10	172.5
11	175.6
12	168.8
13	168
14	166.8
15	164.8
16	168.5
17	162.3
18	161.5
19	162.3
20	164

	RORB gauged data	statistics	
Peak flow (m3/s)	Total flow volume (m3)	Time to peak (hrs)	Time to centroid (hrs)
366.7	3.00E+07	63	56.1

						Si	d statistics	parameters an	RORB calibrated			
	Error (%)	Time to centroid (hrs)	Error (%)	Time to peak (hrs)	Error (%)	Total flow volume (m3)	Empr (%)	Peak flow (m3/s)	CL	IL.	kc	m
	-24.4	42.5	9	54	0.3	3.01E+07	-0.5	364.8	5.85	36	27	0.8
	-21	44,3	-12.7	55	0.2	3.01E+07	0.1	367	5.52	39	28	0.8
12 hr tran	-8.4	51.4	0	63	0.3	3.01E+07	-0.5	364.8	5.85	36	27	0.8
12 hr tran	-5	53.3	1.6	64	0.2	3.01E+07	0.1	367	5.52	39	28	0.8

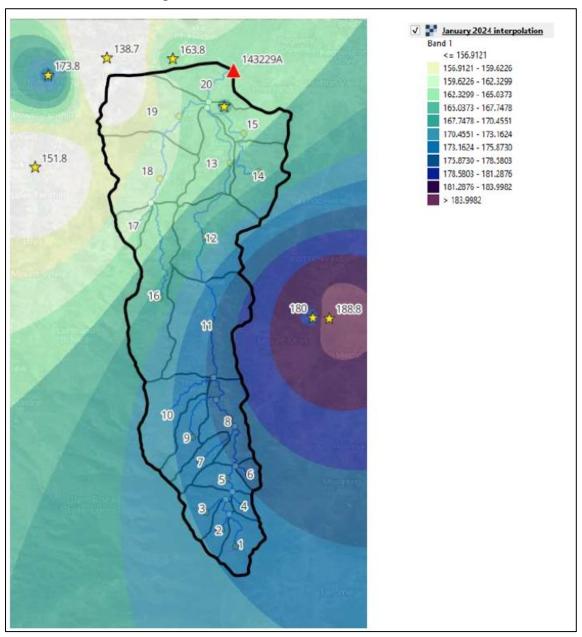
164

Calibrated hydrograph



RORB Storm file

QGIS event rainfall interpolation



Appendix G: Design discharge simulation sets

Summary of results

Current climate conditions

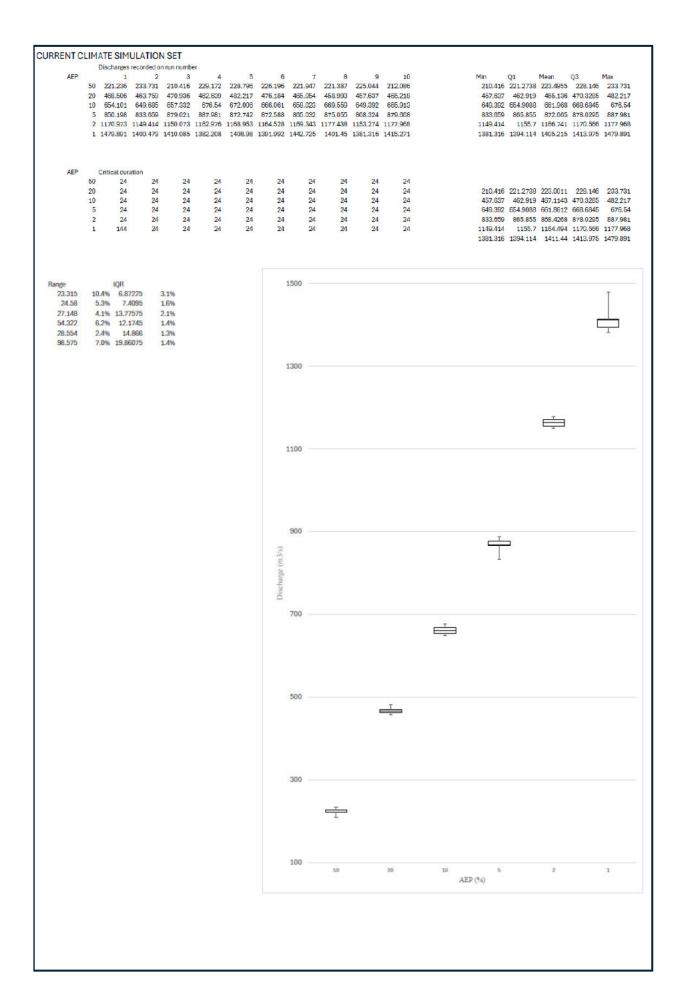
AEP	D	ischarge (m³	/s)
(%)	Median	Minimum	Maximum
50	223	210.4	233.7
20	467.1	457.6	482.2
10	661.9	649.4	676.5
5	868.4	833.7	888
2	1164.5	1149.4	1178
1	1411.4	1381.3	1479.9

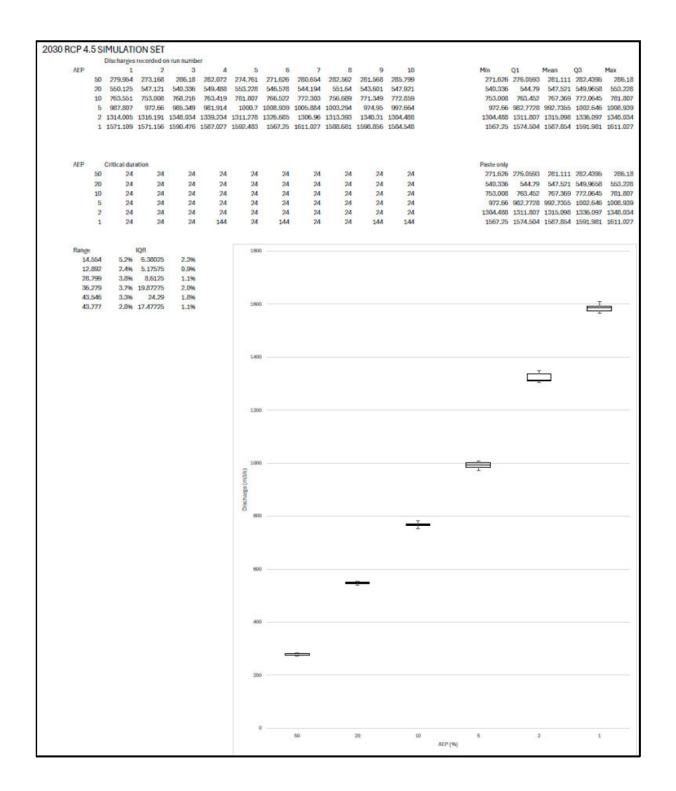
RCP 4.5 Scenarios

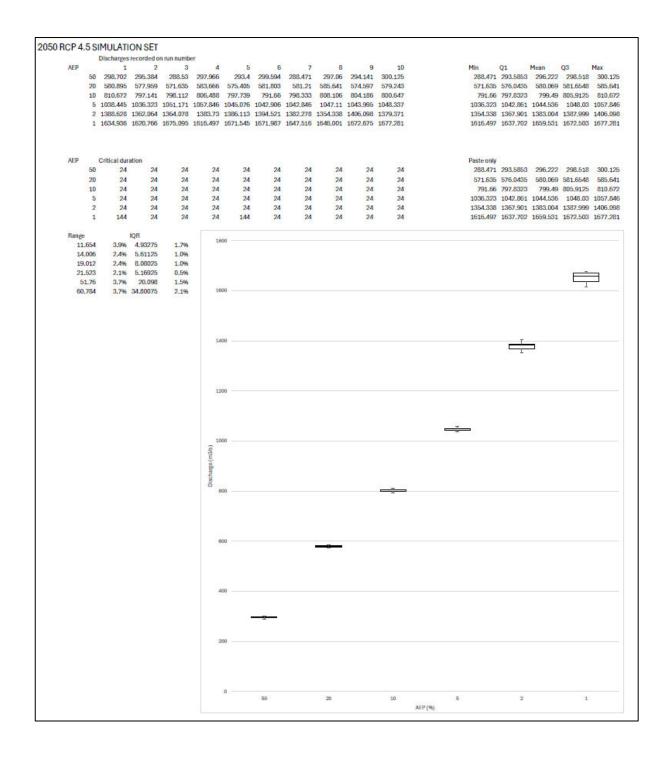
AEP (%)	Discharge 2030 (m³/s)	Increase (%)	Discharge 2050 (m³/s)	Increase (%)	Discharge 2090 (m³/s)	Increase (%)
50	281.1	26	296.2	33	330	48
20	547.5	17	580.1	24	630.4	35
10	767.4	16	799.5	21	865.8	31
5	992.7	14	1044.5	20	1116.7	29
2	1315.1	13	1383	19	1483.7	27
1	1587.9	12	1659.5	18	1772.3	26

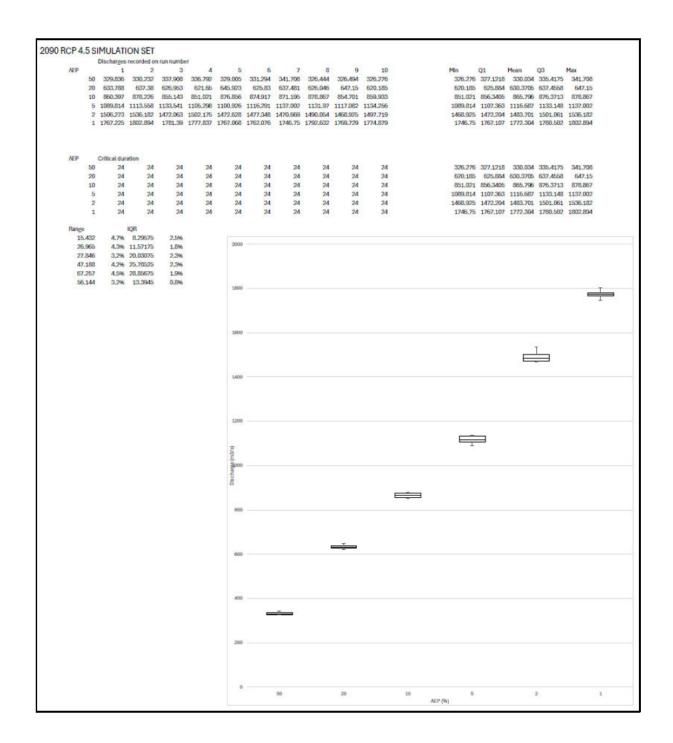
RCP 8.5 Scenarios

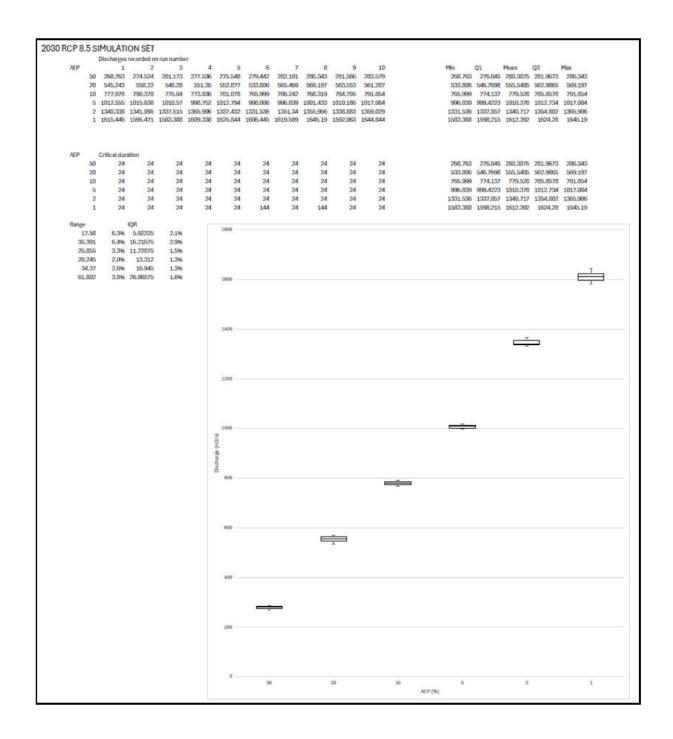
AEP (%)	Discharge 2030 (m³/s)	Increase (%)	Discharge 2050 (m³/s)	Increase (%)	Discharge 2090 (m³/s)	Increase (%)
50	280.3	26	320	43	431.5	94
20	555.5	19	614.9	32	780.9	67
10	779.5	18	842.5	27	1056.5	60
5	1010.4	16	1086.9	25	1341.2	54
2	1340.7	15	1446.6	24	1762.8	51
1	1612.4	14	1736	23	2096.3	49

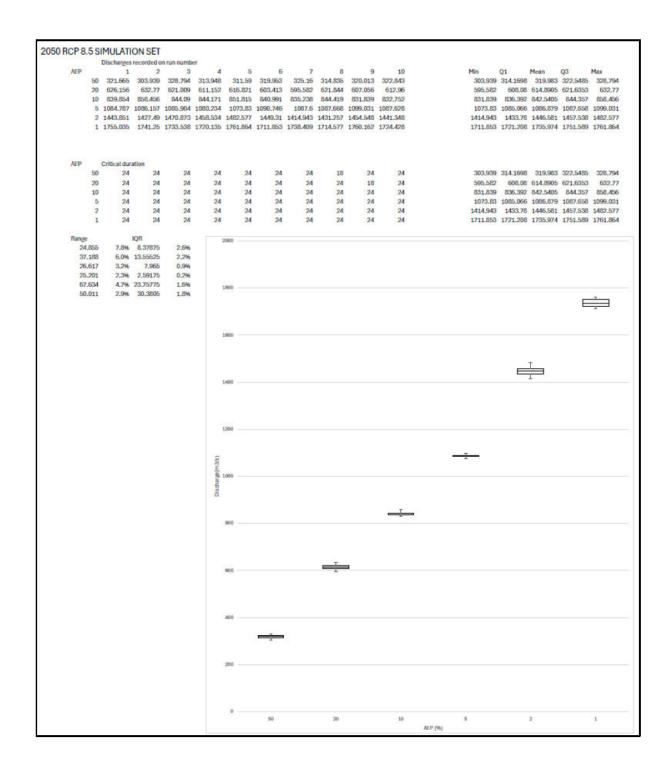


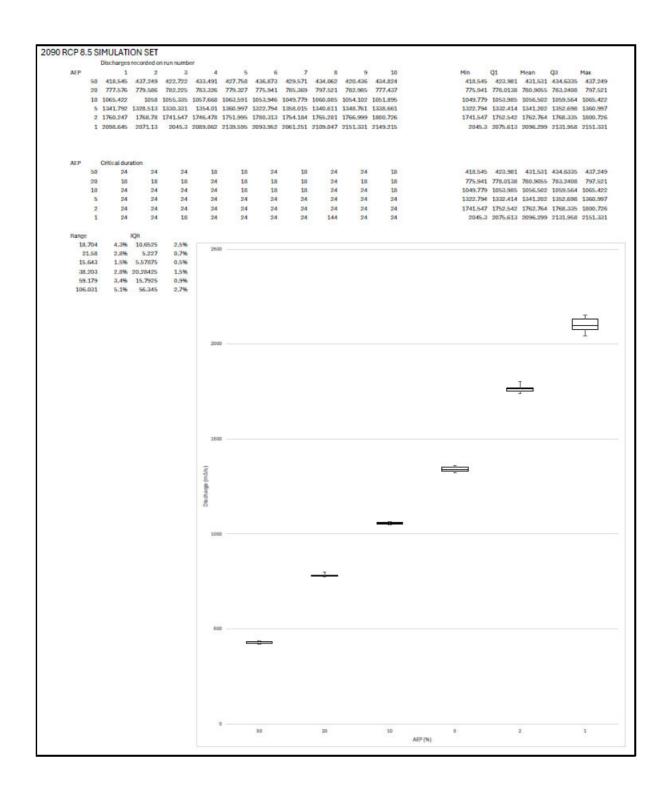


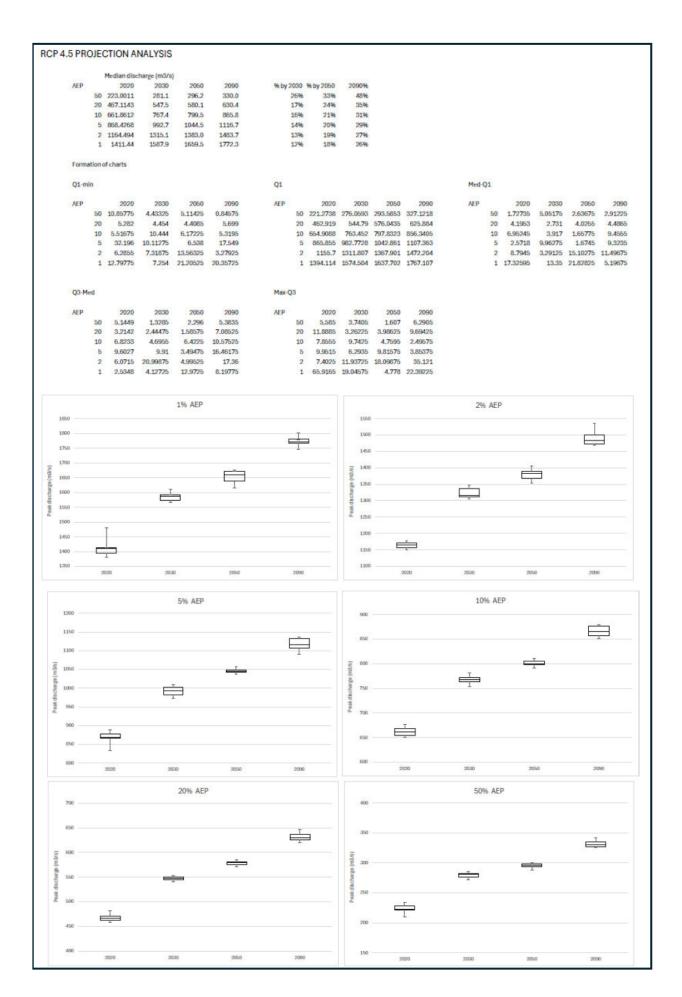


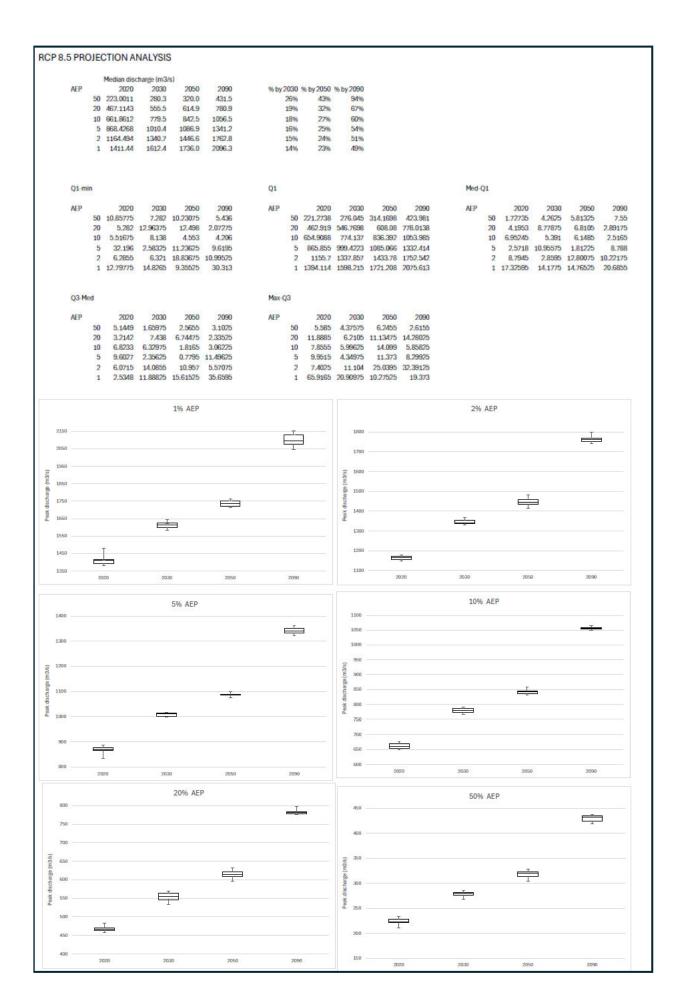


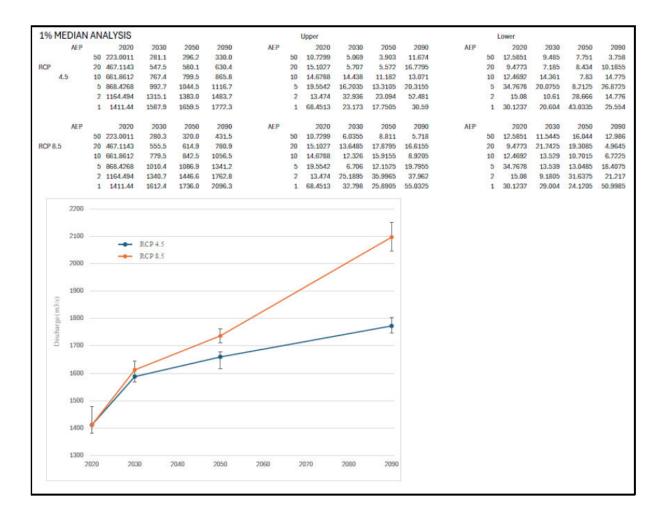












Appendix H: Sensitivity analysis of model parameters using 2020 outputs

	AEP		50	20	10	5	2	i i
	2020 M	ed	223.0011	467.1143	661.8612	868.4268	1164.494	1411.4
KC + 10%			Discharges	recorded o	n run numb	er		
29.7	AEP		1	2	3	4	5	
		50	207.646	201.794	210.813	207.298	206.731	
		20	429.236	433.287	431.989	425.173	439.986	
		10	609.539	615.113	604.845	611.945	632.865	
		5	795.006	795.82	800.967	799.57	796.256	
		2	1089.774	1072.127	1091.725	1077.384	1078.033	
		1	1319.796	1278.987	1302.257	1323.576	1325.423	
	AEP		Critical dur	ation				
		50	24	24	24	24	24	
		20	24	24	24	24	24	
		10	24	24	24	24	24	
		5	24	24	24	24	24	
		2	24	24	24	24	24	
		1	144	36	144	144	144	
			2	5	10	20	50	100
	Min		201.794	425.173	604.845	795.006	1072.127	1278.98
	Q1		206.731	429.236	609.539	795.82	1077.384	1302.25
	Med		207.298	431.989	611.945	796.256	1078.033	1319.79
	Q3		207.646	433.287	615.113	799.57	1089.774	1323.57
	Max		210.813	439.986	632.865	800.967	1091.725	1325.42
	% Med		-7.0%	-7.5%	-7.5%	-8.3%	-7.4%	-6.5%
KC - 10%			Discharges	recorded o	n run numb	er		
	AEP		1	2		4	5	
		50	248.204		238.181		Contraction of the Contraction o	
		20			517.537			
		10			712.198			
		5			920.191			
		2	1321.966					
			1489.297					
	AEP		Critical dur	ation				
		50	24	24	18	24	24	
		20	24	24	24	24	24	
		10	24	24	24	24	24	
		5	24	24	24	24	24	
		2	24	24	24	24	24	
		1	24	24	24	24	24	

			2	5	10	20	50	100
	Min		238.181	501.201	712.187		1219.809	1471.374
	Q1		241.67	511.331	712.198		1238.915	
	Med		247.846				1242.945	
	Q3						1253.216	
	Max		249.848				1321.966	
	Max		243.040	020.040	720.402	320.040	1321.300	1490.037
	% Med		11.1%	10.0%	7.9%	6.6%	6.7%	5.2%
IL+20%			Discharges	recorded o	n run numb	er		
30	AEP		1	2	3	4	5	
		50	187.822	205.803	181.439	202.17	174.577	
		20	438.508	443.731	446.777	435.953	442.622	
		10	616	630.757	648.222	627.861	638.67	
		5		823.341				
		2	1128.232			1144.134		
			1365.584					
		-						
	AEP		Critical dur	ation				
		50	24	24	24	24	24	
		20	24	24	24	24	24	
		10	24	24	24	24	24	
		5	24	24	24	24	24	
		2	24	24	24	24	24	
		1	144	24	24	24	24	
			2	5	10	20	50	100
	Min		174.577		616		1124.892	1365.428
	Q1			438.508			1128.232	
	Med		187.822	442.622	630.757		1130.449	
	Q3			443.731				1382.754
	Max		205.803	446.777	648.222	844.047	1162.218	1398.218
	% Med		-15.8%	-5.2%	-4.7%	-3.2%	-2.9%	-2.9%
IL - 20%			Discharges	recorded o	n run numb	er		
20	AEP		Discharges 1	2			5	
20	7121	50	256.727	_				
			477.754					
			675.883					
			878.994					
							1177.99	
			1424.556					
		-						
	AEP		Critical dur	ation				
		50	24	24	24	24	24	

		(7),100	1000	7.000	25,000	1000		
		20	18	24	24		24	
		10	24	24	24	24	24	
		5	24	24	24	24	24	
		2	24	24	24	24	24	
		1	24	24	24	24	144	
			2	5	10	20	50	100
	Min		242.078	477.754	663.704	868.96	1159.404	1414.927
	Q1		253.411	484.723	666.605	878.994	1166.783	1420.323
	Med			486.612				
	Q3		256,727	487.586	678,458	882,655	1177.99	1424.556
	Max			496.778				
					A. C.			The second second
	% Med		14.1%	4 2%	2 1%	1.3%	0.3%	0.8%
	701104		14.170	7.270	2.170	1.070	0.070	0.070
CI + 200%			Discharges	recorded o	n run numb	ar.		
1.68			Discharges 1	2		4	5	
1.00	ALI	En	201.511	and the second	DESCRIPTION OF THE PARTY OF THE	Secretary 200	Licenses - Cold	
			442.092					
			627.779					
			811.347					
			1120.388					
		1	1369.412	1368.971	1358.419	1357.874	1369.012	
			0					
	AEP		Critical dur			-	2000	
		50	24	24			24	
		20	24	24			24	
		10		24			24	
		5	24	24				
		2	24	24	24	24	24	
		1	24	24	24	144	24	
			2	5			50	
	Min		201.511		619.669			
	Q1		204.143	432.011	627.779	825.414	1122.786	1358.419
	Med		207.489	442.092	630.201	833.331	1123.612	1368.971
	Q3		208.242	442.721	641.15	838.689	1128.164	1369.012
	Max		209.44	448.996	645.251	841.085	1129.2	1369.412
	% Med		-7.0%	-5.4%	-4.8%	-4.0%	-3.5%	-3.0%
CL-20%			Discharges	recorded o	n run numb	er		
1.12	AEP		1	2		4	5	
0.30(2)/60	to one	50	C PROPERTY AND ADDRESS OF	259.237		and the second second	remarks and the	
		20		501.438				
		10		690.326				
1		10	004.330	030,020	700.123	033.203	032.000	

	1	1435.125	1432.342	1427.53	1409.079	1452.443	
AEP		Critical dura	ation				
	50	24	24	24	24	24	
	20	24	24	24	24	24	
	10	24	24	24	24	24	
	5	24	24	24	24	24	
	2	24	24	24	24	24	
	1	24	24	24	144	24	
		2	5	10	20	50	100
Min		233.374	487.583	664.993	901.453	1175.215	1409.079
Q1		243.009	493.858	690.326	902.064	1179.473	1427.53
Med		246.195	501.438	692.585	903.799	1193.772	1432.342
Q3		251.854	503.088	699.209	909.428	1201.824	1435.125
Max		259.237	503.779	703.129	918.103	1217.703	1452.443

Appendix I: Flood frequency analysis – RMC outputs

nmary statist	ics		
Measure	Unit	All data	Low outliers excluded
Record Length	years	33	26
Low Outliers	number	7	7
Minimum	m3/s	0	7.41
Maximum	m3/s	1,387.11	1,387.11
Mean	m3/s	182.86	225.96
Std Dev	m3/s	343.08	370
Skewness		2.5122	2.2935
Kurtosis		8.2562	4.3783
Mean (of log)		1.023	1.8928
Std Dev (of log)		1.8592	0.6591
Skewness (of		-1.2953	0.232
Kurtosis (of log)		3,4322	-0.6233
5%	m3/s	0	7.77
25%	m3/s	7.62	22.88
	m3/s		
50%		53.5	67.54
75%	m3/s	182.1	193.19
95%	m3/s	1,094.35	1,083.66
Annual Maximun			0.00
	charge (m:	Plotting position	Is low outlier
100000000000000000000000000000000000000	127.376	0.2892	FALSE
1992	192.208	0.2289	FALSE
1993	0.567	0.8313	TRUE
1994	0	0.8916	TRUE
1995	7.406	0.7711	FALSE
1996	496.592	0.1084	FALSE
1997	18.882	0.5904	FALSE
1998	0	0.9217	TRUE
1701000			
1999		0.3193	FALSE
2000	15.564	0.6807	FALSE
2001		0.1988	FALSE
2002	0	0.8614	TRUE
2003	7.709	0.7410	FALSE
2004	75.116	0.3494	FALSE
2005	55.205	0.4699	FALSE
2006	17.919	0.6205	FALSE
2007	0	0.9819	TRUE
	53.504	0.5000	FALSE
10.000	255.854	0.1386	FALSE
2.74		0.5301	
2010			FALSE
	1387.11	0.0181	FALSE
2012	63.378	0.4096	FALSE
2013	1041.5	0.0783	FALSE
2014	59.687	0.4398	FALSE
2015	15.68	0.6506	FALSE
2016	0.891	0.8012	TRUE
0.000	248.682	0.1687	FALSE
1		0.7108	FALSE
ZU181			
2018	0	0.9518	INUE
2019		0.9518	TRUE
2019 2020	34.855	0.5602	FALSE
2019 2020 2021	34.855 177.775	0.5602 0.2590	FALSE FALSE
2019 2020 2021	34.855 177.775 1097.72	0.5602	FALSE

Measure	Exp.	Gamma	GEV	GL	GP	EVI	In-Norm	Logistic	Log-Norm	LP3	Norm	P3	Weibu
Location	-6.955	N/A	19,4518	40.8376	-12.1712	27.0431	393.1046	81.3885	1.5213	1.3769	119.0687	180.0819	N/A
Scale	193.9176	601.6235	48.3992	53.8431	57.3812	201.1046		180.7966	0.9655	1.3011	406.6537	240.8627	89.61
Shape	N/A	0.3115	-1.1258	-1.1281	-0.962	N/A	N/A	N/A	N/A	-1.32	N/A	2	0.475
Minimum	-6.96	0	-23.54	-6.89	-12.17	-00	0	-00	0	0	-00	-60.78	0
Maximum	00	00	00	00	00	00	00	00	00	2,230,07	00	00	00
Mean	186.96	187.44	N/A	N/A	1,498.53	143.12	393.1	81.39	97.14	182.2	119.07	180.08	197.9
Std Dev	193.92	335.81	N/A	N/A	N/A	257.93	4,636.17	327.93	266.98	304.98	406.65	240.86	476.2
Skewness	2	3.5832	N/A	N/A	N/A	1.1396	1,675.80	0	29.008	2.5842	0	2	7.416
Kurtosis	9	22.2586	N/A	N/A	N/A	5.4	*******	4.2	6,822.84	10.4196	3	9	112.45
AEP	Discharge (m3/s)												
0.001	1,267.65	2,479.65	93,424.96	105,911.07	49,172.71	1,290.47	17,405.19	1,224.51	17,405.52	4,009.67	1,335.97	1,464.99	4,017.
0.002	1,139.82	2,162.72	42,522.57	48,125.26	24,452.42	1,164.11	11,417.36	1,109.16	11,417.57	3,322.33	1,252.59	1,313.50	3,267.
0.005	970.83	1,751.10	15,001.49	16,930.62	9,692.82	996.91	6,259.16	956.4	6,259.27	2,473.37	1,133.74	1,113.25	2,392.5
0.01	843	1,446.91	6,805.21	7,657.16	4,798.85	870.19	3,811.54	840.38	3,811.60	1,890.71	1,035.66	961.76	1,819.
0.02	715.17	1,151.02	3,073.49	3,442.50	2,362.56	742.99	2,216.79	723.52	2,216.82	1,371.83	928.49	810.27	1,323.
0.05	546.18	777.82	1,057.81	1,172	907.95	573.26	983.25	566.09	983.26	801.45	767.74	610.02	785.6
0.1	418.35	515.83	458.76	500.41	425.63	442.13	477.49	441.93	477.5	468.34	624.92	458.53	469.9
0.2	290.52	282.17	186.87	198.1	185.52	305.43	199.12	307.17	199.12	225.71	451.97	307.04	233.4
0.3	215.74	165.99	103.33	106.59	105.78	219.99	105.98	217.61	105.98	126.01	327.26	218.43	132.4
0.5	121.54	54.19	40.36	39.64	42.16	98.95	37.36	76.81	37.36	43.13	121.1	106.79	45.04
0.7	59.48	11.73	14.31	14.08	14.98	-1.63	13.17	-63.98	13.17	12.81	-85.05	33.25	12.3
0.8	34.86	3.61	5.92	6.85	6.49	-54.5	7.01	-153.55	7.01	5.73	-209.76	4.07	4.92
0.9	13.13	0.49	-1.27	1.69	-0.1	-119.75	2.92	-288.3	2.92	1.71	-382.71	-21.67	1.14
0.95	3.16	0.07	-4.97	-0.25	-2.87	-167.68	1.42	-412.47	1.42	0.58	-525.54	-33.49	0.28
0.98	-2.57	0	-7.75	-1.21	-4.4	-216.3	0.63	-569.89	0.63	0.16	-686.29	-40.28	0.05
0.99	-4.44	0	-9.08	-1.49	-4.89	-246.01	0.37	-686.76	0.37	0.06	-793.45	-42.5	0.01

RMC BAYESIAN POSTERIOR MODE - LP3 Discharge (m3/s) AEP Predictive 97.5% CI 2.5% CI Mode 0.001 5240689.83 1488.16 109219.71 4194.77 0.002 1668918.40 1424.46 38659.31 3447.99 0.005 1272.14 11976.74 2539.62 357245.13 106610.37 1072.56 5152.30 1925.76 0.01 2735.94 1386.27 0.02 32580.86 816.35 0.05 6874.96 471.30 1253.12 802.04 0.1 2033.24 256.41 616.71 465.77 0.2 624.97 108.70 270.52 223.55 0.3 310.27 55.50 133.05 124.77 93.38 39.81 42.91 0.5 15.61 0.7 27.56 2.55 11.37 12.90 13.69 4.92 5.83 8.0 0.58 0.9 6.08 0.05 1.44 1.77 0.95 3.83 0.00 0.39 0.61 0.98 2.68 0.00 0.06 0.17 0.99 2.23 0.00 0.01 0.07 ... 0 100,000 1,000 Annual Exceedance Probability

